

A Ground-Up Data-Driven Approach to Distinguishing Magnetospheric Sources of Geomagnetically Induced Currents > 10 A during the 17 March 2013 Event

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Abstract

We combine wavelet analysis and data fusion to investigate geomagnetically induced currents (GICs) on the Mäntsälä pipeline and the associated horizontal geomagnetic field, BH, variations during the late main phase of the 17 March 2013 geomagnetic storm. The wavelet analysis decomposes the GIC and BH signals at increasing ‘scales’ to show distinct multi-minute spectral features around the GIC spikes. Four GIC spikes > 10 A occurred while the pipeline was in the dusk sector – the first sine-wave-like spike at ~ 16 UT was ‘compound.’ It was followed by three ‘self-similar’ spikes two hours later. The contemporaneous multi-resolution observations from ground-(magnetometer, SuperMAG, SuperDARN), and space-based (AMPERE, TWINS) platforms capture multi-scale activity to reveal two magnetospheric modes causing the spikes. The GIC at ~ 16 UT occurred in two parts with the negative spike associated with a transient sub-auroral eastward electrojet that closed a developing partial ring current (PRC) loop, whereas the positive spike developed with the arrival of the associated mesoscale flow-channel in the auroral zone. The three spikes between 18-19 UT were due to bursty bulk flows (BBFs). We attribute all spikes to flow-channel injections (substorms) of varying scales. We use previously published MHD simulations of the event to substantiate our conclusions, given the dearth of timely in-situ satellite observations. Our results show that multi-scale magnetosphere-ionosphere activity that drives GICs can be understood using multi-resolution analysis. This new framework of combining wavelet analysis with multi-platform observations opens a research avenue for GIC investigations and other space weather impacts.

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1 **A Ground-Up Data-Driven Approach to Distinguishing Magnetospheric Sources of Geomagnetically**
2 **Induced Currents > 10 A during the 17 March 2013 Event**

3
4 **Bhagyashree Waghule¹, D.J. Knipp¹, J. L. Gannon², D. Billet³, S.K. Vines⁴, J. Goldstein⁵**

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11 **Key Points:**

- 12 • Wavelet analysis of GICs at Mäntsälä on 17 March 2013 reveals two features – Pi1/Pi2
13 pulsations superposed on longer duration disturbances.
- 14 • Wavelet decomposition of the GIC and BH signals is consistent with multi-scale
15 magnetosphere-ionosphere activity around GIC spikes.
- 16 • Pi2 pulsations and data fusion suggest mesoscale flow channels (substorm injections)
17 were the underlying cause of four GICs > 10 A.

18

19 **Abstract**

20 We combine wavelet analysis and data fusion to investigate geomagnetically induced currents
21 (GICs) on the Mäntsälä pipeline and the associated horizontal geomagnetic field, BH, variations
22 during the late main phase of the 17 March 2013 geomagnetic storm. The wavelet analysis
23 decomposes the GIC and BH signals at increasing ‘scales’ to show distinct multi-minute spectral
24 features around the GIC spikes. Four GIC spikes > 10 A occurred while the pipeline was in the
25 dusk sector – the first sine-wave-like spike at ~ 16 UT was ‘compound.’ It was followed by three
26 ‘self-similar’ spikes two hours later. The contemporaneous multi-resolution observations from
27 ground-(magnetometer, SuperMAG, SuperDARN), and space-based (AMPERE, TWINS)
28 platforms capture multi-scale activity to reveal two magnetospheric modes causing the spikes.
29 The GIC at ~ 16 UT occurred in two parts with the negative spike associated with a transient sub-
30 auroral eastward electrojet that closed a developing partial ring current (PRC) loop, whereas the
31 positive spike developed with the arrival of the associated mesoscale flow-channel in the auroral
32 zone. The three spikes between 18-19 UT were due to bursty bulk flows (BBFs). We attribute all
33 spikes to flow-channel injections (substorms) of varying scales. We use previously published
34 MHD simulations of the event to substantiate our conclusions, given the dearth of timely in-situ
35 satellite observations. Our results show that multi-scale magnetosphere-ionosphere activity that
36 drives GICs can be understood using multi-resolution analysis. This new framework of
37 combining wavelet analysis with multi-platform observations opens a research avenue for GIC
38 investigations and other space weather impacts.

39

40 **Plain Language Summary**

41 Geomagnetically Induced Currents (GIC) are produced by complex interaction between the
42 Earth's magnetic field and ground composition during intense geomagnetic storms. These two
43 parameters are often related in frequency domain. In this paper, we analyze the GIC signal from
44 the Finnish natural gas pipeline recorded at Mäntsälä during the 17 March 2013 geomagnetic
45 storm. Four spikes > 10 Ampere were recorded between 4:30 – 9:00 PM local time. We use
46 wavelet analysis to learn about the frequencies of GIC spikes and then systematically investigate
47 the observations from ground to space (ground-up approach) to learn what links activity in space
48 to the GICs. Wavelet analysis highlights areas ranging from <1 minute to > 30 minutes, which
49 indicates that higher frequency fluctuations are accompanied with longer duration disturbance.
50 Multi-platform observations help us interpret the physical meaning of the multi-minute (or multi-
51 scale) area in the wavelet plot. We find that multi-scale activity in the magnetosphere and
52 ionosphere, created by fast earthward- flowing particles (magnetotail mesoscale plasma flows),
53 ultimately drove the significant GIC spikes. This new perspective enabled us to link the
54 magnetospheric activity to GICs through observations and previously published simulations and
55 pave a path for future research.

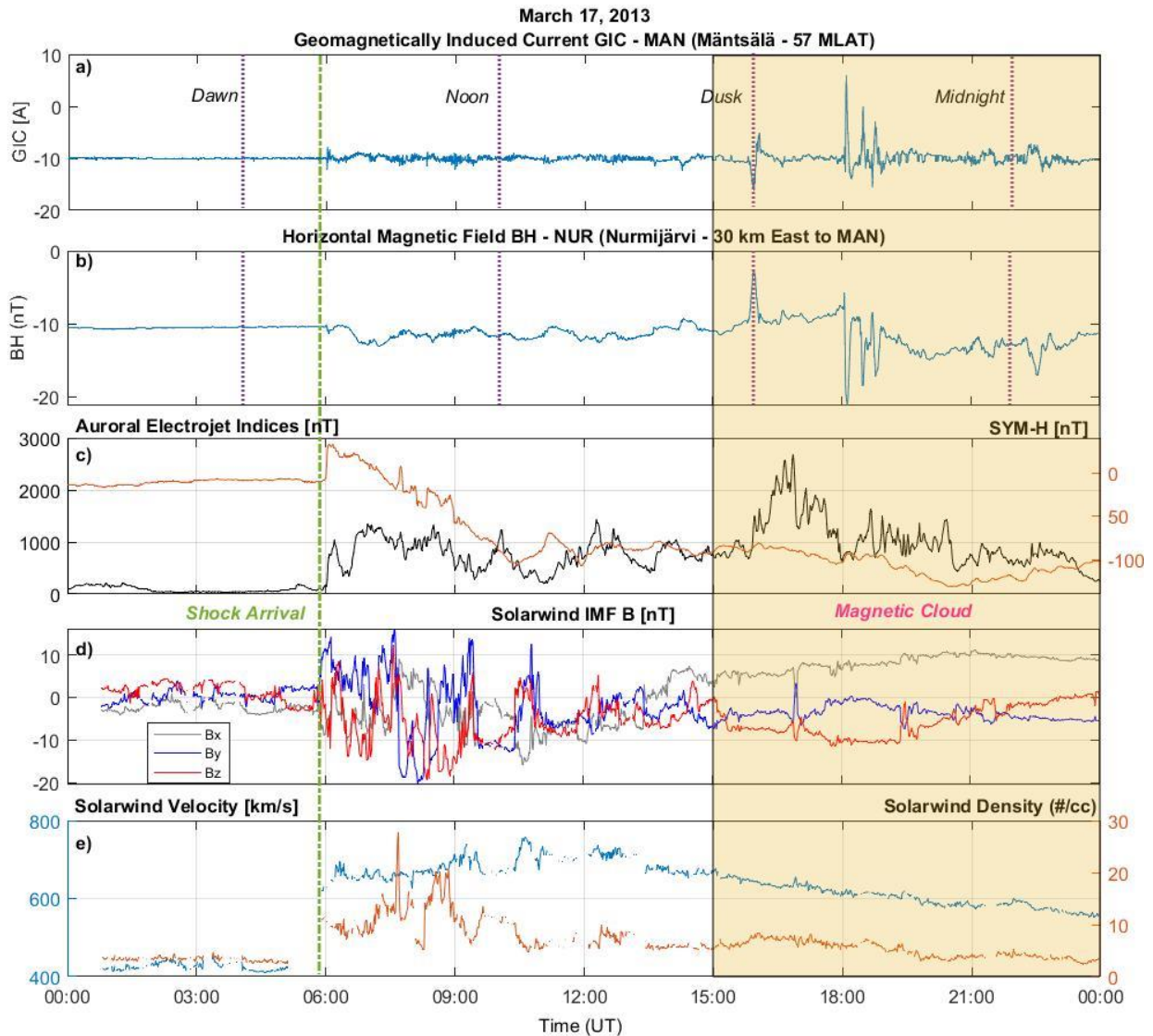
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57 **1. Introduction**

58 Geomagnetically induced currents (GICs) flow near the Earth's surface because of
59 induced geoelectric field. The geoelectric field is related to the magnetic field and ground
60 conductivity in the frequency domain such that the product of frequency spectrum of the
61 northward (southward) magnetic field (**B**) component and the transfer function of the Earth
62 produce the frequency spectrum of the eastward (westward) geoelectric field which drives GICs
63 (Boteler, 1994). Eventually, GICs find a path to close the circuit through long conductive
64 systems (>1km) such as power lines, pipelines, and communication cables, which pose a
65 significant risk to technological infrastructure (Pulkkinen et al., 2017). Hence, understanding the
66 drivers of GICs at different timescales is essential for accurate predictions. Tsurutani & Hajra
67 (2021) surveyed the solar wind conditions for GICs > 10 A in the Mäntsälä pipeline and
68 recommended a deeper investigation into the related near-Earth interactions. Herein, we employ
69 wavelet analysis to decompose the GIC and the horizontal **B** component (BH) time-series at
70 different scales (frequencies). The combined information from wavelet analysis and data fusion
71 of multi-resolution ground and space-based observations is used to explore the magnetospheric
72 source(s) of four GIC spikes (>10 A) recorded at Mäntsälä station of the Finnish natural gas
73 pipeline network, during the CME passage of the 17 March 2013 geomagnetic storm. Although
74 not among the extreme events studied by Juusola et al. (2023), this storm is interesting because
75 wavelet analysis suggests distinct GIC responses to different drivers. We hypothesize that the
76 first GIC spike had a compound source primarily associated with the interplay between partial
77 ring current (PRC), plasmopause, and substorm injection (mesoscale plasma flows), while the
78 other spikes are associated with bursty bulk flow (BBF).

