

Supplementary information

SI.1 Methods

SI.1.1 Meteorological data compilation

QML(QumaLai), TSH (TianshuiHai), SQH (ShiquanHe), ZD (ZhiDuo), MA (MangAi), GZ (GaiZe), WQ (WenQuan), QSH (QingshuiHe), and ZAD (ZaDuo) can be downloaded from the meteorological data service center, China (<http://data.cma.cn/>), whereas BLH (Beilu'He), KXL (Kaixinling), AD (Anduo), XD(XidaTan), and HSX(HuashiXia) were provided by the State Key Laboratory of Frozen Soil Engineering (SKLFSE), China (<http://sklfse.nieer.ac.cn/>). The approach to address the data gaps is described in ref. 1 and 2. We used linear interpolation to address the dataset gaps if they were less than 2 hours, and we used a method described in ref. 3 to address gaps greater than 2 hours but less than 1 day. Furthermore, we used an artificial neural network approach as described in ref. 4 to fill gaps greater than 1 day.

Air temperature was measured by a 2 m tower, and the air temperature sensor was GB10589–89 (Yueke, Chengdu, China). In addition, during 1995–2002, the air sensor was HMP45C (Campbell, Logan, UT 84321–1784, USA), and from 2003 to 2017, the sensor was replaced by HMP155A (Campbell, Logan, UT 84321–1784, USA). From 1975 to 2002, precipitation data were measured by ZGQX–95 (Zhongguoqixiang, Beijing, China), and from 2003 to 2017, precipitation data were measured by T200B (Geonor, Oslo, Norway).

Soil temperatures in the 0–50 cm and 50–100 cm soil layers prior to 1990 were measured manually once or twice per month using a mercury thermometer (with an accuracy of ± 0.2 °C). From 1990, all soil temperatures were measured using a string of thermistors made by the SKLFSE (Lanzhou, China). The soil water contents (SWCs) of the 0–50 cm and 50–100 cm soil layers prior to 1995 were calculated gravimetrically using the ratio of the water mass present to the oven–dried (75 °C, 24 hours) weight of the soil sample 1~3 times per year. Since 1995, SWC has been measured by a pF–meter sensor (GEO–Precision, Environmental Industry Companies, Ettlingen, D-76275, Germany).

All the above data obtained by the different sensors were calibrated by the Meteorological Data Service Center of the SKLFSE, China.

Following Muller's original definition, the maximum active layer thickness (ALT) is the maximum thawing depth in late autumn using a linear interpolation of the temperature of soil profiles between two neighboring points above and below the 0 °C isotherm, as previously described (5).

Growing degree day (GDD) is often used to quantify temperature or heat requirements for plant development and is the cumulative sum of air temperatures above a certain degree. GDD is a heuristic tool in phenology that is widely used in mechanism-based phenological models (6). Here, GDD is calculated as the cumulative sum of daily air temperatures above 0 °C following ref. 7.

SI.1.2 Soil property data compilation

All soil cores were taken as two pseudoreplicate cores within 0.5° × 0.5° (latitude and longitude) and from the same vegetation type. The nine locations (AD, QML, BLH, HSX, TSH, SQH, ZD, MA, and GZ) had 62 plots and 1078 cores that were studied in 1975 and 1978 and from 1995 to 2017 (26 years). The 5 locations of WQ, QSH, KXL, ZAD and XD had a total of 96 plots and 760 cores, which were studied from 2002 to 2017 (16 years).

For each soil core, samples from 0–100 cm were collected by a soil corer (diameter is 5 cm) at 0–5 cm, 5–10 cm, 10–15 cm, 15–20 cm, 20–30 cm, 30–40 cm, 40–50 cm, 50–100 cm, and for the soil between 100 cm to 500 or 550 cm, we used a motorized drill to collect the samples at 50 cm intervals. The samples were collected using a stainless steel ring cutter, with three replicates. The permafrost table was determined by the ice content of the core sampling (5, 8). Near the permafrost table, we collected soil samples every 5 cm in the upper and lower 50 cm ranges. All samples were marked and sealed in a 100 ml steel aluminum box, weighed, frozen at –15 °C, and brought back to the laboratory used for soil property analysis.

52 Soil pH was determined by amperometry (DJS-1C, Leizi, Shanghai, China). Five grams of fresh
53 soil was mixed with CaCl_2 (0.2 M) at a ratio of 1 part soil to 5 parts liquid, and then, the pH of
54 the suspension was measured after 1 hour of shaking for each soil sample of a given year.

55 Soil bulk density was calculated based on total fresh weight, soil water content and total soil
56 volume.

57 Soil organic carbon (SOC) of the air-dried soil samples was analyzed using the wet combustion
58 method, Walkley-Black modified acid dichromate digestion, FeSO_4 titration, and an automatic
59 titrator (9, 10).

60 Total nitrogen (TN) was titrated by the Kjeldahl method (Wu et al., 2016) before 2000, and
61 since 2001, it has been measured by an elementary analyzer (Vario EL Three, Elementar,
62 Germany). All the data transitions between the Kjeldahl method and elemental auto-analysis
63 followed the method described in the lecture of the Chinese ecosystem research network
64 (<http://cern.ac.cn>).

65 The ratio between soil carbon and nitrogen (C:N) was then calculated as the quotient of the
66 SOC and TN concentrations.

67 Chloride (Cl^-) was measured by AgNO_3 titration. Calcium (Ca^{2+}) and magnesium (Mg^{2+}) were
68 measured by EDTA titration. Sodium (Na^+) was measured by a flame photometer (FP6400,
69 JINGMI, Shanghai, China). For more details of the methods, please see the Chinese ecosystem
70 research network, China (<http://cern.ac.cn>), and for primary results, please see Fig. S13

71 Nitrate (N-NO_3^-) was measured by capacity titration (ELIT 8021, Beijing, China) prior to 2000;
72 since 2001, ion chromatography (T6-New Century Spectrophotometer, PUXI, Beijing, China) has
73 been used. The NO_3^- data obtained by the different methods were calibrated by the SKLFSE,
74 China.

75 Ammonium (N-NH_4^+) was analyzed based on soil samples (2008–2017) with the initial moisture
76 content. The soil samples were extracted by 2 M KCl and then quantified using a flow injection
77 analyzer (Autoanalyzer 3 SEAL, Bran and Luebbe, Norderstedt, Germany), and the limit of

78 detection was 0.003 N mg L⁻¹. The NH₄⁺ concentration was subsequently calculated as mg N kg⁻¹
79 dry weight (DW) of soil.

