

# 1 Atmospheric Escape from Earth and Mars: Solar and Solar Wind Drivers 2 of Oxygen Escape

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## 9 Key points:

- 10 • We derive oxygen escape rates from Earth as a function of solar wind parameters from a  
11 reanalysis of Dynamics Explorer 1 data.
- 12 • O<sup>+</sup> escape from Earth varies by a factor of 2 as a function of the solar wind magnetic  
13 field component aligned with the Earth's field.
- 14 • Oxygen escape rates from Mars are not as sensitive to variations in the solar power  
15 components as those from Earth.

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## 17 Index Terms:

18 2756 Planetary magnetospheres (5443, 5737, 6033)

19 2784 Solar wind/magnetosphere interactions

20 2479 Solar radiation and cosmic ray effects

21 2162 Solar cycle variations (7536)

22

## 23 Keywords:

24 Atmospheric escape, Terrestrial planets, solar wind energy, solar irradiance, Mars, Earth

## 25 Abstract:

26 Habitability at the surface of a planet depends on having an atmosphere long enough  
27 for life to develop. The loss of atmosphere to space is an important component in assessing  
28 planetary surface habitability. Current models of atmospheric escape from exoplanets are not  
29 well constrained by observations. Atmospheric escape observations from the terrestrial planets  
30 are available in public data archives. We derive oxygen escape rates from Earth as a function of  
31 solar wind parameters from a reanalysis of Dynamics Explorer 1 data and compare them to  
32 similar data from Mars. The reanalyzed Earth data are consistent with prior reports of escape  
33 from Earth as a function of geomagnetic indices. The reanalysis shows a dependence on the  
34 angle between the solar wind magnetic field and the Earth's magnetic field. It also  
35 demonstrates that oxygen escape rates from Mars are not as sensitive to variations in solar  
36 power components as those from Earth.

37

## 38 Plain Language Summary

39 Habitability of a planet depends on having an atmosphere long enough for life to  
40 develop. NASA and ESA data archives contain information about atmospheric escape from the  
41 terrestrial planets. For these planets oxygen ions dominate atmospheric escape. The data  
42 archives are just beginning to be analyzed and presented in a form that allows comparison with,  
43 and validation of, models of the interaction of stellar winds with exoplanets. We derive oxygen  
44 escape rates from Earth as a function of solar power components from a reanalysis of Dynamics  
45 Explorer 1 data and compare them to similar data from Mars. Our analysis demonstrates that  
46 oxygen escape rates from Mars are not as sensitive to variations in the solar power components  
47 as those from Earth. These data and similar data from Venus will prove to be important  
48 constraints on models of stellar wind/atmosphere interactions and atmospheric escape from  
49 exoplanets.

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## 51 1 Introduction:

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53 The ability of a planet to retain an atmosphere influences whether water can be stable  
54 as a liquid at the planet's surface. Understanding the evolution of planetary atmospheres  
55 requires knowledge of atmospheric escape for different star and planet combinations. A  
56 planet's atmospheric state is the result of source, loss, and modification processes that have  
57 acted on the atmosphere over time. The loss of atmosphere to space is an important  
58 component in assessing planetary surface habitability. Shizgal and Arkos (1996) reviewed the  
59 processes responsible for atmospheric loss from terrestrial planetary atmospheres and  
60 concluded: "The main mechanism for loss of atmospheric constituents appears to be the  
61 ionization of the neutral atmosphere by photoionization ...". Recently Cravens et al. (2017),  
62 Gunell et al., (2018), and others have suggested that dissociative recombination could produce  
63 significant escape of neutral oxygen from Mars.

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65 Escaping oxygen ion observations are available from the terrestrial planets. Recent  
66 reviews of observations and models of escaping ions at Mars (Brain et al. 2017), Venus (Persson  
67 et al. 2020) and Earth (Welling et al., 2015 and Yau et al. 2021) have elucidated the important  
68 role of induced (Venus and Mars) or intrinsic planetary-scale (Earth) magnetic fields in ion  
69 escape. Strangeway et al. (2010) first raised the question "Does a planetary-scale magnetic field  
70 enhance or inhibit ionospheric plasma escape?" Based on comparisons of total escape rates  
71 measured on Earth, Mars, and Venus, Strangeway et al. (2019) concluded that a planetary-scale  
72 magnetic field could enhance ionospheric escape. Gunell et al. (2018) and Ramstad and  
73 Barabash (2021) have reviewed oxygen loss measurements made at Mars, Venus, and Earth as  
74 a function of solar photoionizing flux and solar wind dynamic pressure and reached the same  
75 conclusion. Solar energy inputs, however, are not completely described by the solar wind  
76 pressure and the intensity of the ionizing radiation.

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78 Not only are the magnetic environments of the terrestrial planets different; the  
79 magnitude of solar energy inputs to their atmospheres driving ion escape vary widely. Ion

80 escape is powered by solar energy. This energy includes ionizing radiation directly from the Sun  
81 as well as kinetic and electromagnetic energy carried in the solar wind.