79 Our analysis is informed by prior GIC studies and recent modeling efforts for the event.
80 W.-H. Xu et al. (2022) established the utility of wavelet analysis as a tool for analyzing the 17

81 March 2013 Mäntsälä GICs, relating them to the rate of change of the x-component geomagnetic
82 field (dBx/dt) signal at a reference ground magnetometer at Nurmijärvi, but did not address their
83 causes. Belakhovsky et al. (2019) noted very intense ionospheric vortex-driven GICs in
84 transformers in the Kola Peninsula, north of Mäntsälä during the storm. These spikes were
85 attributed to large-amplitude magnetic pulses that appeared to be part of a nightside substorm
86 current wedge (SCW). Despirak et al. (2022) also studied disturbances on the Karelian-Kola
87 power transmission line for the same date, finding that GICs corresponded to the appearance of
88 successive substorm intensifications.



89

90 **Figure 1:** a) GIC at Mäntsälä; b) BH at Nurmijärvi ~ 30 km east of Mäntsälä; c) Global
 91 Auroral Electrojet index and SYM-H index; d) IMF components; e) Solar wind density and
 92 velocity. Pink lines denote (Magnetic) local time at Mäntsälä (approximately UT+2h). Green
 93 line indicates shock arrival, and the highlighted yellow region indicates interval of magnetic
 94 cloud.

95 Figure 1 provides a ground-to-space view of activity during the 2013 St Patrick's Day
 96 storm. A sudden change in GIC and BH (Fig. 1a-b) and global auroral and RC activity (Fig. 1c)
 97 occurred at ~06 UT, simultaneous with the arrival of a solar wind shock indicated by an increase
 98 in interplanetary magnetic field (IMF) fluctuations (Fig. 1d) and increased solar wind speed and

99 density (Fig. 1e). During the post-shock interval (06-15 UT) an intense, $Dst < 100$ nT,
100 geomagnetic storm developed (Fig. 1c) with significant increases in substorm activity and
101 magnetosphere-ionosphere (M-I) coupling (e.g., Lyons et al., 2016). Despite the post-shock IMF
102 and solar wind variations, the Mäntsälä pipeline experienced only small GICs while it transited
103 the dayside. Verkhoglyadova et al. (2016) and Wu et al. (2016) noted that the IMF stabilized as
104 the leading edge of a magnetic cloud (MC) arrived at $\sim 15:30$ UT, after which the IMF was
105 southward. Four GIC spikes arose while the pipeline moved through the duskside (Fig. 1a). A
106 sinusoid-shaped spike at ~ 16 UT was associated with a sharp increase in BH (Fig.1b) while the
107 other three GIC spikes between 18-19 UT were associated with sharp dips in BH. This interval in
108 the late main phase/recovery phase of the storm has been well studied in context of nightside
109 activity (e.g., Gkioulidou et al., 2014 and Yu et al., 2014)

110 The understanding of magnetospheric dynamics has improved over time through MHD
111 simulations such as those applied in Multiscale Atmosphere-Geospace Environment (MAGE)
112 efforts. Wiltberger et al. (2017) showed an increased plasma pressure in the duskside
113 magnetosphere within $4R_E$. Sorathia et al. (2018), showed peak electron injection at 16 UT
114 followed by ion injections in the recovery phase of the storm. Later, Sorathia et al. (2023)
115 identified plasmashet mesoscale bubbles that penetrate the inner magnetosphere, as an early
116 multi-scale source of the auroral electrojet and dB/dt variations on the ground. Their findings are
117 supplemented by Sciola et al. (2023) who show that bubbles are responsible for at least 50% of
118 the plasma energy enhancement within $6 R_E$ during this strong geomagnetic storm. We note that
119 advances in simulations have been instrumental in associating terms like mesoscale plasma flow,
120 ionization channels, plasma bubbles and bursty bulk flows with substorm and RC injection.

121 Our motivating question is - How would wavelet analysis help with understanding the
122 magnetospheric drivers of the GIC spikes? The frequency of magnetic field perturbations is
123 important for generating GICs because longer-period fluctuations penetrate more deeply into the
124 Earth, whereas the short-period fluctuations remain closer to the surface (Gannon et al., 2017).
125 Hence, compared to a 1-D time series, a 2-D time-frequency analysis (such as wavelet transform)
126 aids in understanding the distribution of fluctuations in the GIC signal and the associated BH
127 fluctuations.

128 Wavelet analysis has been used in climate studies to understand periodic behavior (Yiou
129 et al., 1996; Torrence & Compo, 1998 and references therein), and to understand geophysical
130 time-series (Grinsted et al., 2004). Pulkkinen & Kataoka (2006) used the S-transform method to
131 study the properties of GIC fluctuations in the Finnish natural gas pipeline. Later, Z. Xu (2011)
132 reported that wavelet analysis could distinguish geomagnetic effects produced from various
133 currents in the magnetosphere and the ionosphere, in terms of frequency variations. In the
134 subsequent years, Falayi et al., (2017) explored the spectral characteristics of GICs using
135 continuous wavelet transform (CWT) during several geomagnetic storms; Adhikari et al. (2017)
136 found a positive correlation between GIC, auroral, and RC activities during geomagnetic storms;
137 Khanal et al. (2019) found that long-duration high-intensity substorm activity drives continuous
138 small-amplitude fluctuation in GIC over several days, the cumulative effect of which is
139 important for pipeline corrosion; Orr et al. (2021) analyzed the network response of GICs in the
140 United Kingdom using wavelet transform and found a correlation to auroral electrojets.

141 We add to the literature by using the property of scales from wavelet analysis to learn
142 about the underlying frequencies during peak GICs and data fusion of ground and space-based
143 observations, gathered at different resolutions, for physical interpretation of the CWT results. We

144 adopt the idea of data fusion coined by Hall & Llinas (1997) who defined it as combining data
145 from multiple sensors and related information from associated databases to achieve improved
146 accuracy and more specific inferences than could be achieved using a single sensor alone. This
147 not only helps us overcome the challenge of using datasets from different sources with non-
148 uniform sampling periods but also allows us to analyze them together for an integrated picture.
149 In this paper, we show that wavelet analysis of ground data (GIC and BH) supplemented with
150 systematic fusion of observations and prior modeling, reveals GICs as a natural consequence of
151 multi-scale ionospheric activity driven by magnetosphere dynamics.

152 **2 Data and Method**

153 **Ground-based data:** We use the 10-s GIC data measured in the Finnish natural gas
154 pipeline at Mäntsälä (MAN, 60.6N GLAT / 57 MLAT) on 17 March 2013. The corresponding **B**
155 is measured by the reference magnetometer at Nurmijärvi (NUR, 60.5N GLAT / 57 MLAT) 30
156 km east of MAN (Pulkkinen, Viljanen, et al., 2001; Viljanen et al., 2006). We 10-s resolution
157 International Monitor for Auroral Geomagnetic Effects (IMAGE) magnetometer-derived 2D
158 Equivalent Currents (EC) (Tanskanen, 2009), 1-minute SuperMAG products (Newell &
159 Gjerloev, 2011, 2012; Waters et al., 2015), and 2-minute Super Dual Auroral Radar Network
160 (SuperDARN) products (Greenwald et al., 1995) for assessing the ionospheric activity.

161 It is important to note that MAN sits at the central junction the pipeline that spans 350 km
162 in the east-west direction and about 120 km in the north-south direction. Although, Pulkkinen,
163 Pirjola, et al. (2001) note that the Finnish pipeline is electrically connected to the Russian
164 pipeline, the GICs induced far in the Russian side of the network does not reach very far to the
165 Finnish side. They also note that although the GIC is measured at a single location (MAN), it is

166 influenced by activity over the entire pipeline. In contrast, NUR is not a part of an electrically
167 connected network and hence will have slight differences in the signal fluctuations.

168 **Space-based data:** We use Active Magnetosphere and Planetary Electrodynamic
169 Response Experiment (AMPERE) fitted Field Aligned Current (FAC) data (10-minute averaged)
170 and plots (Anderson et al., 2002; Waters et al., 2001) for assessing the M-I coupling. Energetic
171 Neutral Atom (ENA) 15-minute averaged images from Two Wide-Angle Imaging Neutral-atom
172 Spectrometers (TWINS) instrument provide information about the RC population distribution
173 (McComas et al., 2009). The solar wind observations, Sym-H, and AE indices are from (King &
174 Papitashvili, 2005).

175 For TWINS images, we performed a sensor-dependent flat-field and background
176 subtraction for each image following McComas et al. (2012) a flat-field image was produced:
177 i.e., any flux mismatch at the boundaries between adjacent TWINS sensor heads was removed by
178 subtracting the excess from the higher-flux side. Then, the minimum measured flux along the
179 outermost edge of the flat-field image field-of-view (FOV) was subtracted from the entire image.

180 **Method: Wavelet Analysis + Information fusion**

181 We analyze the 1-D time-series of GIC at MAN and $BH = \sqrt{B_X^2 + B_Y^2}$, at NUR using wavelet
182 analysis techniques. We favor the use of frequency analysis of BH rather than dB_X/dt in our
183 study to capture all underlying frequencies from all directions, which can be important for
184 scientific investigations (Heyns et al., 2021; Watari et al., 2009).

185 **2.1 Continuous Wavelet Transform (CWT)**

186 Given the continuous nature of the signal, and the relationship of the magnetic field to
187 GIC in the frequency domain, CWT emerges as the optimal tool for analyzing this event.

188 Importantly, CWT exhibits high resistance to signal noise (Slavič et al., 2003), making it
 189 particularly advantageous for the analysis of potentially noisy GIC data (Pulkkinen, Viljanen, et
 190 al., 2001). CWT is a convolution of the input data sequence (GIC and BH) with a set of functions
 191 generated by the mother wavelet. We employ the ‘Morlet’ mother wavelet, based on its good
 192 time-frequency localization capabilities (Grinsted et al., 2004; Khanal et al., 2019; Torrence &
 193 Compo, 1998; W.-H. Xu et al., 2022).

194 CWT coefficients $C(a,b)$ are calculated using Eq. 1 such that the continuous wavelet
 195 function, $\psi(t)$ is shifted using the position parameter ‘b,’ across the time-series $x(t)$ with
 196 changing scale factor ‘a’ which stretches the wavelet to provide a 2-D representation of time-
 197 localized oscillations of a 1-D signal.