80

SI.1.3 Plant trait data compilation

For each plot, species identification, vegetation height, landscape types, characteristics of surface drainage and erosion status were recorded. Plant coverage was measured in the early or middle of September at 5 plots (1 m × 1 m) placed with a 500 m radius of the borehole position. In 1975 and 1978, it was measured based observations by experts. From 1995 to 2007, using the pin-point method, a 10 cm × 10 cm nylon mesh was laid in a 1 m × 1 m plot to estimate the plant coverage. From 2008 to 2017, a US agricultural multispectral camera was used to measure the plant coverage. The difference in plant coverage between the different methods was successfully validated by linear regression ($r^2 > 0.9$.)

Plant height was measured by a 10-point frame sample (wind speed < 0.5 m/s). Forty-five degrees slowly downward from the pin, when the pin first hit the target grass leaf, was recorded and calculated as the individual plant height, with 5 repeats. The average of 50 samples (10 points × 5 repeats) was recorded as the plant height. Finally, the mean height for the plot was calculated for any given year.

The aboveground biomass in September was quantified in three 20 cm × 20 cm subplots within each plot during September, for a total of 15 subplots per site. At each subplot, aboveground biomass was cut using scissors, packed in paper bags and returned to the laboratory, where all samples were filtered through a 1 mm × 1 mm sieve to remove soil, and plant samples were dried at 75 ± 2 °C for 72 hours to calculate the aboveground biomass.

To investigate the root vertical distribution pattern of the plant species in the community, a 1.5 m depth pit was dug in five 1.5 m × 1.5 m quadrats to measure the maximum root depth after all of the above work was finished. Since 2008, this approach has been replaced by a special soil sample drill device (diameter is 15 cm) designed by SKLFSE, China.

The vertical root distribution was based on visually observed fresh roots from the flow water immersion soil core and consisted of three replicates per plot. A mean value was calculated for the maximum root per plot.

Additional deep soil profile samples were collected in areas with retrogressive thaw slumps near study sites. Non of the actual study sites reveal such features. At these sites, roots were followed to the surface and associated with species-specific living plants. Maximum root lengths which could be linked to individual plants were recorded.



The photo shows a retrogressive thaw slump. Thawing permafrost and a collapsed landscape has exposed soil profiles to a depth of about 5 m. The main root zone was within the upper 0.5 m and the active layer depth 2.8 m. Species-specific roots were found at least 2.4 m below the surface (shown with an yellow dashed line) and related to *Kobresia littledalei* C. B. Clarke and *Oxytropis pauciflora* Bunge. The location is 55 km North of ZAD (ZaDuo) and described July 13, 2021.

122

123 **SI.1.4 Stable isotope ^{15}N data compilation**

124 To quantify if the ammonium released during permafrost thawing can be taken up by plants,
125 210 plots (including 30 control plots) of 6 sites for both group A (ZD, HSX, and XDT) and group B
126 (BLH, AD, and KXL) were set on 5th to 18th September 2017, which is close to the time of
127 maximum active layer depth and corresponds to the end of the aboveground growing season
128 and onset of plant senescence. N release from the permafrost was mimicked by injecting a
129 trace amount of the stable isotope ^{15}N into the soil at the depth of the thawing permafrost
130 front. At each plot isotopically labeled N, 1 g $^{15}\text{N}\text{-NH}_4\text{Cl}$, 99 atom%, was dissolved in 50 g
131 deionized water and wrapped in polyvinyl chloride capsule. To avoid interruption to the ground
132 surface environment, a capsule was launched into the thawing permafrost front by a hollow,
133 narrow slope steel drill (Supplementary Fig. S12), and then, a radar (precise: 0.02 m for 500
134 MHz; MALA ProEx, MALA, Sweden) was used to confirm the location of the capsule and plant
135 sample plot.

136 To assess the time-scale of isotope nitrogen uptake by plants, the dominant plant roots of the
137 0-50 cm soil layer and aboveground mass (including leaves and stems) were collected within
138 33×33 cm quarters immediately above the capsule after treatment for 10~15 days, one year,
139 two years, three years and four years. There were six primary dominant species at 6 sites,
140 including 3 deep-rooted species, which were *Carex moorcroftii* Falc. Ex Boott (*Carex*), *Kobresia*
141 *tibetica* Maxim. (*Kobresia*), and *Oxytropis glacialis* (*Oxytropis*), and 3 shallow-rooted species,
142 which were *Heteropappus bowerii* (Hemsl.) Griens. (*Heteropappus*), *Leontopodium*
143 *nanum* (Hook. f. et Thoms.) Hand.-Ma (*Leontopodium*), and *Saussurea arenaria* Maxim
144 (*Saussurea*). It was not possible to macroscopically separate roots for all plant species, but
145 species could be sorted as deep-rooted and shallow-rooted plant species by hand.

146 The stable isotope ^{15}N is naturally present in very low quantities in plants. By adding a trace
147 amount of ^{15}N , N uptake can be measured by determining the excess ^{15}N content compared to
148 the naturally occurring ^{15}N found in the total N pool. Thus, excess ^{15}N , or ^{15}N enrichment, was

calculated as the difference in atom% ¹⁵N between the samples from the treated (isotopically labeled) and untreated (natural abundance) plots:

$$atom\%_{excess} = atom\%_{sample} - atom\%_{nature\ abundance}$$

Atom% is calculated as:

$$atom\% = \frac{100 \times (\delta + 1,000)}{\delta + 1,000 + \frac{1,000}{R_{st}}}$$

where δ is the $\delta^{15}\text{N}$ for the sample (‰), and R_{st} is the $^{15}\text{N}/^{14}\text{N}$ ratio of the international reference material (0.003676). Samples with naturally abundant $\delta^{15}\text{N}$ were collected from the untreated plots at specific species and harvest times. In this study, potential seasonal changes in the natural abundance of $\delta^{15}\text{N}$ in plants and the significant levels of isotopic enrichment showed that the potential temporal variation in $\delta^{15}\text{N}$ was negligible. Excess ^{15}N is expressed as the concentration of excess ^{15}N per N in plant leaves and roots.

SI.1.5 Data quality control

We conducted a multistep hierarchical data-cleaning method (11) to provide quality control for the compiled data as follows:

First, soil property data and plant trait data were checked for improbable values, with the goal of excluding likely errors or measurements with incorrect units but without excluding true extreme values. We manually checked plant trait values and excluded only those that were obviously erroneous based on our expert knowledge of these species.