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83 The direct solar electromagnetic radiation below 90 nm ionizes the neutral atmosphere  
84 to produce the ionosphere from which ions escape. The importance of solar extreme ultraviolet  
85 radiation below 90 nm (EUV) and its variation parameterized by the solar  $F_{10.7}$  index in driving  
86 heavy ion escape is well established for Earth (e.g. Yau et al 1985, 1988, 2021). The relative  
87 importance of solar wind electromagnetic and kinetic energy in driving atmospheric escape is  
88 not well established. Recently Lockwood (2019) determined that the solar wind  
89 electromagnetic flux provides  $\sim 10\%$  of the power input to the Earth's magnetosphere and  
90 argues that it contributes to the plasma processes that drive geomagnetic activity and  
91 consequent ion escape.

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93 Measurements of the solar wind electromagnetic energy flux, i.e. the Poynting flux (S)  
94 and kinetic energy flux (K) have been available from solar wind monitors since 1970 (Feldman  
95 et al. 1978). These solar wind data were initially difficult to obtain and use so early Earth  
96 investigators focused on ground based geomagnetic indices such as  $K_p$ ,  $D_{ST}$ , and  $A_E$  which  
97 parameterize solar geomagnetic activity driven by changes in solar energy inputs (e.g. Yau et al.  
98 1988). More recently investigators have begun to parameterize atmospheric escape as a  
99 function of a single component of solar wind energy: the solar wind dynamic pressure (e.g.  
100 Ramstad and Barabash, 2021). Shillings et al. (2019) have reported escape fluxes as a function  
101 of coupling parameters that include elements used to calculate the Poynting flux. Earth-centric  
102 coupling parameters are not applicable to Mars. To our knowledge atmospheric escape has  
103 only been measured as a function of both solar wind electromagnetic and kinetic energy fluxes  
104 incident on Earth's magnetosphere by Lennartsson et al. (2004) and at Mars by Schnepf et al.  
105 (2024).

106

107 Lennartsson et al., (2004) hereafter called Lenn04 reported energetic  $H^+$  and  $O^+$  escape  
108 rates from the Earth's magnetosphere using data from NASA's Polar satellite for three years  
109 (1996-1998) during a solar minimum period as a function of Poynting flux (S) and solar wind  
110 kinetic energy flux (K). The data were further sorted for intervals where the solar wind magnetic  
111 field was parallel or anti-parallel to Earth's magnetic field. Data sampling limitations precluded  
112 an analysis of the escape rates as a function of solar ionizing radiation. They found increasing  $O^+$   
113 escape rates with increasing S and K intensity for both orientations of the solar wind magnetic  
114 field. The limited number of intervals where both escaping ion observations from the Polar  
115 satellite and solar wind data were simultaneously available compounded with the lack of data  
116 from a solar maximum interval limited the conclusions that could be drawn from the data.

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118 Our ultimate objective is to use escaping oxygen observations from terrestrial  
119 planets to partially validate models attempting to describe atmospheric loss from planets  
120 of other stars. Progress in understanding the influence of solar energy inputs on the ion  
121 escape from terrestrial planets requires comprehensive observations of all input power sources  
122 and escaping ion fluxes from each planet. A first step in this process is to use archived solar

123 wind and escaping ion data at Earth to expand the data base of escaping ions as a function of all  
124 solar drivers: solar ionizing radiation, Poynting flux (S) and kinetic energy flux (K) at Earth.  
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## 126 2 Data:

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128 The ion data we use to calculate escape rates are from the Energetic Ion Mass  
129 Spectrometer (EICS, Shelley et al. 1981) on the Dynamics Explorer 1 satellite (DE 1, Hoffman and  
130 Schmerling, 1981) acquired from September 1981 to February 1991 spanning Cycle 21 and 22.  
131 We use archived hourly average solar wind parameters from the IMP-8 satellite to calculate the  
132 solar wind power components S and K as well as the solar wind magnetic field direction. A  
133 model of the daily solar spectrum (Chamberlin et al. 2020) is integrated over two wavelength  
134 ranges to obtain measures of the solar ionizing radiation power:  $Q_{EUV}$  (0-45nm) and  $I_{EUV}$  (0-90  
135 nm).  
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### 137 2.1 Dynamics Explorer 1

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139 The DE 1 orbit has an apogee of 23,000 km, a perigee of 570 km and an inclination of  
140 90°. Its orbital plane has a local time drift period of 12 months and a line-of-apsides drift period  
141 of 18 months. 96-s cadence  $O^+$  flux and 1 sigma uncertainty based on the number of samples in  
142 15 broad energy bins sensitive to energies from 10 eV/q to 20 keV/q and 14 pitch angle bins in  
143 NASA's CDF format were acquired and processed. Each interval was characterized by five data  
144 quality flags. Only the highest quality data acquired above 6,000 km were used. Data for each  
145 96s interval were normalized to an altitude of 300 km and binned into 12 magnetic local time  
146 (MLT) and 16 Invariant latitude (INVL) bins.  
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### 148 2.2 IMP 8