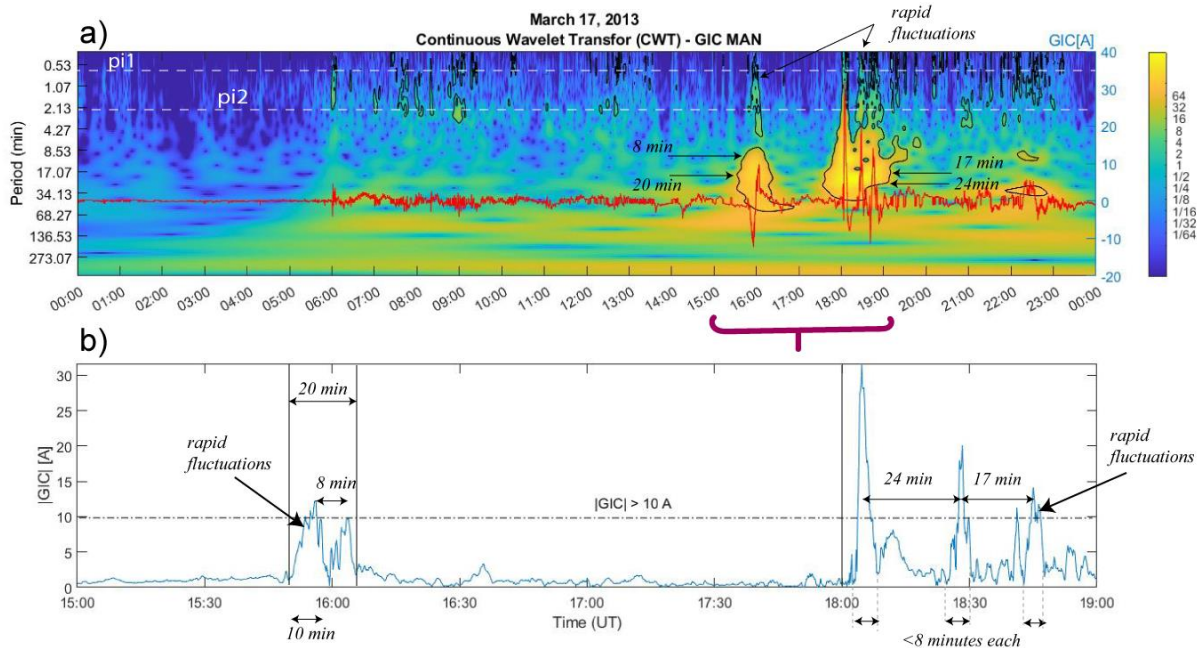
198
$$C(a, b) = \frac{1}{\sqrt{(a)}} \int_{-\infty}^{\infty} x(t) \psi \left(t - \frac{b}{a} \right) dt \dots\dots\dots (1)$$

199 The heatmap in Fig. 2a shows the power spectrum of the CWT of GIC, calculated using Eq. 2,

200
$$\text{Wavelet power Spectrum} = |C(a, b)|^2 \dots\dots\dots (2)$$

201 To differentiate between underlying periodicities and noise, Torrence & Compo (1998) created a
 202 statistical significance test for determining the significance of the time-localized oscillation in the
 203 wavelet power spectra. Fig. 2b shows the zoomed version of the GIC amplitude time-series

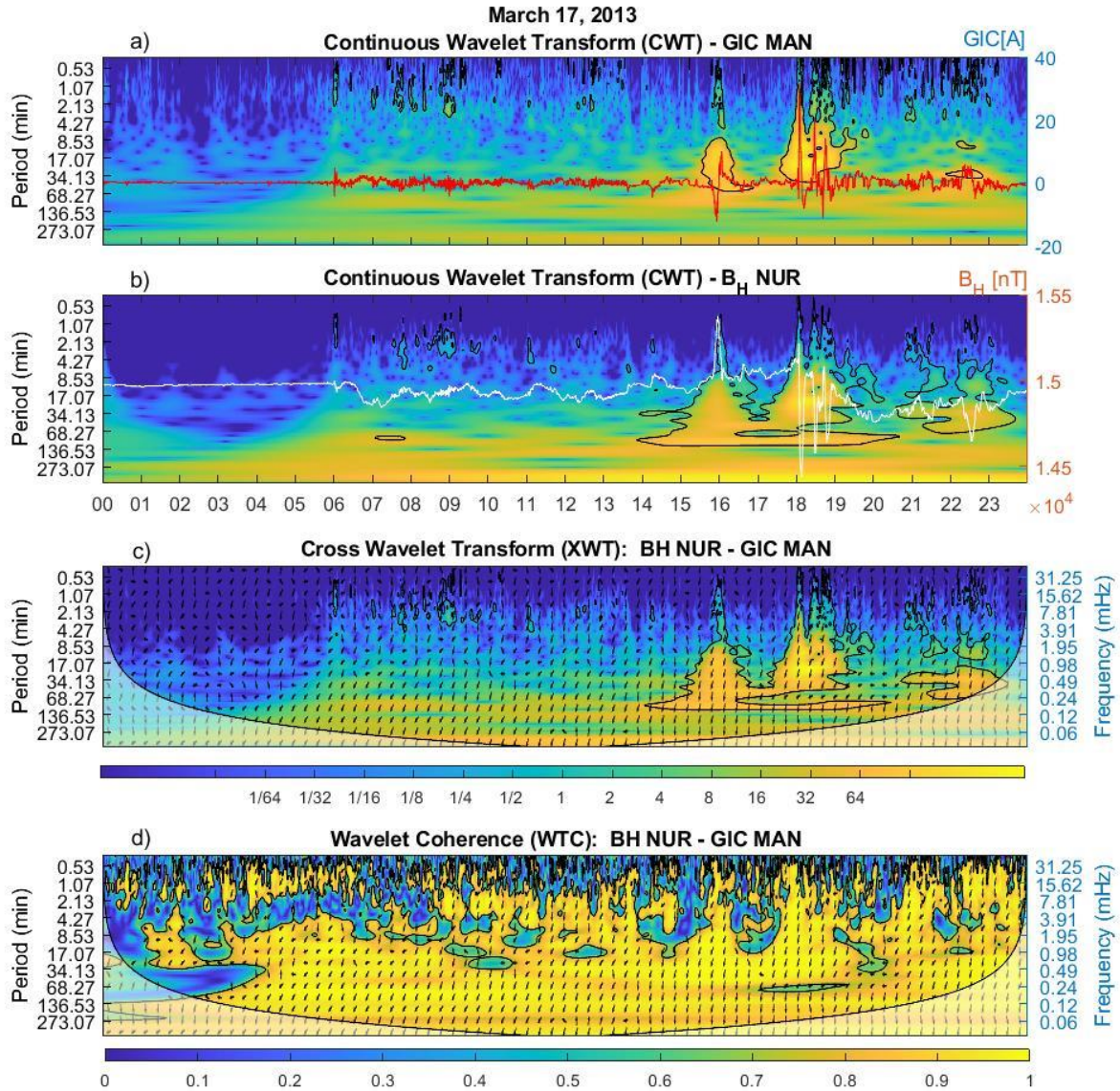
204 between 15-19 UT to confirm that CWT shows significant identifiable periodicities.



205

206 **Figure 2: a) Continuous Wavelet Transform (CWT) of GIC shown as heatmap overlaid with**
 207 **GIC time-series. The left axis indicates period [minutes]; the right axis indicates GIC [A].**
 208 **Yellow indicates high wavelet power. Black contour lines indicate time-frequency oscillations**
 209 **that are statistically significant (95% confidence). The frequency boundaries for Pi1 and Pi2**
 210 **pulsation ranges are shown in white. b) GIC time-series (absolute values) zoomed on 15:00-**
 211 **19:00 UT to show the underlying periodicities. Annotated arrows in a) show key periods**
 212 **identified in the time-series.**

213 **2.2 Cross Wavelet Transform (XWT) and Wavelet Coherence (WTC)**



214

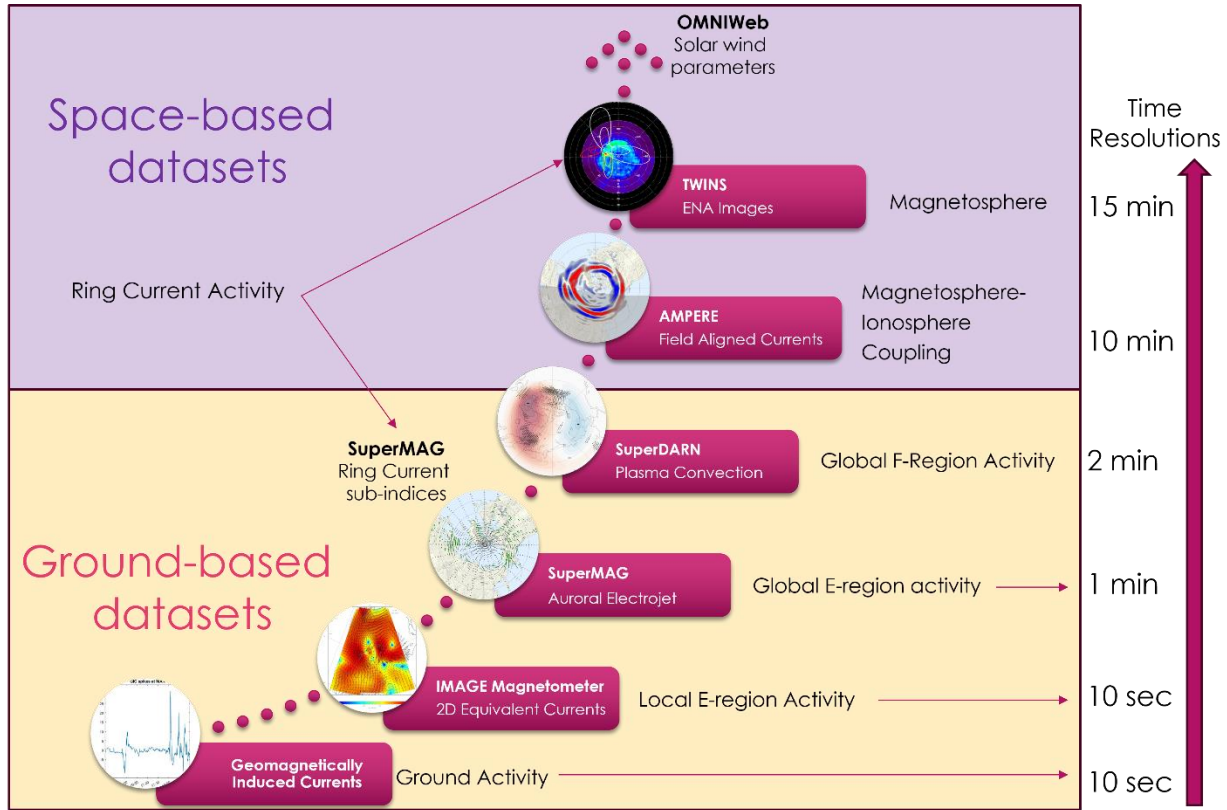
215 **Figure 3: Wavelet analysis - a) Repeat Fig. 2a. b) CWT of BH at NUR overlaid with BH Time-**
 216 **series [Right axis indicates the magnitude of BH [nT]]; c) Cross Wavelet Spectrum (XWT) of**
 217 **GIC and B-field with phase arrows indicating lag between the time-series; d) Wavelet**
 218 **Coherence of (WTC) of GIC and B-field indicating the correlation between the two time-**
 219 **series. Left Axis indicates period [min]. The black arrows in (c) and (d) indicate the phase angle**
 220 **between the two time-series. Translucent highlights at bottom left and right corners show cone of**
 221 **influence.**

222 Grinsted et al. (2004) extended the statistical test to cross-wavelet analyses. We adapt
223 their code (<https://grinsted.github.io/wavelet-coherence>) to generate Fig.3. The associated
224 equations for creating the plots have been described in detail by the authors. This method
225 ascertains that the time-frequency oscillations of any signal are significantly different from the
226 background red noise power spectrum (that captures randomness typically found in geophysical
227 time-series) and that the power spectrum of GIC is neither noise nor random. Fig. 3a-b show the
228 CWT of GIC and BH respectively.