Second, for each case, identifying whether a given observation (x) was likely to be erroneous (that is, ‘error risk’) by calculating the difference between x and the mean (excluding x) of the taxon (species or site) and then dividing by the standard deviation of the taxon were conducted since the standard deviation of a trait value is related to the mean and sample size. We checked individual records against the entire distribution of the observations of trait data and removed any records with an error risk greater than 8 (that is, a value more than 8 standard deviations away from the soil property/plant trait mean value). For individual species across the sites, we estimated a species mean per location and removed observations for which the species (plot-

level) mean error risk was greater than 3 (that is, the species mean of that site was more than 3 standard deviations away from the species mean across all plots from the location). Finally, we compared individual species records directly to the distribution of the values at the plot level. We excluded values above an error risk of 2.25. After the above procedure was conducted, 124 data points (4.7%) were removed for soil properties, and 316 observations/data points (3.8%) were removed for plant traits. In all cases, we visually checked the excluded values against the distribution of all observations for each species to ensure that our trait and soil property data cleaning protocol was reasonable. Furthermore, we found that 124 excluded data points (4.7%) were removed because soil properties were mainly based on the soil surface pH value (71%) and soil surface NO_3^- data (22%), and 316 data points were removed because plant traits were mainly based on the aboveground biomass of September (86%), which was caused by waste or the activity of small mammals, such as *Ochotona curzoniae*.

SI.1.6 Spatial autocorrection and uncertainty estimates

Because the plots included in this study were not uniformly dispersed over the Tibetan Plateau, we examined the potential of locations, such as locations TM, TGL, KXL and BLH, in relatively concentrated plots to drive the overall patterns. To do this, we first examined spatial autocorrelation in the trends in weather parameters, soil properties and plant traits using Moran's I in the R package ltools (12). Moran's I test results ranged from -1 to 1, where -1 indicates a perfectly dispersed variable, 0 indicates a randomly distributed variable, and 1 indicates that the data are perfectly autocorrelated. We observed that the results ranged from 0.001 to 0.018, which suggests that in this study, some variables were weak but had no significant spatial autocorrelation.

We conducted an analysis to compare trends, confidence intervals, and significance of trends over two time periods, 1995–2005 ($n = 154$) and 2006–2017 ($n = 168$), to assess whether different locations during observation years influenced the overall trends in active layer warming, active layer moisture, permafrost thawing, soil properties (0-50 cm, 50-100 cm, and 100-permafrost front), and plant traits we observed. For each time period, we used a subset of

locations that had data for at least 80% of the years within the defined time period. Following the established methods of ref. 13, we calculated a yearly anomaly in the active layer, soil properties, and plant traits for each location as the difference between each year's observation and the long-term mean. We then averaged these anomalies across all locations and used 1) the probability density function to test for a normal distribution and 2) linear regression to calculate the slope, significance, and confidence intervals of these averaged anomalies.

SI.1.7 Climate, permafrost dataset temporal assimilation and group split

To generate a dataset with consistent time scale (year) resolution within and among climate characteristics, soil properties, and plant traits, we extracted and averaged air temperature, total precipitation, soil temperature, and soil moisture for the growing season (from ¹ May to ³⁰ September) and nongrowing season (from ¹ October to ³⁰ April) at specific locations.

To directly compare the different effects of permafrost conditions on plant community traits under climate warming, we used the observed ALT time trends (from 1975 to 2017), and the 14 study locations were split into two groups. Group A consisted of TSH, QSH, SQH, GZ, MA, WQ, ZAD, XD, ZD, and HSX, with significantly positive increasing trends (ALT significantly increased with time under climate warming). Group B consisted of AD and BLH without significant changes and KXL and QML with significantly negative trends (ALT significantly decreased with time under climate warming). To test whether the split standard was not an artifact of our group method, we applied a control plot and fitted it by linear least squares. The slope was 85.88, the correlation coefficient (r^2) was 0.8, and $p < 0.01$, which demonstrated the sensitivity of the changes in ALT in air temperatures.

SI.1.8 Characterizing trends in climate and permafrost

For each location, we calculated the mean ALT, mean soil temperature at 0-50 cm, mean soil temperature at 50-100 cm, mean soil moisture at 0-50 cm, and mean soil moisture at 50-100 cm in a given year for our defined growing season to obtain a mean annual value. These variables and the mean annual air temperature for each location were then used to calculate 1) the long-term trends for climate change, active layer warming, active layer moisture, and permafrost thawing and 2) group level (group A and group B) mean annual soil temperature at 0-50 cm, soil temperature at 50-100 cm, soil moisture at 0-50 cm, soil moisture at 50-100 cm

and variation in ALT. The difference between group A and group B was tested by a *t test*, and the significance level was 0.05. All time intervals were calculated using the Mann–Kendall test (trend package, R 4.1). We excluded datasets that had < 10 years. The control plots that measures the rate of air temperature per decade ($\delta T_{air}/\text{decade}$) and the rate of ALT per decade ($\delta \text{ALT}/\text{decade}$) were used to assess the sensitivity of permafrost thawing to climate warming and fitted by the linear least-squares method. The significance level was $p < 0.05$, and the confidence level was 95%.

SI.1.9 Characterizing trends in soil properties and soil nutrition

We calculated the mean soil bulk density, SOC concentration, TN concentration, C/N, Cl^- concentration, NO_3^- -N concentration, NH_4^+ -N concentration, Na^+ concentration, Ca^+ concentration, and Mg^+ concentration. For each location, we calculated the mean of the above properties for soil profiles (0-50 cm, 50-100 cm, and 100 cm permafrost front) in a given year to obtain a mean annual value. The definition for the 100 cm permafrost front, as it is not possible to precisely judge or distinguish by eye in the field, was the 100 cm permafrost front soil core that included 5~10 cm permafrost. Mean annual soil properties and soil nutrition were then used to calculate 1) the long-term trends for the 0-50 cm, 50-100 cm and 100 cm permafrost fronts and 2) the variation in soil properties and soil nutrition for group A and group B. The difference between group A and group B was tested by a *t test*, and the significance level was 0.05. All trends were calculated by linear regression, and the slope was estimated by the linear least-squares method. The significance level was $p < 0.05$, and the confidence level was 95%.

SI.1.10 Characterizing trends in plant traits

For each location, we calculated the mean plant coverage, aboveground biomass in September, root maximum length, mean plant height (September), and plant community species richness in a given year (in our defined growing season period) to obtain a mean annual value. These variables were then used to 1) calculate the long-term trends for individual locations and 2) calculate the group-level plant trait time trend for group A and group B. All trends were calculated using linear regression and fitted by the least-squares method; the significance level was $p < 0.05$, and the confidence level was 95%.