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150 Selected one hour resolution solar wind parameters from the Interplanetary Monitoring  
151 Platform 8 (IMP 8, Feldman et al. 1978 and references therein) were obtained and processed.  
152 Specifically, we used the solar wind kinetic energy flux density (K) and the solar wind  
153 electromagnetic energy flux density (S, Poynting flux)  
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$$155 \quad K = \frac{1}{2}(M_h N_h + M_{he} N_{he})V^3 \quad (1)$$

$$156 \quad S = VB^2 \cos^2 \theta / \mu_0 \quad (2)$$

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158 where  $M_h$  and  $M_{he}$  are the  $H^+$  and  $He^{++}$  atomic mass,  $N_h$  and  $N_{he}$  the  $H^+$ , and  $He^{++}$  number  
159 density,  $V$  the solar wind velocity,  $B$  the IMF magnitude and  $\theta$  the angle between the solar wind  
160 velocity and magnetic field (Maggiolo et al, 2022).  $\mu_0$  is the magnetic permeability of free space.  
161 One-hour resolution magnetic field data in GSM coordinates were used to calculate the IMF  
162 clock and cone angles. The clock angle is the projection of the magnetic vector onto the  $B_y B_z$   
163 plane. The cone angle is the projection of the magnetic vector onto the  $B_x B_y$  plane.

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The IMP 8 satellite had a 30  $R_E$  orbit and so is often in the magnetosphere where reliable calculations of solar wind K and S are impossible. There is a magnetospheric flag included in the NASA data set indicating when IMP 8 is not in the magnetosphere. Unfortunately, the magnetospheric flag is not available after Jan 1, 1988. We selected only data where the magnetospheric flag was 6 or greater indicating the highest quality S and K values. Following Lenn04, parameters obtained from the IMP 8 satellite were time shifted 15 minutes to account for solar wind travel time to the magnetosphere.

### 173 2.3 Solar Ionizing Radiation

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The  $F_{10.7}$  radio flux has been a standard proxy for solar EUV irradiance due to its ease of ground-based measurements and the long historical daily dataset ranging back to 1947. More recent studies have shown the break in the  $F_{10.7}$  index and EUV irradiance relationship, especially for higher solar activity (Schonfeld et al., 2015, 2019). This is due to the  $F_{10.7}$  solar radio flux being comprised of two distinct formation mechanisms, namely the bremsstrahlung radiation, which does correlate well with EUV irradiance, and the gyromagnetic radiation, which has no relation to the EUV irradiance; therefore, the  $F_{10.7}$  index is not a pure measurement of solar EUV irradiance and fails to accurately represent its variability at times when dominated by the second component.

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Rather than using the solar  $F_{10.7}$  index to parameterize the rates of escaping oxygen ions, use integrals of the solar irradiance spectrum over two different wavelength ranges. Daily estimates of the solar irradiance spectrum at 1-nm resolution below 190 nm are available from the Flare Irradiance Spectral Model Version 2 (FISM2, Chamberlin et al. 2020). This model uses a combination of observations and solar indices to construct daily average solar spectra (0-190 nm) from 1947 to present.

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We use the integral over the spectral range that produces ionization in the atmosphere (i.e. below 90 nm) and designate it  $I_{EUV}$ . Schnepf et al. (2024), Dong et al. (2023) and others have used this index to parameterize oxygen ion escape from Mars.

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We also use an integral over the wavelength range that produces photoelectrons that in turn ionize neutral constituents of the upper atmosphere. The integral of the solar spectrum below 45 nm is identified as  $Q_{EUV}$ . Strickland et al. (2004) and Meier et al. (2015) demonstrated that  $Q_{EUV}$  values determined from integrating solar spectra agreed well with those determined from remote sensing of the O/ $N_2$  ratio in Earth's airglow.

### 202 2.4 Data Processing

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Data from DE 1 and IMP 8 are not continuous. There are 348,281 intervals with high quality DE 1 observations above 6,000 km. Of these 70,568 have simultaneous IMP 8 observations suitable for calculating S and K. Figure 1 panels A, B, and C display respectively