229 The XWT between BH and GIC identifies common time-frequency oscillations and
230 shows the relative lag between the two time-series as small black arrows overlaid on the color
231 map. The phase arrows represent the following: right (in-phase), left (anti-phase), down (BH
232 leading GIC by 90°), and up (GIC leading BH by 90°). WTC resembles the traditional
233 correlation coefficient but is superior to that since it shows a correlation between the two signals
234 at different frequencies. Fig. 3c shows the power of XWT coefficients with similar color
235 schemes as CWT. Fig. 3d shows the WTC colormap with yellow indicating a high correlation
236 and blue indicating a low correlation.

237 The CWT, XWT, and WTC have edge artifacts because the wavelet is not completely
238 localized in time. It is therefore useful to introduce a Cone of Influence (COI) in which edge
239 effects cannot be ignored, which is shown as gray highlighted region in the bottom right and
240 bottom left corners of the plots. The COI is removed from panels d and e to show the time-series
241 clearly (GIC and BH respectively), but it occupies the same area as (c) and (d).

242 **2.3 Ground-Up Approach to Information Fusion**



243

244 **Figure 4: Summary of datasets used for the Ground-Up Approach using ground and space-**
 245 **based observations. Right side shows the time resolutions of the observations.**

246 Fig. 4 illustrates the ground-up approach with datasets that sample the activity at different
 247 resolutions and altitudes, thus providing support for physical interpretation of the significant
 248 multiscale periodicities identified in wavelet plots. We use three tools in this paper – wavelet
 249 analysis for time-frequency perspective, global maps for spatial perspective, and keograms for
 250 localized-temporal perspective. The waveform (1-D time-series) and wavelet transform (2-D
 251 time-frequency heatmaps) reveal underlying frequencies in the GIC and BH signal. Spatial maps
 252 around the time of GIC spikes provide a local/global context of ionospheric and magnetospheric
 253 activity. Keograms for the evolution of AMPERE-derived FAC and IMAGE 2D-EC provide
 254 insights into the magnetospheric drivers and their duration.

255 **3 Results**

256 The first GIC spike has a sinusoidal waveform with 12 A peak at 15:56 UT (G1a) and
257 9.8A peak at 16:04 UT (G1b). For reasons described later, we refer to this as a ‘compound’ GIC.
258 The largest GIC peak of 31.65 A occurs at 18:04 UT (G2). Two more GICs >10 A occur at 18:28
259 (G3) and 18:46 UT (G4). We discuss the observational differences between these spikes in the
260 rest of the section.

261 **3.1 Spectral Similarities and Differences in Wavelet Analysis**

262 The CWT heatmaps in Figs. 2a and 3a-b show minimal fluctuations in the GIC and BH
263 spectra before the shock arrival at 06 UT; thereafter intensified high-frequency variations are
264 present during the passage of the pre-CME sheath (06 UT - 10 UT) while the MAN/NUR region
265 is on the dayside. The large amplitude GIC spikes occur in tandem with spikes in BH after the
266 magnetic cloud arrives at ~15:30 UT. Both heatmaps show two distinct multi-minute regions of
267 time-frequency oscillations - one at ~16 UT (<1 min to ~30 min period, and the other between
268 18-19 UT (<1 min to ~1h period). XWT power in Fig. 3c shows phase arrows (down and/or
269 down-left) within the 95% significance contour indicating that BH leads (down) the GIC
270 response and is out-of-phase (left) after 15 UT. This is consistent with W.-H. Xu et al.'s (2022)
271 results and confirms the physical intuition of BH variations driving GICs. The WTC plot (Fig.
272 3d) shows a low correlation in the two time-series at higher frequencies before the shock arrival
273 and a high correlation thereafter. The correlation is much higher across all frequencies during
274 GIC spikes. That is: the duration and variation of activity affecting BH is also reflected in the
275 GIC signal.

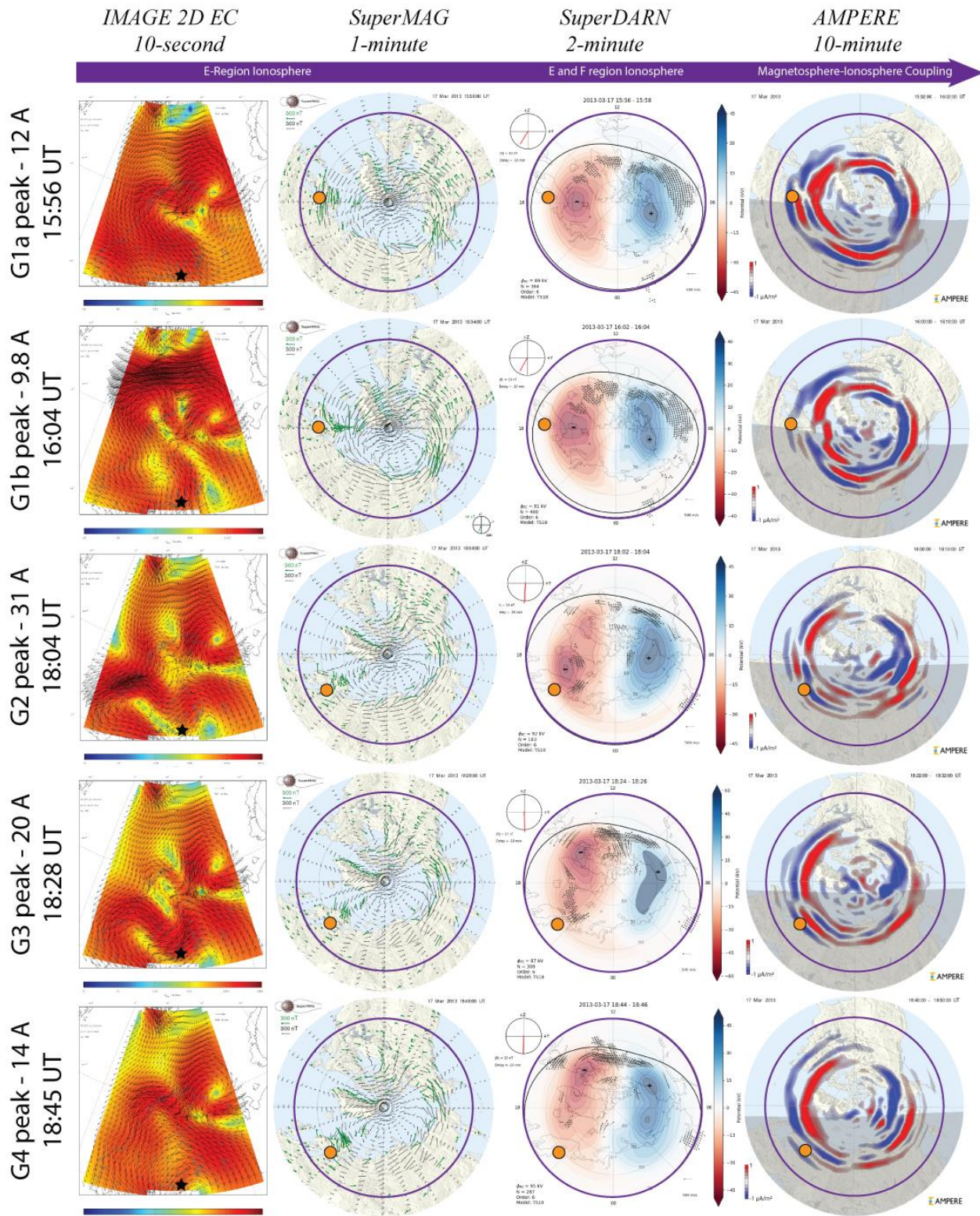
276 The key takeaway from the spectral analysis is that around 16 UT, the ‘compound’ GIC
277 spikes occur over 20 minutes with peaks having an 8-minute separation. Subsequently, three

278 periodic spikes (G2-G4) occur over 40 minutes; G2 and G3 are 24 minutes apart and G3 and G4
279 are 17 minutes apart. The higher frequency fluctuations in the range Pi1 and Pi2 (periods < 40 s
280 and 40-150 s, respectively) shown in Fig. 2a have been associated with substorm onset (Saito,
281 1969 and references therein) which is discussed in Section 4. We suspect there is more Pi2
282 fluctuation in the GIC signal compared to BH (or dBx/dt from W.-H. Xu et al. (2022)) due to the
283 integrated effect over the pipeline.

284 **3.2 Ionospheric Activity and Coupling with Magnetosphere**

285 The combination of signal waveform and spectral heatmap suggests different driver(s) for
286 the GIC spike around 16 UT compared to spikes between 18-19 UT. We explore the ionospheric
287 activity and the associated coupling with the magnetosphere to learn about the different drivers
288 in Fig. 5. The first column shows the local IMAGE 2D-EC spatial map over Finland while the
289 second column shows the global map of the SuperMAG equivalent currents in the northern
290 hemisphere. The third and fourth columns show polar plots of the NH SuperDARN convection
291 maps and AMPERE FACs, respectively. We organize the spatial maps in increasing altitudes and
292 temporal resolutions, which also capture multiscale M-I activities.