SI.1.11 Linkage between permafrost thawing and nitrogen release

At ambient sites, linear regression was used to identify 1) the nitrogen source that was the major contributor to the 50-100 cm NO_3^- concentration variation, as we hypothesized that 0–50 cm was a proxy for the nitrogen source from active layer warming and 93% SOC was attributed to the 0-50 cm and 100 cm permafrost front proxies for the nitrogen source from permafrost thawing, and 2) the relationship between the variation in the NO_3^- concentration within the active layer (0-50 cm, 50-100 cm, and 100-permafrost front) and permafrost thawing (variation in ALT). The r^2 and significance level were calculated by the least-squares method. The significance level was $p < 0.05$, and the confidence level was 95%.

To assess the robustness of the inferred enhancement in NO_3^- concentration within the active layer from permafrost thawing, we used a contour map to delineate the variation in nitrate concentration (NO_3^-) within the active layer with permafrost thawing change (variation in ALT) from 1995 to 2017 (SF 15; Surfer 12.0). Kriging was used to grid NO_3^- concentration data at 10 cm intervals within 0-500 cm during 1995 to 2007; x axis is year, y axis is depth of active layer, Z axis is NO_3^- , Z axis is linear, there is no transform and inflate conveys hull by 0.

SI.1.12 Causality analysis between maximum root length and nitrogen increase at 50-100 cm

To demonstrate the causal relationship between the maximum root length increase and the nitrogen concentration increase of 50-100 cm, we ran the causality test by convergent cross mapping with the rEDM package of R 4.1. The optimal value of the embedding dimension E was 2, estimated by the function of the SSR pred boot. Convergent cross mapping demonstrated that an increase in soil nitrogen of 50-100 cm caused the maximum root length change (increasing), and the time lag was approximately 2 years. This result was consistent with field observations of isotope ^{15}N where labeled N rejected to the front of permafrost could be taken up by the deep-rooted plant in two years. Furthermore, convergent cross mapping was used to improve our understanding of the relationship between climate change or permafrost thawing or active layer warming or active layer soil moisture and maximum root length. The results were assessed by the r^2 and p value, with a significance level <0.05 .

SI.1.13 Multiple regression analysis of drivers of plant trait trends

We conducted a multiple regression analysis of the air temperature, total precipitation, active layer temperature, active layer moisture, variation in ALT, soil nitrogen variation within the active layer (variation in N-NO_3^- concentration and N-NH_4^+ concentration), and drivers of observed plant trait (maximum root length, species richness, and aboveground biomass in September) trends at each location. Predictors in the analysis included growing season air temperature, growing season total precipitation, growing season soil temperature of 0-50 cm, growing season soil temperature at 50-100 cm, growing season soil moisture at 0-50 cm, growing season soil moisture at 50-100 cm, growing season soil nitrogen (including N-NO_3^- and N-NH_4^+) at 0-50 cm, growing season soil nitrogen at 50-100 cm, growing season soil nitrogen at the 100 cm permafrost front, variation in ALT, nongrowing season air temperature, nongrowing season soil temperature at 0-50 cm, nongrowing season soil temperature at 50-100 cm, annual air temperature, annual soil temperature at 0-50 cm, and annual soil temperature at 50-100 cm.

All variables were standardized by Z-scores to facilitate a comparison of model coefficients across the variables with different units. Selection variables were entered stepwise. We verified that multicollinearity was not a problem by checking that the variance inflation factor was well below ten for all variables (14). We used the leaps R package to select subset models including all predictors and two-way interactions and selected the fitted model having the lowest Akaike information criterion (AIC). The results are described by the coefficient (r^2) and p value, and the significance level <0.05 .

SI.1.14 Structure equation model (SEM)

Piecewise SEM was examined to identify 1) the pathway through which climate change potentially affects plant growth and 2) the difference between the direct and indirect effects of temperature, water balance and soil nutrition on plant growth. The mean value at the site level for the growing season was used in the SEM analysis and split into two groups, A and B (supplementary Fig. S1). Variables that only demonstrated a significant correlation ($p<0.05$) with plant traits in the multiple regression analysis were pooled into the SEM. Nutrition variation was represented by the nitrogen variation, whereas ammonium (NH_4^+) variation

317 delineated the nitrogen released from the permafrost. Ultimately, 7 variables were used in the
318 SEM.

319 All variables were standardized (mean zero, unit variance) using Z-scores. Then, principal
320 component analysis (PCA) was used to summarize the structure between plant growth and
321 driver parameters. We assumed linear Gaussian relationships between variables included in the
322 model, which tested normality with density plots for each variable (15).

323 As plant uptake of nitrogen released from permafrost has a two-year time lag, the final climate
324 change and plant trait dataset was from 1997 to 2017, and the soil nitrogen dataset was from
325 1995 to 2015. We fit separate models window-by-window from 0 to 5 years both for the
326 growing season and nongrowing season to 1) test whether the time lag of 2 years in the SEM
327 was not an artifact setting and 2) account for possible time lag effects of climate change
328 variables (i.e., nongrowing season air temperature and soil temperature) and nitrogen released
329 from the permafrost thawing front during plant aboveground senescence.

330 To examine pathways by which climate change potentially affects plant growth and to
331 differentiate between direct and indirect effects of temperature (air temperature, growing
332 degree days, soil temperature at 0-50 cm, soil temperature at 50-100 cm, and variation in ALT),
333 water balance (soil moisture at 0-50 cm and soil moisture at 50-100 cm) and soil nutrition (N-
334 NO_3^- and N-NH_4^+ concentration in that 0-50 cm, 50-100 cm, and 100 cm-permafrost front) on
335 plant growth, we used piecewise SEMs.

336 Before the SEM analysis, all variables for the 14 locations were 1) split into two groups based on
337 the variation in ALT under climate warming, and the details are presented in Supplementary
338 Fig. S1 and 2) only demonstrated a significant correlation with plant traits ($p < 0.05$) in the above
339 multiple regression analysis and were pooled into the SEM. For each group of variables, the
340 mean value at the level of location was used in SEM analysis. As the plant uptake of nitrogen
341 released from permafrost has a two-year time lag, the final climate change and plant trait
342 dataset was from 1997 to 2017, and the soil nitrogen dataset consisted was from 1995 to 2015.
343 Then, we used maximum likelihood estimation to interpolate the deficiency of data (16).
344 Considering regional climate and using a priori knowledge, we included only growing season

climate variables in our SEM, and nongrowing season and annual climate variables were excluded from the analyses. Although plant growth relationships with dormant season climate are occasionally reported, this information was not included in the SEM analyses.