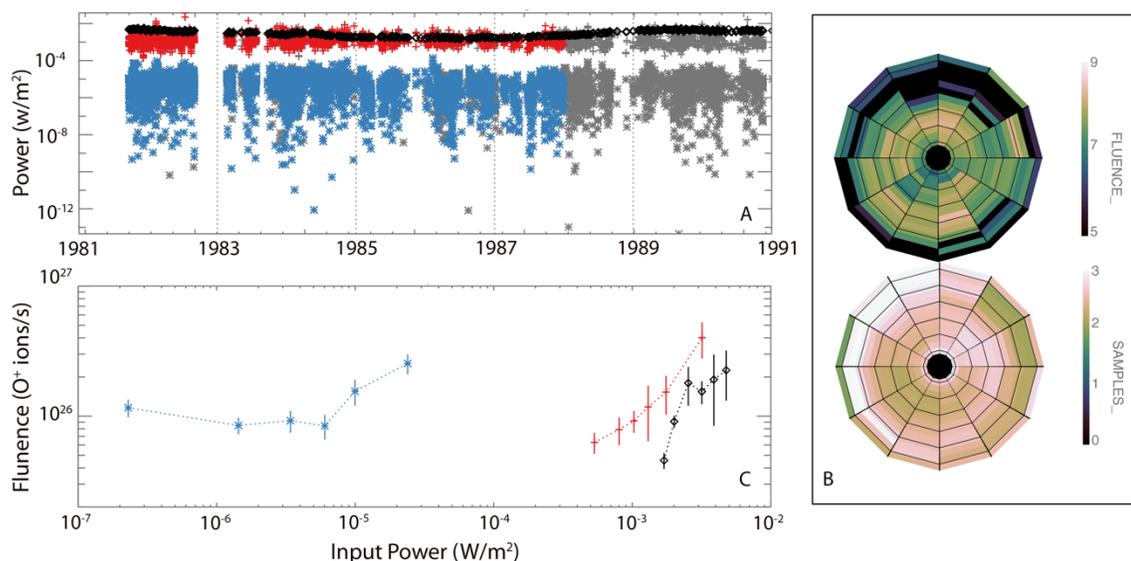
207 the input, intermediate, and final data. Figure 1B shows how the individual DE 1 escaping O<sup>+</sup>  
208 observations normalized to 300 km are binned in 14 INVL bins between 58 and 90 degrees and  
209 12 two-hour MLT bins. The software calculating INVL available at the time the DE 1 data were  
210 archived did not return INVL values greater than 85°. Peterson et al. (2008) have shown that  
211 the net escaping O<sup>+</sup> from the polar cap is negligible. Figure 1B shows that sampling in  
212 INVL/MLT space was non-uniform. The maximum and minimum number of samples in an  
213 individual bin were 2,389 and 83 respectively. Integrating the map of escaping O<sup>+</sup> fluxes in  
214 Figure 1B, and doubling it for the two hemispheres, gives  $1.3 \times 10^{26} \pm 8.3 \times 10^{25}$  O<sup>+</sup> ions per  
215 second escaping the Earth. The escape rate is consistent with the results of Yau et al. (1988)  
216 who used a lower resolution, preliminary version of the EICS data.

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218 To determine the variation in net O<sup>+</sup> escape rates as a function of solar energy inputs,  
219 the data shown in Figure 1B were further sub-divided into six S, K, I<sub>EUV</sub>, or Q<sub>EUV</sub> INVL/MLT maps  
220 with approximately equal numbers of samples. This division provides optimum sampling and  
221 resolution for our purposes.

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223 The O<sup>+</sup> escaping fluence and estimated uncertainty as a function of S, K, and Q<sub>EUV</sub> shown in  
224 Figure 1C are obtained by projecting the fluence at each INVL/MLT bin for each map onto a  
225 spherical surface at 300 km altitude (6670 km geocentric radius) and then integrating over the  
226 surface. Median values of S, K, and Q<sub>EUV</sub> for each of the six levels used in each map are shown  
227 in units of W/m<sup>2</sup>. We note that if the escape rate were independent of S, K, or Q<sub>EUV</sub> the escape  
228 rate would be constant and equal to  $1.3 \times 10^{26}$  ion/s obtained from the average map shown in  
229 Figure 1B. Figure 1C shows variations of O<sup>+</sup> escape with an increase of each of these variables.  
230 There appears to be a threshold value of the hourly average Poynting flux (S) above which it has  
231 an impact on O<sup>+</sup> escape. The O<sup>+</sup> escape rate appears to depend exponentially on the hourly  
232 average kinetic energy flux (K). The O<sup>+</sup> escape rate increases more rapidly at the lower values  
233 of Q<sub>EUV</sub>. The dependence of escaping O<sup>+</sup> as a function of I<sub>EUV</sub> is almost identical to that of Q<sub>EUV</sub>  
234 and not shown in Figure 1C.



235 Figure 1. Input (A), intermediate (B) and final (C) data. Panel A: Input solar wind Poynting flux S  
 236 indicated by \* symbols and kinetic energy flux K indicated by + symbols. Intervals where high-  
 237 quality S and K values are available are shown as blue and red symbols respectively; the other  
 238 intervals are shown as grey symbols. Solar irradiance integral powers  $Q_{EUUV}$  and  $I_{EUUV}$  are shown  
 239 as black diamonds and X's respectively. Panel B: Top: DE 1 escaping  $O^+$  fluxes in units of the  
 240 logarithm of ions/cm<sup>2</sup>-s normalized to 300 km are binned in 14 INVL bins between 58 and 90  
 241 degrees and 12 two-hour MLT bins. Bottom: logarithm of the number of data samples in each  
 242 MLT-INVL bin. C:  $O^+$  escaping flux and estimated uncertainty as a function of S (\*), K (+), and  
 243  $Q_{EUUV}$  (diamond). Uncertainties are shown as vertical lines.