293



294

295 **Figure 5: Matrix of observations (Top to Bottom) Ionospheric activity and coupling with**
 296 **magnetosphere during the instances of peak GICs labeled as G1a, G1b, G2, G3, G4. (Left to**

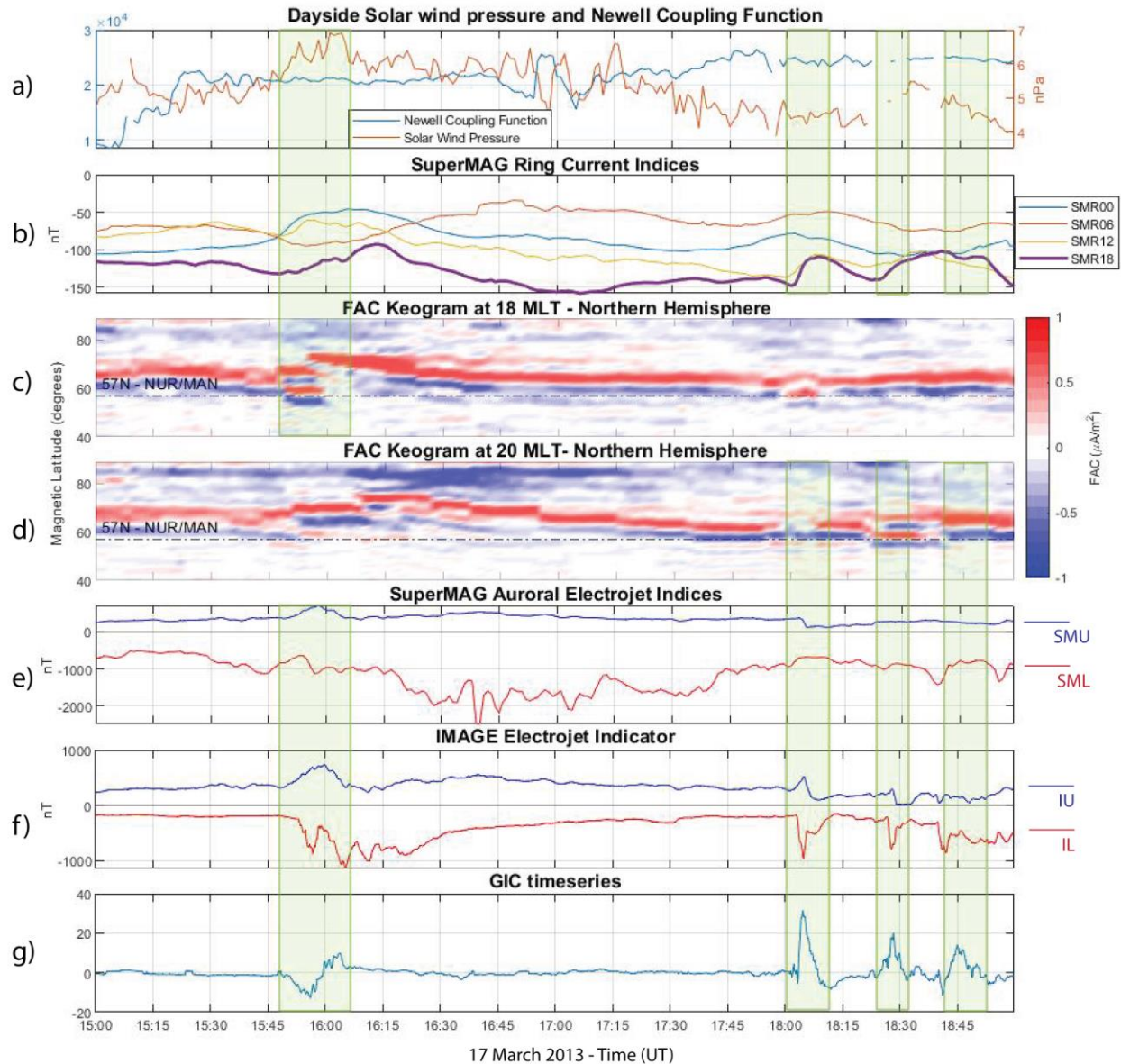
297 *right) IMAGE 2D Equivalent Currents (EC) [NUR marked with black star], SuperMAG plots*
298 *with green vectors indicating electrojet pattern, SuperDARN convection maps with black*
299 *vectors showing plasma velocity. AMPERE-derived upward (red) and downward (blue) FACs,*
300 *Outlines mark 50 MLAT on all plots. Location of MAN marked as orange circle.*

301 The interval at ~16 UT, shows two GIC extrema occurring within 20-minutes. The top
302 row of Fig. 5 shows NUR station (black star) at 15:56 UT under a region of strong equivalent
303 current at the equatorward edge of the IMAGE 2D-EC map. SuperMAG suggests a strong
304 localized eastward electrojet (eEJ) over MAN/NUR, with a broad westward auroral electrojet
305 (wEJ) poleward of the location. This is consistent with the rising trend in BH. Although the
306 SuperDARN convection map lacks duskside plasma velocity vectors, possibly due to the
307 expanded auroral oval, the line-of-sight (LoS) spectrogram of plasma velocity (see supporting
308 information Fig. S1) shows near-range echoes at Hankasalmi station, indicating activity in the E-
309 region and suggestive of an enhanced eEJ. AMPERE-derived FACs show a highly structured
310 layered pattern (alternating reds and blues) from the pole to the equator, with MAN/NUR
311 beneath the transition of upward and downward FACs. For context, typically, Region 1 (R1)
312 FAC points into the ionosphere on the dawnside (blue) and out of the ionosphere on the duskside
313 (red) whereas Region 2 (R2) FAC (equatorward of R1) points downward on the duskside and
314 upward on the dawn-side (Iijima & Potemra, 1976). The DMSP SSUSI emission maps (see Fig.
315 S2) show duskside precipitation stretching from 65°-75° MLAT, indicative of highly structured
316 ionospheric conductivity.

317 The second row of Fig. 5 follows the same sequence for 16:04 UT. The IMAGE 2D-EC
318 gained a significantly structured north-south component. Similarly, the SuperMAG plot reveals a
319 narrow equivalent-current channel north of MAN. The AMPERE-derived FACs show that the
320 layered structure has subsided and the R1 FAC at 18 MLT has moved poleward. The DMSP

321 Southern Hemisphere emission map (Fig. S2) shows a bright spot just above 65° MLAT and
322 slightly west of the 18 MLT line.

323 The bottom three rows in Fig. 5 suggest that MAN/NUR were under the influence of
324 shearing eastward and westward electrojets during G2-G4. The SuperDARN convection maps
325 show that a mesoscale plasma vortex structure (identified by the tight curl) forms and dissipates
326 over the three timestamps. FACs appear to have a complex upward and downward structure. The
327 sequence of observations shown in Fig. 5 suggests different magnetospheric mechanisms
328 affecting the ionosphere at G1(a and b) compared to G2-G4. We discuss these in Section 4.



329

330 **Figure 6: Ground-to-Space dynamics between 15-19 UT. a) SuperMAG-derived parameters**
 331 **Newell Coupling Function and solar wind dynamic pressure; b) SuperMAG Ring current**
 332 **indices at four local times; c,d) FAC keogram at 18 MLT and 20 MLT respectively. Red**
 333 **indicates upward FAC, Blue indicates downward FAC. e) SuperMAG Auroral Electrojet Index**
 334 **Upper (SMU/eastward) and Lower (SML/westward) values for global substorm activity; f)**
 335 **IMAGE Magnetometer-derived Electrojet Indicator Upper (IU) and Lower (IL) values for**
 336 **local substorm activity; g) GIC signal. Green highlights mark the duration of GIC spikes.**

337

Additional parameters that add insight to the state of magnetosphere during these GIC

338

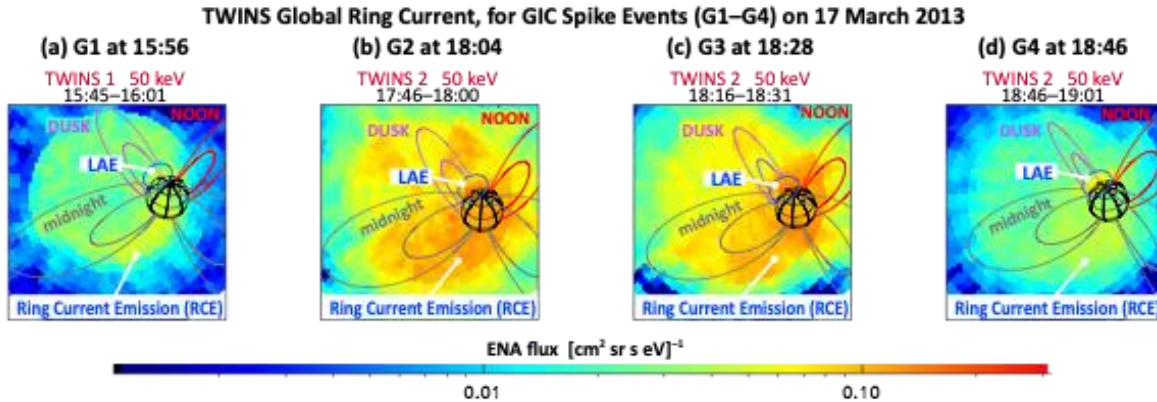
spikes are shown in Fig. 6. The rate of magnetic flux addition (Fig. 6a) quantified by the Newell

339

coupling function (Newell et al., 2007) increases from 1×10^4 Wb/s to 2×10^4 Wb/s from 15:00 to

340 15:30 UT. Such a flux increase, followed by the rising solar wind pressure (brown curve), should
341 eventually trigger magnetic reconnection on the night side and consequent enhancement of the
342 duskside RC. Figure 6b shows the sector-wise RC proxies. We focus on the SMR18 sub-index
343 since MAN is at dusk. A decrease in the duskside RC intensity around 16 UT is followed by RC
344 intensification for 2 hours before reducing again. In Fig. 6c and d, we show FAC keograms at 18
345 MLT and 20 MLT, respectively, to focus on the localized coupling during the G1a-b and
346 subsequent spikes. These FACs drive the ionospheric currents (electrojets) and affect the global
347 substorm values (Fig. 6e) of the Auroral Electrojet, which is represented by the SuperMAG SML
348 and SMU indices. These indices are associated with global westward and eastward electrojet
349 activity respectively. The local substorm indices IL and IU derived from IMAGE magnetometer
350 (Fig. 6f) show different features than the global indices suggesting localized effects over the
351 pipeline (Fig. 6g).

352 The negative GIC peaking at 15:56 UT (G1a) corresponds to 1) intensification of local
353 eEJ (SMU and IU) and dip in local wEJ (IL), 2) downward (blue) FAC split by upward FAC,
354 and c) decreasing RC intensity at midnight, dusk, and noon. The positive GIC peaking 8 minutes
355 later at 16:04 UT (G1b) corresponds to 1) decaying local eEJ with a second dip in IL, and 2)
356 poleward movement of R1 FAC with weakened equatorward layers. This suggests a major
357 reconfiguration in the regional M-I coupling. During G2-G4, there is a decrease in global auroral
358 indices, complex FAC structure, and fluctuating RC intensity at dusk but an increase in intensity
359 at midnight. These spikes rise and fall thrice over a duration of 40 minutes, again, captured by
360 CWT as significant periods (Fig. 2). From Fig. 6, we infer that the G2-G4 spikes are a result of
361 localized M-I coupling.