After the data were compiled, a conceptual model was constructed to investigate complex (i.e., direct and indirect) relationships among plant growth and temperature, water balance, and soil nitrogen of both groups of responders. Soil moisture at 0-100 cm was used as a surrogate for the water balance among precipitation, ice melt, unfrozen water, and evaporation in the model. Because the maximum root length ranged from 50 cm to 100 cm for group A, the soil moisture and soil temperature at 0-100 cm were both chosen. For group B, the maximum root length ranged from 0 cm to 50 cm, and the soil moisture and soil temperature both ranged from 0-50 cm. To explore the potential mechanisms behind plant growth responses associated with changes in temperature, moisture and nutrition, the same SEM structure variables were applied to groups A and B.

As the nitrogen variation was mainly caused by active layer warming and permafrost thawing, we grouped the variation in nitrogen into two parts: 1) one part was related to active layer warming, as the ammonium (NH_4^+) data were sampled from 2008 to 2017 and nitrate (NO_3^-) data were sampled from 1995 to 2017, the nitrate (NO_3^-) variation at 0-100 cm accounted for 73% of the total nitrogen variation at 0-100 cm, and to avoid the uncertainty in data weight, here only the nitrate (NO_3^-) concentration in the 0–100 cm (averaged by NO_3^- concentrations at 0–50 cm and 50–100 cm) layer was used for further analysis with the data on the ammonium (NH_4^+) concentration in the 0–100 cm layer being excluded; 2) the other part was related to permafrost thawing. NH_4^+ variation accounted for 93% of the nitrogen variation in the 100 cm permafrost front. To avoid uncertainty in data weight, only the ammonium (NH_4^+) variation was included to delineate the nitrogen released from permafrost.

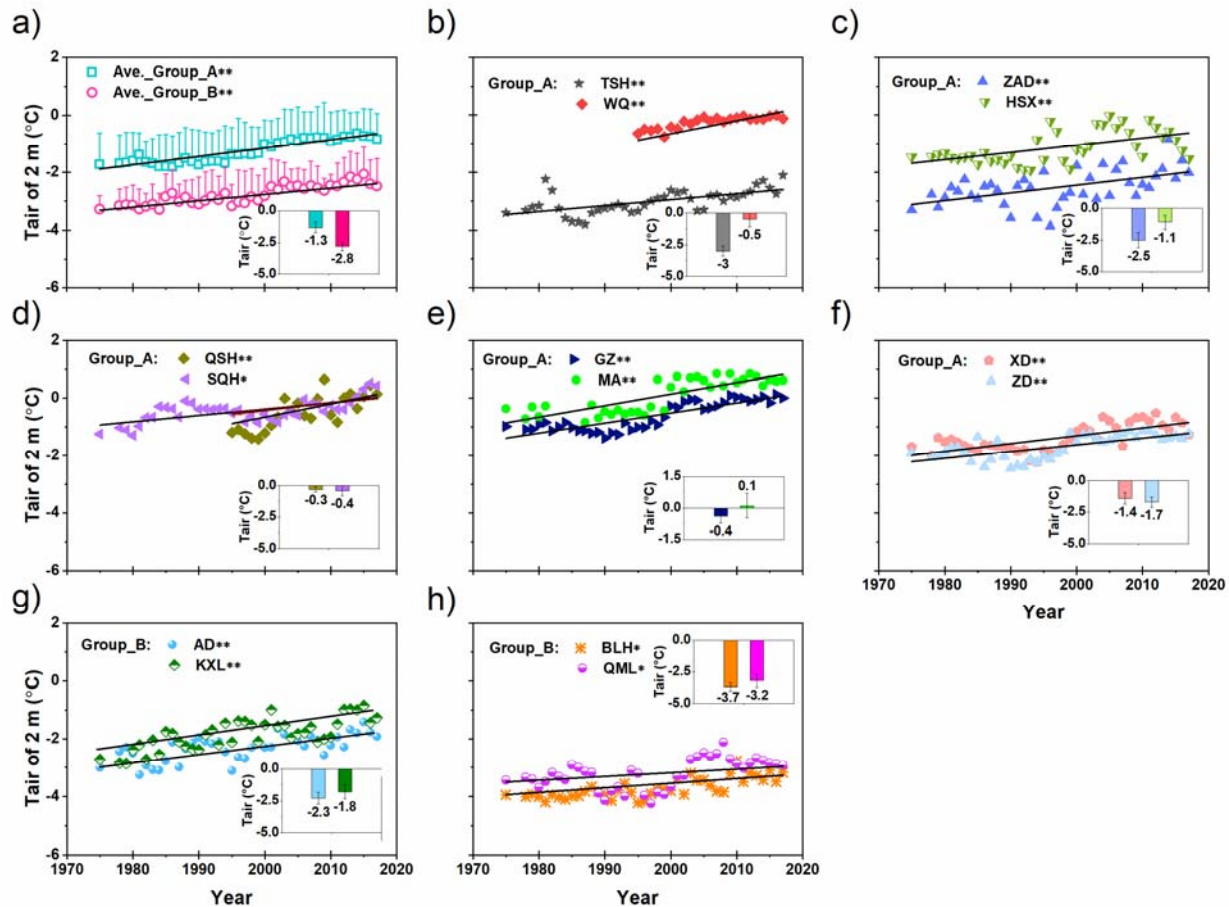
In the end, 7 variables were used in the SEM: (1) the mean air temperature (MAT) during the plant growing season; (2) the mean soil water content at 0–100 cm (SWC; averaged by 0–50 and 50-100 cm) during the growing season; (3) growing degree days, which is a cumulative sum of daily air temperature above 0 °C; (4) active layer warming, which was proxied by the mean

soil temperature of 0–100 cm (T_{soil} ; averaged by soil temperature at 0–50 and 50–100 cm) and NO_3^- concentration at 0–100 cm; (5) permafrost thawing, which was proxied by the maximum ALT and NH_4^+ concentration of the 100 cm–permafrost front; (6) plant growth, including aboveground biomass in September and species richness; and (7) the maximum length of roots, which was the maximum root depth.