### 245 3 Discussion

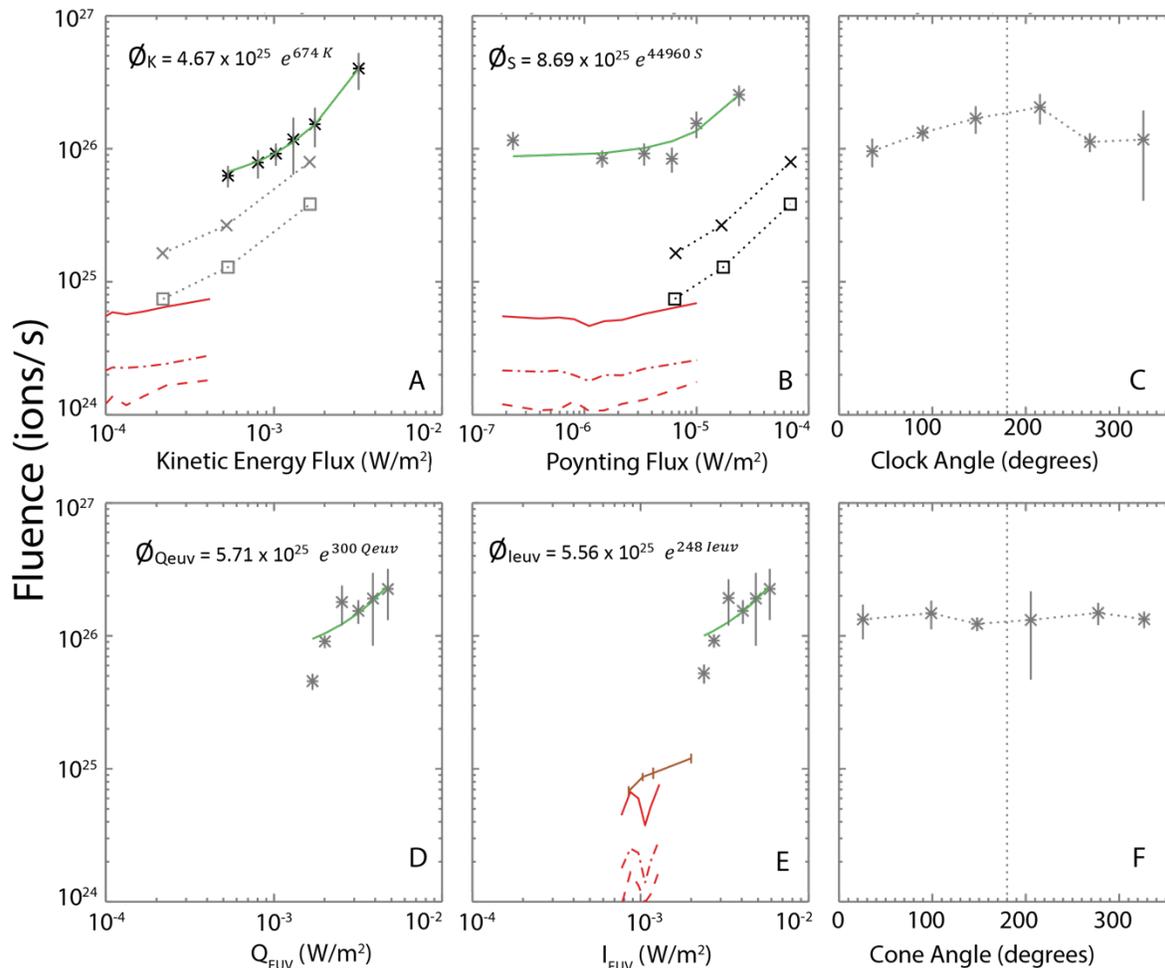
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 247 We are interested in comparing ionospheric escape rates from the terrestrial planets as  
 248 a function of all solar power inputs. NASA and ESA have supported multiple missions that have  
 249 observed plasma escaping the terrestrial planets. The available data on oxygen ion escape rates  
 250 as a function of all solar power inputs, however, is limited. Specifically, published data on the  
 251 effect of variations in solar wind Poynting flux (S) on atmospheric escape is comprehensive for  
 252 only Mars (Schnepf et al. 2024), partially available for Earth (Lenn04), and non-existent for  
 253 Venus. Strangeway et al. (2005) and others have correlated outflowing ions with the locally  
 254 measured Poynting flux. We do not use the Poynting flux observed at DE 1 which is different  
 255 than that observed in the solar wind (Rodríguez-Zuluaga, et al. 2022)

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 257 Data obtained on the DE 1 satellite from NASA's archives have been reprocessed to  
 258 produce estimates of the variation of  $O^+$  escape from Earth as a function of components of solar

259 power input. The results of this reanalysis are shown in Figure 1C which presents the rate of  
260 escaping  $O^+$  as a function of  $S$ ,  $K$ , and  $Q_{EUV}$ . The components of solar and solar wind power per  
261 unit area incident on the Earth and its magnetosphere vary over 5 orders of magnitude over the  
262 interval investigated.  $Q_{EUV}$  and  $I_{EUV}$  are used here to parameterize variations in the power input  
263 from solar ionizing radiation.  $I_{EUV}$  is included for comparison with the Mars data reported in  
264 Schnepf et al. (2024) and Dong et al. (2023). We note that the Schnepf et al., data in Figure 2E  
265 do not include data from the early mission which had the highest values of  $I_{EUV}$  included in the  
266 Dong et al. data.