362

363 **Figure 7: TWINS ENA images of 50keV particles showing large-scale ring current (RC)**
 364 **activity when GIC spikes (G1 – G4) are recorded.** In each image, the Earth is depicted, and
 365 dipole field lines are drawn at L=4 and 8 (Red = noon MLT, purple = dusk MLT). Each $4^\circ \times 4^\circ$
 366 ENA image is integrated over 15 minutes. Each image includes a ring current emission (RCE)
 367 and low altitude emissions (LAE).

368 The TWINS ENA images provide a large-scale perspective of the inner magnetospheric
 369 state and RC dynamics. Figure 7 shows snapshots of 50 keV ENA images, at the nearest
 370 available times to the four GIC spikes. Each image contains the following information – Earth’s,
 371 the two dipole field lines at L=4 and 8 with color coding to indicate MLT: red (noon), purple
 372 (dusk), and grey (midnight and dawn). Each pixel indicates line-of-sight (LOS) integrated ENA
 373 flux, accumulated over 15 min. On 17 March 2013, both TWINS 1 and 2 imagers frequently
 374 observed elevated background counts, most likely from local (to TWINS) energetic ions
 375 penetrating past the collimator plates that are supposed to keep these ambient ions out
 376 (McComas et al., 2012). Hence, we performed a sensor-dependent flat-field and background
 377 subtraction for each TWINS image described in the Data and Method section.

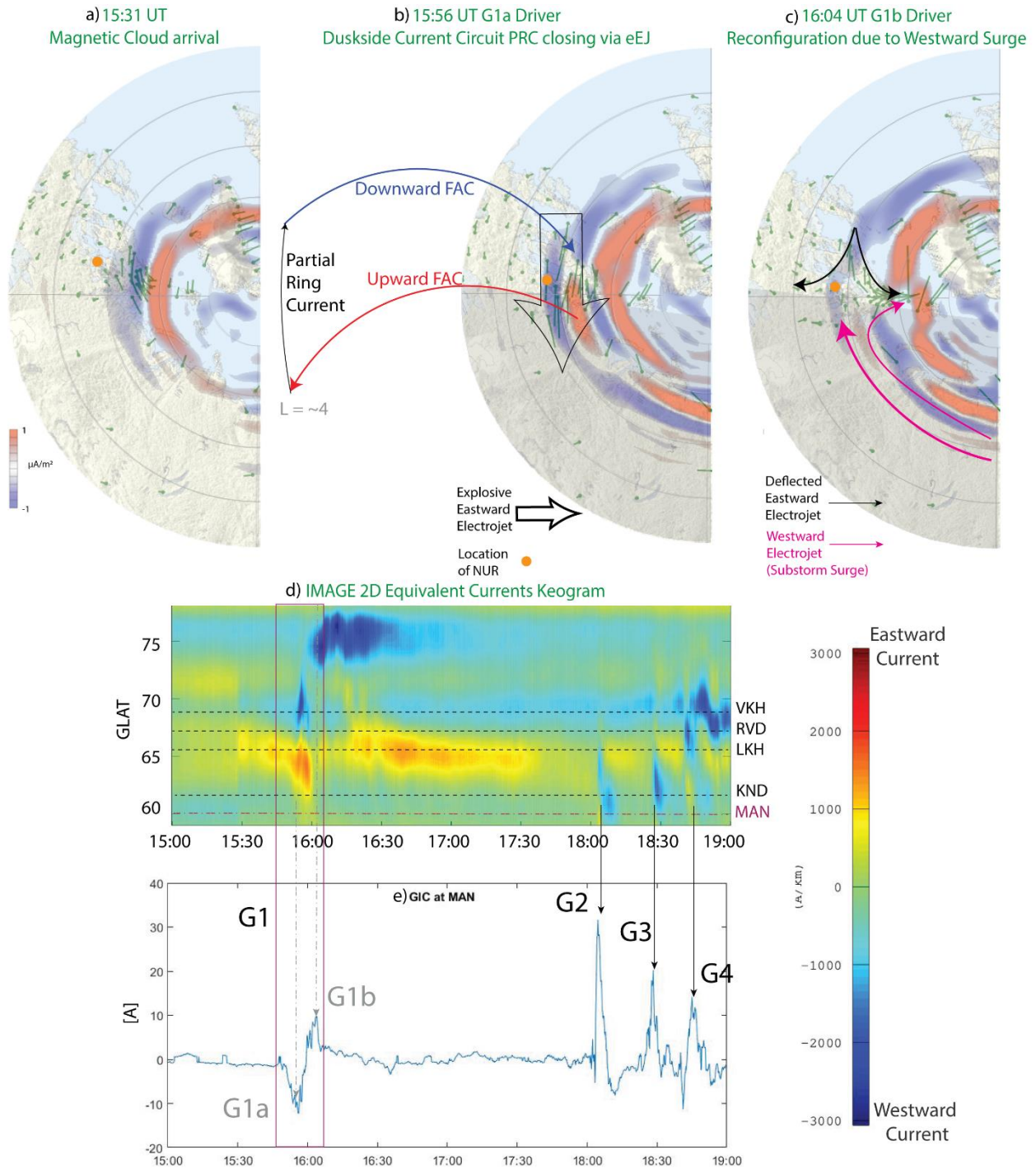
378 During all GIC spikes, the presence of Low Altitude emission (LAE) suggests significant
 379 ion precipitation from the duskside RC, even though the ENA flux appears weaker there. The
 380 reason could be anisotropic duskside ion pitch angle distributions (PADs) generating fewer
 381 ENAs in the directions of the TWINS locations (Goldstein et al., 2012), rather than from a

382 weaker RC. Near the time of G1a (Fig. 7a), TWINS observed a weak RCE and modest LAE at
383 dusk. This LAE enhancement was not observed in the image 15 minutes before, indicating new
384 RC activity enhanced the precipitation. Near the times of the larger GIC spikes of G2–G4,
385 TWINS imaged a significant global RC enhancement, as well as an LAE intensification. For the
386 largest GIC peak, G2, the RCE was enhanced most strongly in the post-midnight sector, but also
387 in patches of enhanced ENA flux near dusk (Fig. 7b). The non-uniformity of the duskside
388 increase may reflect an MLT-dependent PAD evolution that modulates the ENA intensity, but
389 also may be (at least partly) an artifact from the background subtraction. The closest-in-time
390 images to G3 (Fig. 7c) and G4 (Fig. 7d) likewise depict global RC changes concomitant to the
391 strength of the GICs. That is, these TWINS imaging observations confirm the difference in RC
392 activity corresponding to the GIC spikes, indicating magnetospheric sources for the GICs. Søråas
393 et al. (2018) provide strong supporting evidence based on NOAA and MetOps satellite particle
394 precipitation data. They note that the isotropic particle precipitation (PP) in the upper
395 atmosphere is related to particles injected from plasmashet (BBF) into the inner magnetosphere
396 (up to $L = 2.7$) and hence is a proxy for RC injections. They further find an increase in ENA flux
397 around 15-19 UT in the equatorial region and elevated PP in the RC energy range over MAN
398 (50-60 MLAT). Hence, the duskside LAE signal in Fig. 7 can be explained by precipitation
399 originating from RC injections.

400 **4. Discussion**

401 We consider impinging mesoscale plasma flows as likely GIC sources. Based on
402 available data and models, we argue that the magnetospheric source(s) of the compound GIC
403 spikes at 16 UT was a larger and more energetic flow burst than those driving the self-similar
404 GIC spikes between 18-19 UT.

405 **4.1 Drivers of GIC peak at ~16 UT**



406

407 **Figure 8 a-c): AMPERE-derived FACs superposed on SuperMAG plot with 50% transparency**
 408 **at 15:56 UT (G1a) and 16:04 (G1b) to interpret their drivers. Arrows indicate the direction of**
 409 **eastward and westward ionospheric current surges** d) **IMAGE 2D-EC keogram with 15 to 19**
 410 **UT. MAN latitude marked in red dotted line. Transformer stations (VKH, KND, RVD, LKH) in**
 411 **the Kola Peninsula (Belakhovsky et al. 2019) marked in dotted black lines, e) Repeat of Fig 6g.**

412 *G1 highlighted in the box. Arrows show the connection between overhead ionospheric activity*
413 *driving GICs.*

414 For context, Fig. 8 (top row) shows polar views of the near-dusk equivalent current
415 system with AMPERE FACs overlaid. Figure 8a displays the current pattern as the magnetic
416 cloud passed Earth when the duskside was dominated by an Ijima-Potemra FAC pattern, typical
417 of a well-developed RC. The negative portion of G1 (Figs. 8b and e) coincided with a major
418 current reconfiguration and a sudden duskside eEJ enhancement. The positive portion of G1
419 (Figs. 8c and e) coincided with another major current reconfiguration and a sudden development
420 of a poleward equivalent current channel and an enhanced wEJ at high latitudes. Recall also that
421 the CWT (Figs. 2 and 3a-b) showed distinct Pi1/Pi2 pulsations during the 30-minute interval
422 around 16 UT, which are indicative of substorm activity in the auroral and subauroral regions
423 (Kepko & Kivelson, 1999; Milling et al., 2008 and references therein).

424 ***Event G1a.*** Figure 8b (15:56 UT, G1a) shows the most intense eEJ developed near 60°
425 MLAT close to the southern end of the pipeline. The nearby NUR magnetometer recorded a
426 positive deviation in BH lasting for ~30 min (Figs. 1b and 3b), consistent with the rising IL
427 index in Fig. 6f. With the already strong RC (Fig. 6b), the *prima facie* evidence is that the G1a
428 spike was driven by a sudden PRC closure (see the current loop in Fig. 8b) that increased
429 eastward current in the duskside ionosphere. Supporting this is the sudden TWINS LAE (Fig. 7a)
430 and conjugate DMSP F16 and F17 particle data for the SH (not shown) that confirm new RC
431 precipitation (10-40 keV ions) equatorward of -60° MLAT. Below we build a case that supports
432 G1a and G1b activity as driven by an impinging mesoscale flow channel that influenced current
433 systems across a range of latitudes.

434 Kamide & Fukushima (1972) championed the idea of a PRC closing via an eEJ in the
435 duskside ionosphere as part of long-lived, storm-time-enhanced convection event. In contrast, the
436 pre-16 UT eEJ enhancement was short-lived and seemingly disrupted by a major duskside
437 current reconfiguration within 10 minutes. Grafe et al. (1998) and Feldstein et al. (1999)
438 designated similar transient, duskside eEJs as ‘explosive’ events and noted an association with
439 substorm-driven wEJ onset at higher latitudes, although the exact timing and relationship
440 between the electrojets was unclear. Chen et al. (2020) reported that strong substorm injections
441 generate Electromagnetic Ion Cyclotron (EMIC) waves in the dusk sector. This leads to an
442 inference that wide flow channel penetrated the duskside inner magnetosphere and created waves
443 that triggered RC precipitation at the duskside plasmopause bulge/plume as suggested by
444 Trakhtengerts & Demekhov (2005) and Spasojevic & Fuselier (2009) thus, potentially
445 contributing to the formation of a PRC-driven eEJ at $L \leq 5$. A superposed epoch analysis by
446 D’Onofrio et al. (2014) showed that eEJ enhancement precedes wEJ enhancement during
447 moderate substorms.