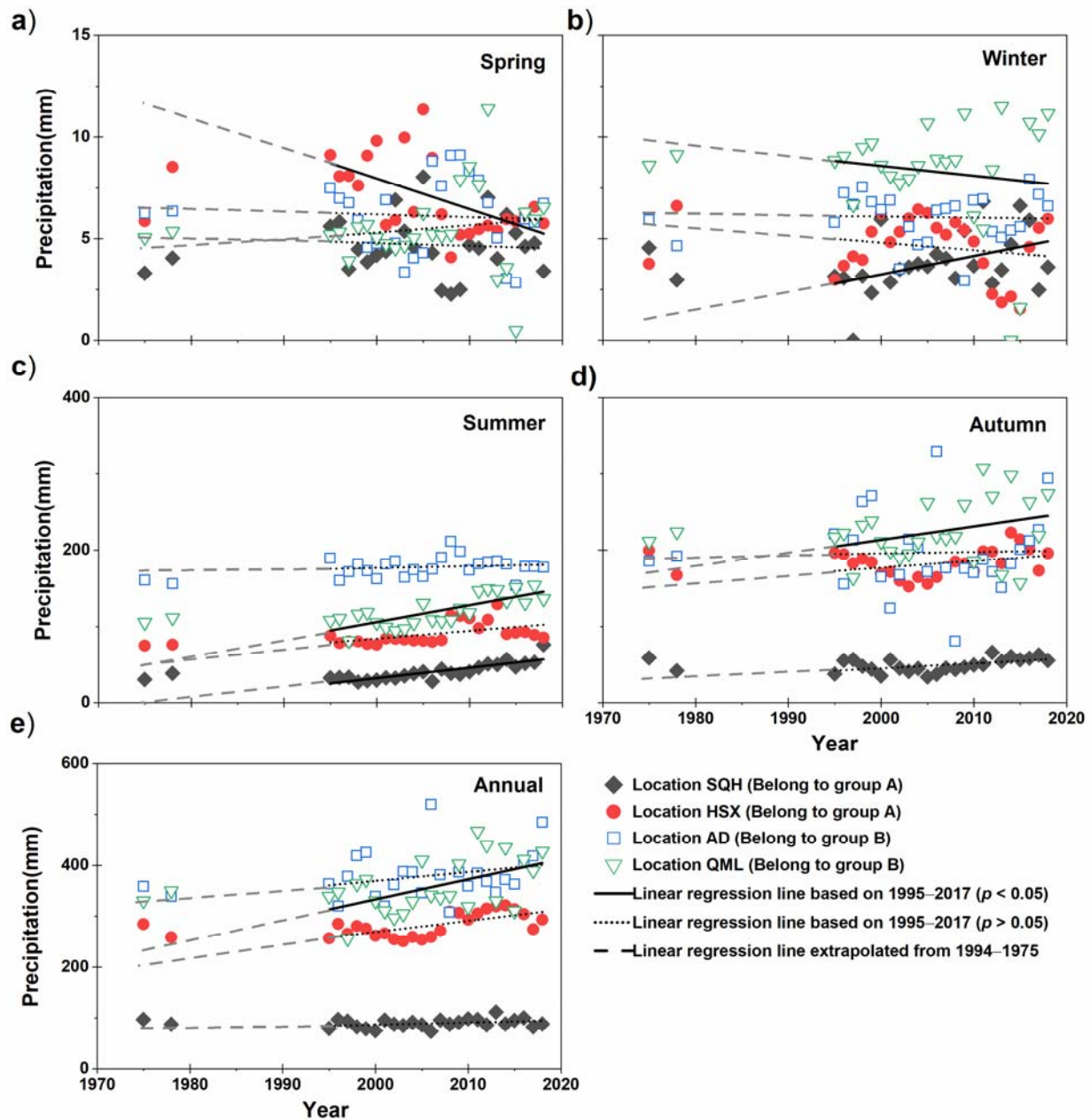
To test whether the time lag (2 years) in the SEM was not an artifact and account for possible time lag effects of climate change variables (i.e., nongrowing season air temperature and soil temperature) and the nitrogen released from the permafrost thawing front when plants experienced aboveground senescence, we fit separate models window-by-window from 0 to 5 years both for the growing season and nongrowing season.

We then created an expression of community structure compatible with the SEM. We assumed linear Gaussian relationships between the variables included in the model that tested each variable for normality with density plots. All variables were standardized (mean zero, unit variance) using Z-scores. Then, PCA was used to summarize the structure between plant growth and climate drivers by FactoMineR packages of R 4.1. All variables were entered before SEM analysis to facilitate the interpretation of parameter estimates.

To test the goodness-of-fit of our SEMs, we used chi square (χ^2), degrees of freedom ($d.f.$), and the root-mean-square error of approximation (RMSEA; the model had a good fit when the RMSEA was indistinguishable from zero to 0.05). A path coefficient is analogous to the partial r^2 or regression weight and describes the strength and sign of the relationships between two variables (16). We calculated the standardized total effects of all drivers on the plant growth of active layer warming and permafrost thawing attributes. The net influence that one variable had upon another was calculated by summing all direct and indirect pathways (effects) between two variables. All SEM analyses were conducted using the piecewise SEM package of R 4.1.