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268 The three components of incident power shown in Figure 1C have different entry paths  
269 to the magnetosphere and affect  $O^+$  escape rate differently. The power in solar irradiance ( $Q_{EUV}$   
270 and  $I_{EUV}$ ) is transferred to the ionosphere below about 500km whereas the power in both  $K$  and  
271  $S$  is transferred to a much larger area: i.e. the Earth's magnetopause at altitudes above about  
272 60,000 km. The dependence of the ion escape is different for these parameters. Poynting Flux  
273 ( $S$ ) appears to have a threshold value below which it has no effect on the oxygen escape rate.  
274 The oxygen escape rate as a function  $Q_{EUV}$  and  $I_{EUV}$  (not shown in Figure 1) has different rates of  
275 increase at low and high values of  $Q_{EUV}$  and  $I_{EUV}$ . In addition, the ratios of maximum to minimum  
276 escape rates over the  $Q_{EUV}$ ,  $S$ , and  $K$  parameter ranges are different being 4.9, 3.0, and 6.4  
277 respectively.

278  
279 The Earth's large scale magnetic field has a strong influence on the paths available to  
280 incoming solar wind kinetic and electromagnetic energy fluxes. Lenn04 reported a factor of 2  
281 variation of  $O^+$  escape rates as a function of  $S$  and  $K$  for positive and negative values of IMF  $B_z$ .  
282 The reprocessed DE 1 data have a similar dependence on IMF  $B_z$  as shown below.



283  
 284 Figure 2. Escaping Earth oxygen rates as a function of S, K, Solar wind clock angle,  $Q_{EUV}$ ,  
 285  $I_{EUV}$ , and solar wind cone angle from this analysis, Schnepf et al. (2024), Dong et al. (2023), and  
 286 Lenn04. Data from Earth are shown in black and green lines and symbols; data from Mars are  
 287 shown in red and brown. Data from this analysis are shown as stars with associated  
 288 uncertainties shown as vertical lines. Exponential fits to the data are shown as solid green lines  
 289 with parameters of the fit in the upper left of the panel. The Mars data reported by Schnepf et  
 290 al. are shown in panels A, B, and E as red lines:  $O^+$  as dashed lines,  $O_2^+$  as dot dashed lines and  
 291 total O as a solid line. The Mars total oxygen escape data from Dong et al. (2023) are shown in  
 292 Panel E as a brown line with vertical red bars indicating the uncertainty. Apogee data above  
 293 50,000 km from Lenn04 are reported in panels A and B as X and box symbols. Intervals where  
 294  $B_z$  is strongly southward ( $B_z < -3nT$ ) are encoded by X symbols. Intervals where  $B_z$  is strongly  
 295 northward ( $B_z > 3 nT$ ) are encoded by box symbols. Earth  $O^+$  escape rates as a function of solar  
 296 wind clock and cone angles are presented in panels C and F.

297  
 298 Figure 2 presents  $O^+$  escape rates as a function of K, S,  $Q_{EUV}$ ,  $I_{EUV}$ , IMF clock and cone  
 299 angle. These data are compared and contrasted with available observations from Earth and  
 300 Mars. Panels C and F show that the average oxygen escape rate varies only as a function of one  
 301 component of the solar wind magnetic field, the component aligned with the Earth's magnetic

302 field,  $B_z$ . Panel C shows that the peak escaping flux is near a cone angle of  $180^\circ$  ( $B_z$  negative,  $B_y$   
303 near 0) a factor of 2.1 greater than the minimum near  $0^\circ$ , consistent with Lenn04. It is well  
304 known that geomagnetic activity is strongly dependent on the direction and intensity of the  
305 interplanetary magnetic field (IMF) relative to the Earth's dipole axis ( $B_z$ ). Masunaga et al.  
306 (2013) have shown that the escape of oxygen from Venus does not depend on the IMF  
307 magnetic field direction. Dong et al. (2015, 2023), Schnepf et al. (2024) and others have shown  
308 that oxygen escape from Mars is best organized in Mars Solar Electric (MSE) coordinates  
309 defined by the solar wind electric field ( $-V \times B$ ) which is perpendicular to both the solar wind  
310 velocity and magnetic field. There are no observations of variations in oxygen escape from Mars  
311 as a function of the IMF magnetic field direction relative to any Martian geographic or magnetic  
312 direction nor is such a dependence expected. The data presented here affirm the conclusions of  
313 Gunnel et al. (2018), Strangeway et al. (2019), and Ramstad and Barabash (2021) that a  
314 planetary-scale magnetic field enhances ionospheric escape.  
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316 The range of K and S values sampled in Lenn04 shown in panels A and B differ from the  
317 ones sampled in the DE 1 data and reported here. Maggiolo et al. (2022) report one-hour  
318 average values for S and K obtained from the NASA archives. The variability of S and K in the  
319 1996-1998 interval investigated by Lenn04 is greater than that in the 1981-1991 interval  
320 reported here. Lenn04 reported values when K values were about  $\frac{1}{2}$  and S values about 2 times  
321 those reported in Figures 1C and 2A and B. The Lenn04 analysis was based on 15-minute  
322 average S and K values rather than the 1-hour values considered here. We conclude that  
323 sampling biases introduced by the limited intervals for which  $O^+$  escaping fluxes and solar wind  
324 S and K values were available are major contributors to the higher values of S and lower values  
325 of K sampled by Lenn04.  
326