448 Using EISCAT and IMAGE magnetic observatories and several Russian observatories, as
449 well as DMSP particle data, Feldstein et al. (1999) determined that the equatorward side of eEJ
450 was bounded by the plasmopause projection to ionospheric height, and that the poleward
451 boundary of eEJ was the ionospheric projection of the plasmashet inner boundary. During
452 substorms the wEJ widened to fill the auroral zone and at times forced the auroral boundary
453 poleward. They reasoned that the storm time eEJ in the evening sector of the magnetosphere
454 linked to processes in the inner magnetospheric regions adjacent to the plasmashet inner
455 boundary.

456 Feldstein et al. (2006) studied the 25 September 1998 storm and found that eEJ had
457 contributions from a PRC closure at lower latitudes as part of an intense substorm-sub auroral
458 polarization stream (SAPS) interaction. Yang et al. (2012) simulated a strong plasma bubble
459 injection impinging on the nightside ionosphere (23 MLT). One obvious effect in the simulation
460 was the creation of a SCW two-loop (SCW2L) circuit of evolving R1 and R2 currents that match
461 the pattern evolution shown in Figs. 8b and 8c. Mishin et al. (2017), studying both the 25
462 September 1998 event and the 17 March 2013 event, established a causal relationship between
463 fast RC injection via mesoscale flow channels, a SCW (westward moving front creating Pi2
464 pulsations), and duskside SAPS via an SCW2L.

465 Prior studies of the 17 March 2013 event reveal that: 1) near dusk, the plasmasphere was
466 eroded to $< 4 R_E$ (Krall et al., 2017) leading to Pc1 fluctuations and injection signatures in the
467 Russian (50° - 60° GLAT) sector (Potapov et al., 2017); 2) the RC was well-developed and
468 supported an active SAPS channel near $L = 4$ (Ferdousi et al., 2019; Lin et al., 2021); and 3)
469 Sorathia et al. (2018)'s simulation shows a wide plasma injection (mesoscale flow channel) in
470 the evening sector, although the modeled injection impinged slightly later in MLT. Thus, Figs.
471 8a-d along with results in Section 3 support the idea of 'compound' 16 UT GIC with negative-
472 then-positive spikes (Fig. 8d) driven by a dusk-region bubble injection/substorm. The injection
473 first disturbed the RC and then the auroral zone. While we have insufficient information to
474 determine how the large-scale current systems were being modified by mesoscale drivers to
475 produce the GICs, we anticipate that additional coupled mesoscale simulations (e.g. Bao et al.,
476 2023) could provide insight into the related physics.

477 **Event G1b.** Just before 16 UT significant global and local changes in the geomagnetic
478 field occurred. The equivalent current and FAC patterns shown in Figs. 8a and 8b were

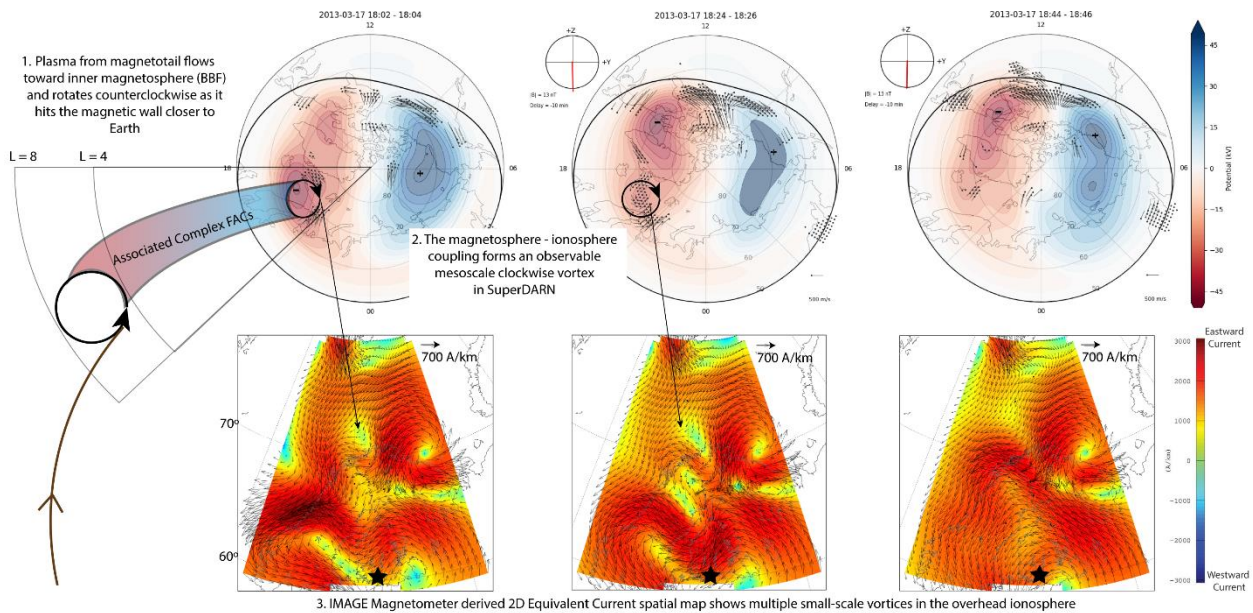
479 disrupted. We associate these changes with event G1b. Over Finland and Kola regions,
480 meridional equivalent currents rapidly replaced the east-west currents. From 15:56 to 16:04 UT
481 on the Mäntsälä pipeline, the minimum-to-peak GIC variation was 22 A; in the same interval,
482 there was a 70 A variation at Vykhodnoi power station in the Kola Peninsula. Figure 8c displays
483 the currents for 16:04 UT (G1b) and shows an intense poleward current channel over the north
484 end of the pipeline and the Kola peninsula, where strong ground magnetic fluctuations were
485 observed (Belakhovsky et al., 2019; Despirak et al., 2022). The narrow current channel,
486 bracketed by up-down FACs (Fig. 8c) and a poleward displacement of the R1-FAC (Figs. 6d) are
487 consistent with the description of auroral streamer development during arrival of plasma flow
488 bursts (e.g. Hubert et al., 2007; Zou et al., 2022). It was also during this time that significant Pi2
489 fluctuations appeared in the MAN CWT (Fig. 2a). Although we lack auroral imagery during this
490 time frame, we offer streamer development near the north end of the pipelines as a plausible
491 cause of G1b. The sharp enhancement in POES/MetOps particle precipitation data supports this
492 idea (see Figure S3),

493 Two essential findings arise from this synthesis: 1) based on observational evidence, we
494 assert that afternoon eEJ, perhaps amplified by substorm dynamics, can lead to significant GIC
495 spikes near dusk and 2) based on circumstantial evidence we assert that the meridional
496 equivalent current channel and related FACs are a further (substorm) manifestation of the
497 energetic mesoscale plasma flow and a source of GICs. In the next section, we address a set of
498 related GIC spikes that may be more typical of the storm recovery phase.

499 **4.2 Driver of GIC peaks between 18-19 UT**

500

Simplified Illustration of Bursty Bulk Flow Channel Driving 3 Periodic GIC spikes between 18-19 UT



501
 502 **Figure 9: Illustration of BBF creating three meso-scale vortices (SuperDARN) in the topside**
 503 **ionosphere leading to smaller-scale vortices (IMAGE 2D-EC) that eventually lead to the three**
 504 **periodic GIC spikes.**

505 Approximately two hours after G1, three periodic spikes arise when MAN is in the post
 506 dusk region (~20 MLT). In Fig. 5, the SuperMAG plots show MAN situated under opposing
 507 eastward and westward current vectors. Movie S1 clarifies that the vectors rotate
 508 counterclockwise as localized mesoscale vortices three times. The direction of the vortices is
 509 consistent with the SuperDARN clockwise vortex-like structure. These vortices occur
 510 simultaneously with local electrojet index spikes (Fig. 6f) and GICs. In the SuperDARN
 511 convection map, a meso-scale vortex is located between ~65-70 MLAT and 18-21 MLT
 512 (poleward of MAN). Using $L = \frac{r}{\cos^2(\lambda)}$ where λ is the geomagnetic latitude and $r = 1\text{Re}$, the 60-
 513 70 MLAT in the ionosphere traces into the equatorial magnetosphere at a distance of ~4 to 8 Re
 514 in the post-dusk region. In this region Sorathia et al. (2018, 2023)’s MHD simulation movies
 515 show increased bursty bulk flow (BBFs) between 18-19 UT.

516 Figure 9 illustrates the idea of BBFs driving GICs, as suggested by Wei et al. (2021) who
517 reported that BBF-associated FACs alter ionospheric currents and create dB/dt variations.
518 Keiling et al. (2009) note vortices in equivalent ionospheric currents could be related to
519 magnetospheric plasma flows. BBFs are rapidly flowing plasma channels from magnetotail to
520 the inner magnetosphere. As BBFs approach closer to Earth, the flow speed reduces and
521 eventually hits ‘a magnetic wall’. This causes the plasma to rotate away from Earth to form a
522 vortex-like structure and merge into the RC (Yu et al. (2017) and references therein). We
523 observe clockwise (cw) plasma vortex in the SuperDARN (Fig 9 top row) convection maps.
524 Based on M-I coupling, we can infer that the cw-rotation in SuperDARN convection map is due
525 to a counterclockwise plasma rotation in the equatorial magnetosphere. Velocity direction (\mathbf{v}) is
526 governed by $\mathbf{v} = \mathbf{E} \times \mathbf{B}$, where \mathbf{v} rotates clockwise and \mathbf{B} points toward the earth’s surface in the
527 high-latitude ionosphere, which results in \mathbf{E} directed toward the center of the vortex. Assuming
528 equipotential lines, the ionospheric \mathbf{E} has the same direction in the equatorial magnetosphere, but
529 \mathbf{B} is parallel to the earth’s surface at the equator, thus the plasma circulation is counterclockwise.
530 The 10-minute FAC maps in Fig. 5 (rows 3-5) shows a layered structure making it difficult to
531 decipher the exact dynamics, but the simulation by Sorathia et al. (2023) shows the complex
532 evolution of BBF-associated FACs in the 20 MLT sector, where we note the vortices. To
533 summarize Fig. 9, the BBFs in the post-dusk magnetosphere create meso-scale vortices in the
534 ionosphere inferred from SuperDARN (and SuperMAG) via FACs, that likely result in rapidly
535 changing small-scale vortices observed in IMAGE 2D-ECs, forming three periodic GIC spikes.