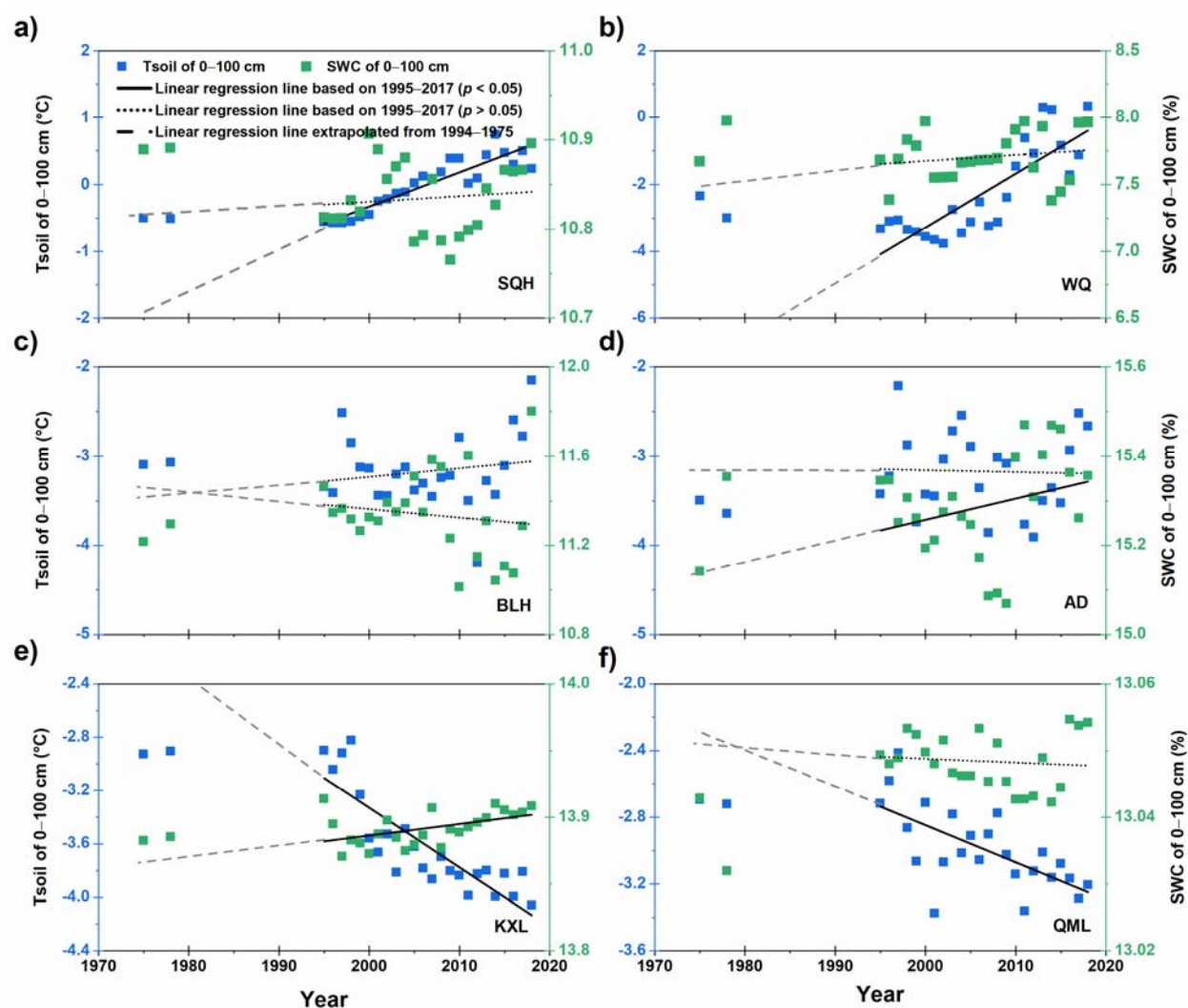


Supplementary Figure S1. Temporal changes in annual air temperature (Tair) at a height of 2 m (MAAT) on the Tibetan Plateau from 1975 to 2017. **a** is shown MAAT temporal changes for group A locations and group B locations from 1975 to 2017. For group A and B separation details, please see Supplementary Figure S1. Error bars denote SE, $n = 8$ for 1975–1994 and $n = 10$ for 1995–2017 for group A, and $n=4$ for group B. Subfigures **b**, **c**, **d**, **e**, and **f** show the annual air temperature data of group A locations of TSH, WQ, ZAD, HSX, QSH, SQH, GZ, MA, XD, and ZD, respectively. Subfigures **g** and **h** show the annual air temperature data of group B locations of AD, KXL, and BLH, QML, respectively. Solid lines for 1975–2017 indicate significant changes ($p < 0.01$, **; $p < 0.05$, *). The insets show the mean annual air temperature (MAAT) for group A, group B and specific locations, which with different color columns from 1975 (locations WQ and QSH from 1995) to 2017.



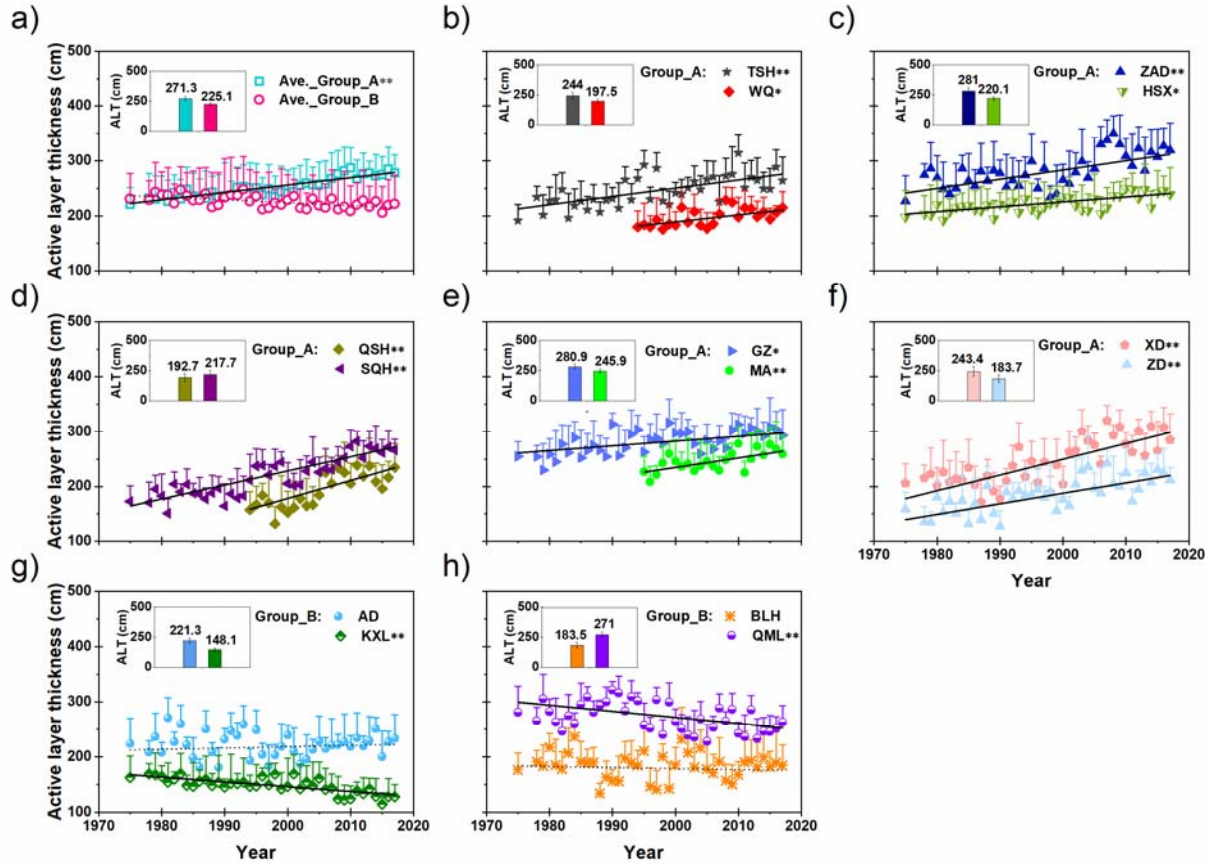
Supplementary Figure S2. Precipitation of 4 representative types of 14 locations (locations SQH and HSX belonged to the Group A; Locations AD and QML belonged to the Group B) in the permafrost region of the Tibetan Plateau from 1975 to 2017. Subfigure a, b, c, d, and e is used to designate spring, winter, summer, autumn, and annual total precipitation, respectively.