327 Cully et al., (2003) noted that differences in the magnitude of estimated escape rates  
328 arise from many factors including differences in altitude and energy range sampled. The  
329 difference between the magnitudes of the average total  $O^+$  escape rates from DE1 and Polar in  
330 panels A and B are consistent with the analyses presented in Peterson et al., (2006) and  
331 Lenn04..  
332

333 For analysis of the dependence on S and K Lenn04 considered only intervals with  
334 extreme values of the  $B_z$  direction. In Figure 2B, there appears to be a threshold power of  $2 \times$   
335  $10^{-6} \text{ W/m}^2$  below which the escape rate is constant in the DE 1 data. The rate of increase in  $O^+$   
336 escape as a function of S above this value is similar for both the DE 1 data reanalyzed here and  
337 the values reported in Lenn04. Because of limited data, Lenn04 was forced to calculate total  
338 escape in only three S ranges. We conclude that the variation of  $O^+$  escape as a function of S  
339 and K reported here and by Lenn04 are not inconsistent.  
340

341 We next address the fits to S, K,  $Q_{EUV}$ , and  $I_{EUV}$  power input levels versus escaping oxygen  
342 rates shown on the panels A, B, D, and E of Figure 2. Exponential fits to the data are shown as  
343 solid red lines. The exponential fits for S and K show that the observed rate of escaping  $O^+$   
344 increases with increasing incident S and K power.  
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346 At Mars, the major escaping ions  $O^+$  and  $O_2^+$  have been reported as a function of  $S$ ,  $K$ ,  
347 and  $I_{EUV}$  by Schnepf et al. (2024) and of  $I_{EUV}$  by Dong et al. (2023). The total Mars oxygen mass  
348 escape is the sum of the  $O^+$  rate and twice the  $O_2^+$  rate. These values are consistent with other  
349 reported rates as a function of other parameters (e.g. Ramstad and Barabash, 2021 and Brain et  
350 al. 2017). The variation of escaping oxygen as a function of  $I_{EUV}$  from Mars shown in Figure 2D is  
351 more than a factor of two smaller than it is from Earth. Direct comparison of Mars and Earth  
352 results in panels A, B, and E illustrates that oxygen escape rates from Mars are not as sensitive  
353 to variations in the solar power components,  $S$ ,  $K$ , and  $I_{EUV}$  as those from Earth.

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355 The exponential fits for solar irradiance power inputs  $Q_{EUV}$  and  $I_{EUV}$  shown in Figure 2D  
356 and E do not completely characterize the variation. At low values of the irradiance power the  
357 rate of increase in escaping  $O^+$  is greater per unit of input power than at higher values. These  
358 threshold and saturation effects have not been previously reported. Yau et al. (1988) reported  
359 an exponential fit for the variation of escaping  $O^+$  as a function of solar power parameterized by  
360 the  $F_{10.7}$  index. Cully et al. (2003) also reported an exponential fit to the  $O^+$  escape rate as a  
361 function of the solar irradiance power index  $F_{10.7}$ . The full DE 1 data set, not the present one,  
362 limited to only the data intervals where observations of  $S$  and  $K$  are available could provide  
363 more insight into the dependence of the  $O^+$  escape rate on solar EUV power.

364  
365 The weaker response to solar drivers on Mars has been noted by Forbes et al. (2006).  
366 They compared the contemporaneous response of Earth's and Mars' thermospheres to the 27-  
367 day variation of solar flux parameterized by the  $F_{10.7}$  index. Their analysis considered the  
368 differences in ionospheric cooling processes in the atmospheres of Mars, Earth, and Venus.  
369 They found that the thermospheric temperature response at Mars to variations in  $F_{10.7}$  was  
370 about half of that at Earth. They also suggested that the thermospheric temperature response  
371 of Venus during intervals of enhanced  $F_{10.7}$  indices would be less than that of Mars and  
372 significantly less than that of Earth.

373  
374 Oxygen ion escape from Earth is a complex process. The suggestion of different  
375 efficiencies of solar irradiance power in driving  $O^+$  escape at low and high levels of irradiance  
376 power input at Earth is not unreasonable. The standard model of oxygen energization at Earth  
377 involves multiple processes acting from the ionosphere to the plasma sheet (Yau et al. 2021,  
378 Peterson et al. 1992, Miyake et al. 1993 and others) to create and provide the 10-eV energy  
379 required for an  $O^+$  ion to escape.