536 Juusola et al. (2009) studied the ionospheric signature of BBF using the IMAGE 2D-EC
537 maps and Cluster satellite-data. They found that the average duration of BBF was 8 minutes (also
538 seen in Fig. 2b) accompanied by a northwest channel at high latitudes (62-76 MLAT) creating a

539 pair of small-scale upward and downward FACs. Recently, Engebretson et al. (2024) found that
540 BBF drive short-lived high amplitude dB/dt spikes (like 17 March 2013) using the Canadian
541 ground-magnetometers and THEMIS satellite-data in the 23-01 MLT sector, consistent with
542 Juusola et al's (2009) findings. They estimate that a flow channel, during quiet conditions,
543 creates an oval of ~ 180 km (east-west) by 90 km (north-south) in the ionosphere, which is an
544 order of magnitude smaller than the vortex we note in Fig 9. We suspect that the signature
545 reported by these authors can be amplified in moderate storm by creating either wider flow
546 channels or the mesoscale vortices are superposed on smaller-scale vortices, which would be
547 consistent with the multi-minute spectral area in CWT.

548 The dearth of duskside in-situ satellite observations on 17 March 2013 the BBF
549 interpretation an inference, but the presence of Pi1/Pi2 pulsations and the <8-minute duration of
550 spikes (Fig. 2a) provide confidence. First, these substorm-associated pulsations are present in BH
551 and GIC heatmaps (Figs. 3a-b) when the local westward (substorm) currents are enhanced (Fig.
552 6f). These BBFs produce small-scale wedgelets that combine to form a large-scale SCW (Palin
553 et al., 2016) while also enhancing the RC (Fig. 7b-d) (Goldstein et al., 2020; Sciola et al., 2023).
554 The flows that penetrate the inner magnetosphere create a compression, which has been observed
555 to create oscillations of 1-2 minutes periodicity on the ground (Kepko et al., 2001; Lysak et al.,
556 2015), like Pi2 pulsations in Figs 3a-b. Hence, the observations presented herein, supported by
557 previous studies, suggest that BBFs generate magnetospheric vortices, which connect to the
558 ionosphere by FACs, whose ionospheric closure can cause GICs. Our work extends the
559 relationship from BBF driving short-lived large amplitude GMD directly to GICs.

560 Belakhovsky et al. (2019) reported substorm-associated ionospheric vortices over the
561 Kola Peninsula between 15-19 UT. (See Fig. 8d for latitudes of the transformer stations

562 discussed in Fig. 7 of their paper.) Simultaneously, the low-latitude Wp index (Nosé et al.,
563 2012) showed elevated substorm (Pi2) activity. Low-latitude Pi2s are now known to be
564 associated with FACs produced by BBF braking (Cao et al., 2008). Importantly, Figure 7a of Z.
565 Xu et al. (2017) provides a global context of the mid-to-low latitude Bx dynamics wherein we
566 find the fingerprint of the 4 mesoscale channels in the dusk sector from 16-18 UT. We contend
567 that the high-latitude vortices, low-latitude substorm indicators, and the sub-auroral G1-G4
568 disturbance in the Mäntsälä pipeline were latitude-spanning effects of a series of mesoscale flow
569 injections or impingements. Minor timing differences can be attributed to the difference in
570 overhead currents, ground and material conductivity, system configurations, and the north-south
571 orientation of the powerline (Despirak et al., 2022).

572 We gather the following key insights from this case study. After the magnetic cloud
573 arrived at 15:30 UT, the strong, steady negative Bz initiated a new cycle of magnetotail
574 reconnection. The interaction between duskside substorm-associated plasma bubble injection,
575 RC, and plasmasphere likely created a transient current circuit closing the PRC via mid-latitude
576 eEJ driving G1a. Immediate duskside current reconfiguration produced G1b. As the storm
577 progressed from late main- to recovery-phase and the pipeline moved from dusk toward
578 midnight, small/mesoscale earthward plasma flow (BBFs) took over to form vortices in the mid-
579 latitude ionosphere to create three large spikes. We conclude the discussion by offering a novel
580 perspective - during this event (and perhaps other events), the pipeline (long conducting
581 material) acted as a huge antenna, capturing multi-scale ionospheric activities.

582 **5. Contributions and Future Work**

583 Our work addresses the RC-GIC association gap identified by Ganushkina et al. (2017).
584 We suggest storm-time RC dynamics and the mesoscale flows the impinge poleward of the
585 duskside RC play a crucial role in creating sub-auroral GICs. Furthermore, Yu et al. (2022)
586 identified a modeling gap for integrating different spatial and temporal scales to achieve a
587 comprehensive understanding of the entire system. Our approach can provide observational
588 support to such modeling efforts.

589 Clearly, more work lies ahead. A follow-on wavelet analysis on the IMAGE
590 magnetometer chain, especially on all three components, during this event would provide more
591 insight into the sequence of driving current systems. Also, extending our analysis to conjugate
592 magnetometer (not shown, MAW – -70.67 MLAT) in SH, which shows larger perturbation than
593 NUR around 16 UT, would further illuminate the M-I coupling aspects of hemispheric
594 asymmetry. Analyzing other GIC events similarly will answer a crucial question: Under what
595 circumstances are GICs generated locally (such as 17 March 2013), vs globally (such as the
596 October 2003 Halloween event (Swedish Transformers – (Pulkkinen et al., 2005); New Zealand
597 transformers - (Marshall et al., 2012, 2017); South African Transformers - (Bernhardi et al.,
598 2008)) in addition to MAN)? What causes them to be short- vs long-lived? A co-occurrence
599 study, extending our approach to multiple locations, for a particular event can be useful for
600 assessing the state of the magnetosphere during GIC spikes that arise globally during later events
601 from Solar Cycle 24 onwards when more ground and space-based observations are available.
602 Moreover, a concerted effort between the simulations and our ground-up approach can provide a
603 better understanding of the topic, important for accurate GIC prediction.

604 Ground-based data played a pivotal role in our paper and hence we highlight the need for
605 a continuous ground-based stream of data. Our extent of interpretation is limited by the 10-
606 second data which provides substorm onset related information. However, Potapov et al.
607 (2017)'s frequency analysis of higher cadence magnetometer data (0.2 to 5 /sec) enabled them to
608 identify a relationship of Pc1 waves with plasmopause dynamics. This underscores the need for
609 higher sampling frequency data, which not only appears to keep relative peak errors below 10%
610 for predicting GICs (Grawe & Makela, 2021), but also appears to make the strongest
611 contribution at magnetic latitudes $<60^\circ$ (Hartinger et al., 2023) such as the Continental United
612 States and Europe.

613 **6. Conclusions**

614 The March 2013 St Patrick's Day storm provides a unique opportunity to identify the
615 magnetospheric root causes of the GIC spikes recorded on the Mäntsälä pipeline. The pipeline
616 and other ground-based observatories were in the right place at the right time to capture and
617 quantify the space weather impacts of this storm. The time-frequency perspective provided by
618 wavelet analysis shows spectral features spanning seconds to ~ hour around the four GICs >
619 10A. These multi-scale fluctuations are captured across ground- and space-based observations
620 with different resolutions. Together, CWT and data-fusion of multi-platform observations paint a
621 robust picture of the nature and scale of M-I coupling leading to significant GICs. Based on
622 supporting data and prior MHD modeling, we find that the first 'compound' GIC was likely the
623 result of a mesoscale flow channel interacting with the ring current and the plasmasphere. The
624 interaction manifested in the ionosphere as a transient eastward electrojet closing a partial ring
625 current on the duskside. In only a few minutes the same mesoscale flow channel produced a new
626 FAC structure, strong poleward auroral currents, and a > 20 A change in the pipeline GICs. Two
627 hours later, magnetospheric BBFs created intense small-to-mesoscale ionospheric vortices
628 leading to three periodic spikes, the largest of which was > 30 A. The CWT plot shows
629 significant Pi2 fluctuations, which have been associated with such bursts. Hence, we find
630 substorm injection of varying scales to be the underlying cause of these spikes. Our CWT
631 approach provides a framework for further research on the impact of GICs on different
632 technologies by offering new insights into M-I-ground coupling during geomagnetic storms.
633 With the improved simulation capabilities, we hope to gain a deeper understanding of the GIC
634 drivers. A similar analysis of other GIC events remains in the realm of future work.

635 **Acknowledgments**

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638 data, IMAGE for NUR magnetometer and 2D equivalent current keogram, SuperMAG and
639 SuperDARN collaborators, AMPERE team and the Science Data Center for providing data
640 products derived from the Iridium Communications constellation, enabled by support from the
641 NSF, TWINS team for ENA data, and OMNIWeb for the solar wind data and derived products.
642 Special thanks to Ari Viljanen for providing crucial feedback on the ideas and Steve Milan for
643 providing the IDL code for reference to create FAC keograms. K. Sorathia, M. Wiltberger and R.
644 Marshall provided useful feedback on our work.

645

646 **Open Research**

647 Mäntsälä Finnish pipeline (https://space.fmi.fi/gic/man_ascii/) is used to gather GIC data.
648 The recommended magnetometer data at the NUR station is gathered for the corresponding
649 geomagnetic field (https://space.fmi.fi/image/www/?page=user_defined). SuperMAG is used to
650 generate 90-degree rotated vector plots to give a sense of the auroral electrojet and for the SML,
651 SMR, and SMU indices and sub-indices (<https://supermag.jhuapl.edu/>). Active Magnetosphere
652 and Planetary Electrodynamic Response Experiment (AMPERE) data is used from
653 (<https://ampere.jhuapl.edu/>) for generating Field Aligned Current patterns. The solar wind data
654 is retrieved from OMNIWeb (https://omniweb.gsfc.nasa.gov/form/omni_min.html). TWINS
655 data are accessible to the public at <https://cdaweb.gsfc.nasa.gov/>. Fitted SuperDARN data can be
656 downloaded from Globus, instructions of which are provided here: [https://superdarn.ca/data-](https://superdarn.ca/data-products)
657 [products](https://superdarn.ca/data-products).

658 For supporting information, Defense Meteorological Satellite Program (DMSP) Special Sensor
659 Ultraviolet Spectrographic Imagers (SSUSI) was used to generate plots to understand the spatial
660 distribution of the particle precipitation (https://ssusi.jhuapl.edu/gal_Aur)(Paxton et al., 2002)
661 and other DMSP data was referred from
662 http://cedar.openmadrigal.org/static/experiments3/2013/dms/17mar13/plots/s16_13mar17_1.htm
663 #20 . MetOps2 particle precipitation (Yando et al., 2011) plots were generated using
664 <https://cdaweb.gsfc.nasa.gov/>

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