## 381 4 Summary and Conclusions

382  
383 It is not yet possible to observationally determine what planetary characteristics (e.g.  
384 mass, rotation rate, distance from the star, magnetic field, or atmospheric composition) most  
385 strongly influence atmospheric escape or retention. Analysis of similarities and differences in  
386 the oxygen escape rates from terrestrial planets as a function of  $S$ ,  $K$ , and  $Q_{EUV}$  place new and  
387 significant constraints on models of stellar wind/exoplanet interactions and atmospheric escape

388 from exoplanets. We can then use these partially validated models to extend these constraints  
389 to the interactions of exoplanet atmospheres to their stellar environments

390

391 The magnitude and composition of solar energy inputs to terrestrial atmospheres vary  
392 widely. To date only the study of Schnepf et al. (2024) of data from the MAVEN mission to Mars  
393 has considered the influence of all components of solar energy input in driving atmospheric  
394 escape. This paper extends the observational constraints on the influence of solar drivers of  
395 atmospheric escape to Earth by combining observations from DE 1 and IMP 8. The solar drivers  
396 include solar irradiance and energy in the solar wind. Here solar wind electrodynamic power  
397 (the Poynting flux ( $S$ )), solar wind kinetic energy flux ( $K$ ), and two measures of solar irradiance  
398 power  $Q_{\text{EUV}}$  (0 – 45 nm) and  $I_{\text{EUV}}$  (0 – 90 nm) are used. The direction of the interplanetary  
399 magnetic field was also considered. The results of the re-analysis of observations of escaping  
400  $\text{O}^+$  observed on the DE 1 satellite as a function of  $S$ ,  $K$ , and  $Q_{\text{EUV}}$  were presented in Figure 1C.  
401 Differences with a previous analysis by Lenn04 of the dependence of  $\text{O}^+$  escape rates on  $S$  and  $K$   
402 were identified in Figure 2 and discussed.

403

404 Lenn04 reported a factor of 2 variation of  $\text{O}^+$  escape rates for positive and negative  
405 values of IMF  $B_z$  which was confirmed in the reanalysis (Figure 2C).

406

407 Lenn04 did not directly compare the influence of solar irradiance on  $\text{O}^+$  escape rates  
408 because of limited data. Data in Figure 2D and E show that, contrary to the results of prior  
409 analysis, the  $\text{O}^+$  escape rate is not well characterized by a single exponential increase with  
410 increasing solar irradiance. There is a higher rate of increase at low values of  $Q_{\text{EUV}}$  and  $I_{\text{EUV}}$  (i.e. a  
411 threshold effect).

412

413 The data presented above from Earth are compared to the data obtained from Mars by  
414 Dong et al. (2023) and Schnepf et al. (2024) in Figure 2. Figure 2 shows that oxygen escape rates  
415 from Mars are not as sensitive to the solar power components as those from Earth. Specifically:

416

- 417 • The variation of oxygen escape from Mars as a function of  $I_{\text{EUV}}$  reported by Dong et al.  
418 (2023) is larger than the corresponding variation of escape from Mars as a function of  $S$  and  
419  $K$  but smaller than that observed at Earth.
- 420 • Oxygen escape increases at both Mars and Earth as a function of increasing  $K$  but at a  
421 significantly reduced rate at Mars.
- 422 •  $\text{O}^+$  escape at Earth as a function of  $S$  appears to have a threshold value of  $2 \times 10^{-6} \text{ W/m}^2$   
423 below which the escape rate is constant.

424

425 Data exist to extend these comparisons to include Venus (Persson et al. 2020) but have  
426 not yet been presented as a function of solar wind Poynting flux ( $S$ ), kinetic energy flux ( $K$ ), or  
427 solar ionizing radiation.

428

429 The data present here demonstrate that the processes most important in driving  
430 oxygen escape respond differently to solar drivers on Mars and Earth

431

432

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434

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441 determinations are better organized by  $Q_{EUV}$  than  $F_{107}$

442

### 443 Open Research

444

445 The data used in this investigation were obtained from public data archives. The DE 1  
446 data are described by Peterson (2005) and available from NASA CDAWEB (Candey and Kovolick,  
447 2024). The IMP 8 data from which the S and K indices are described by King and Papitashvili  
448 (2004) and Papitashvili and Candey (2024). NASA points to Papitashvili et al. (2020) for access to  
449 the data. We found them difficult to find there. We suggest using the data at the NASA web  
450 page (2024). The  $Q_{EUV}$  and  $I_{EUV}$  indices were derived from daily 1 nm resolution solar irradiance  
451 spectra obtained from FISM daily 1 hour data page on the LASP Interactive Solar iRadiance  
452 Datacenter (2024). The daily 1 nm resolution spectra are described by Chamberlin et al., 2020.

453

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