420

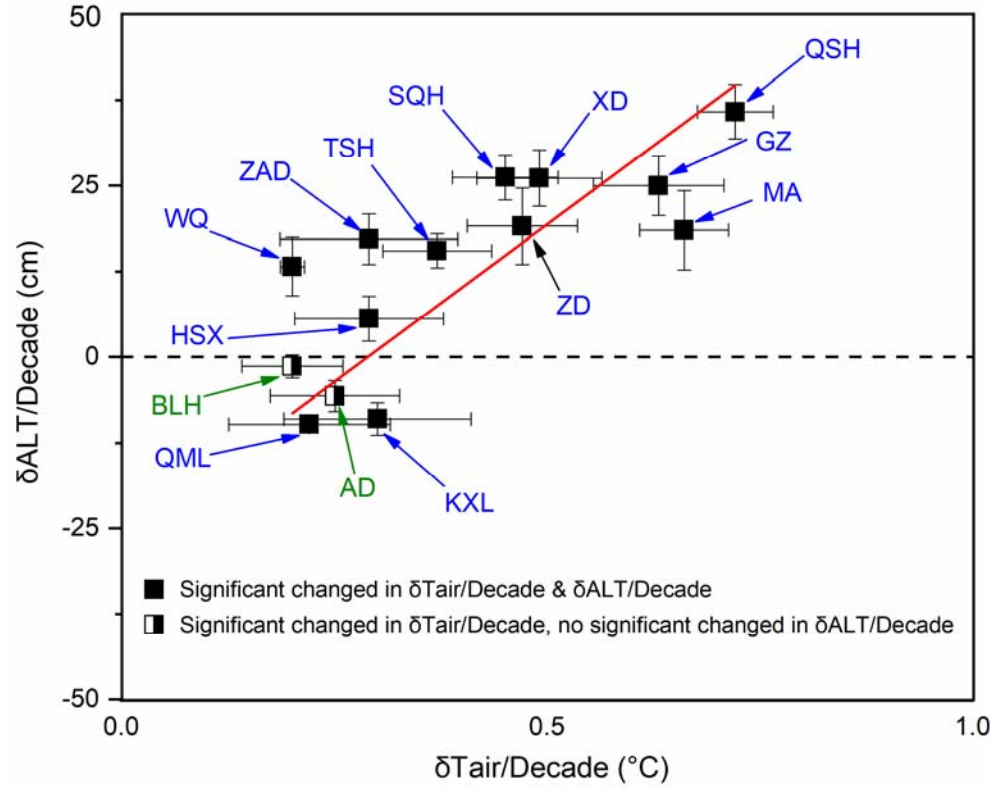


421

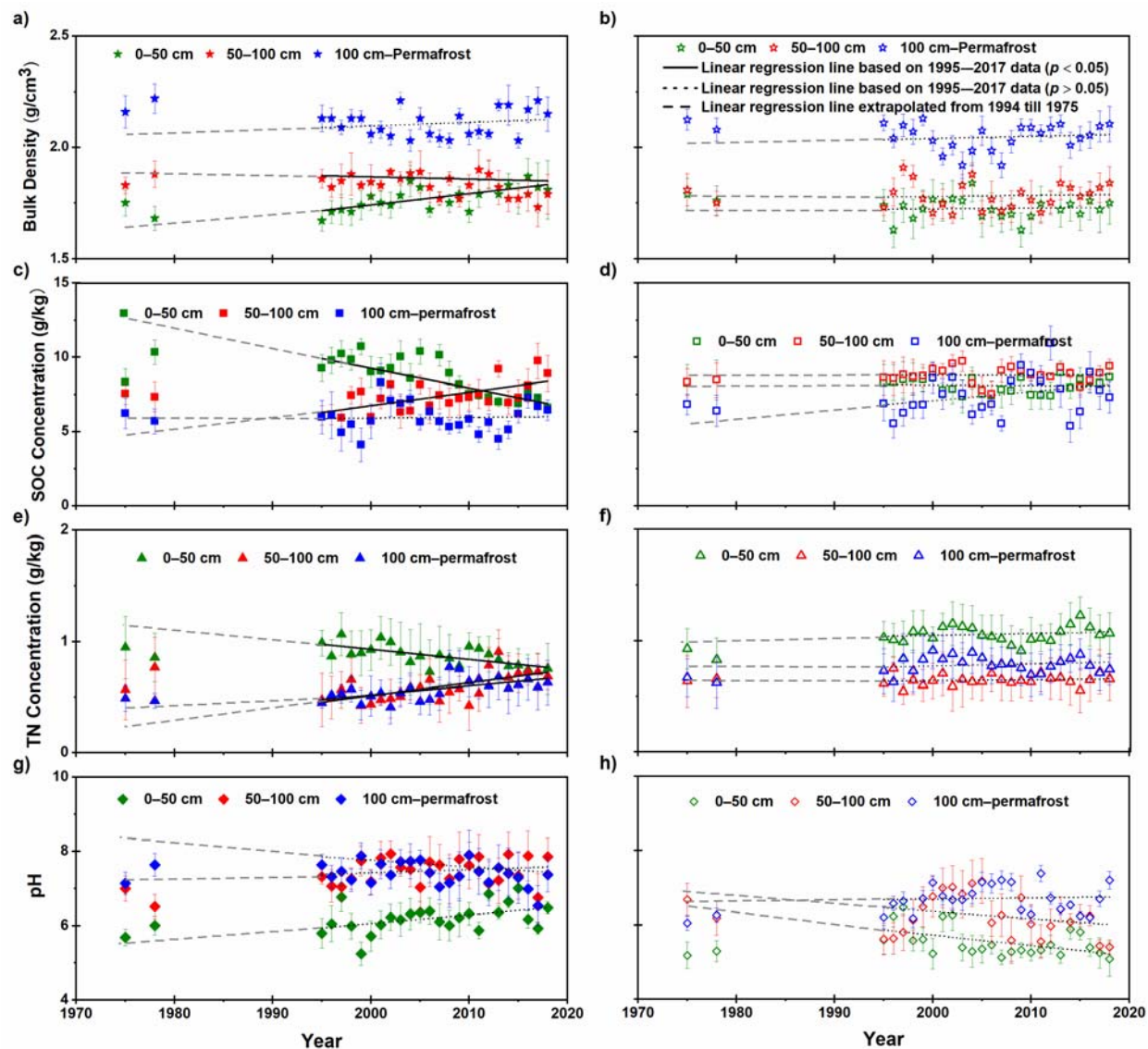
422 **Supplementary Figure S3.** Mean annual soil temperature (Tsoil) and soil water content (SWC)
 423 of 0–100 cm changes from 1975–2017 for six type locations on Tibetan plateau (Figs. S9 a-f).
 424 Location SQH, WQ belonged to group A; location BLH, AD, KXL, QML belonged to group B.



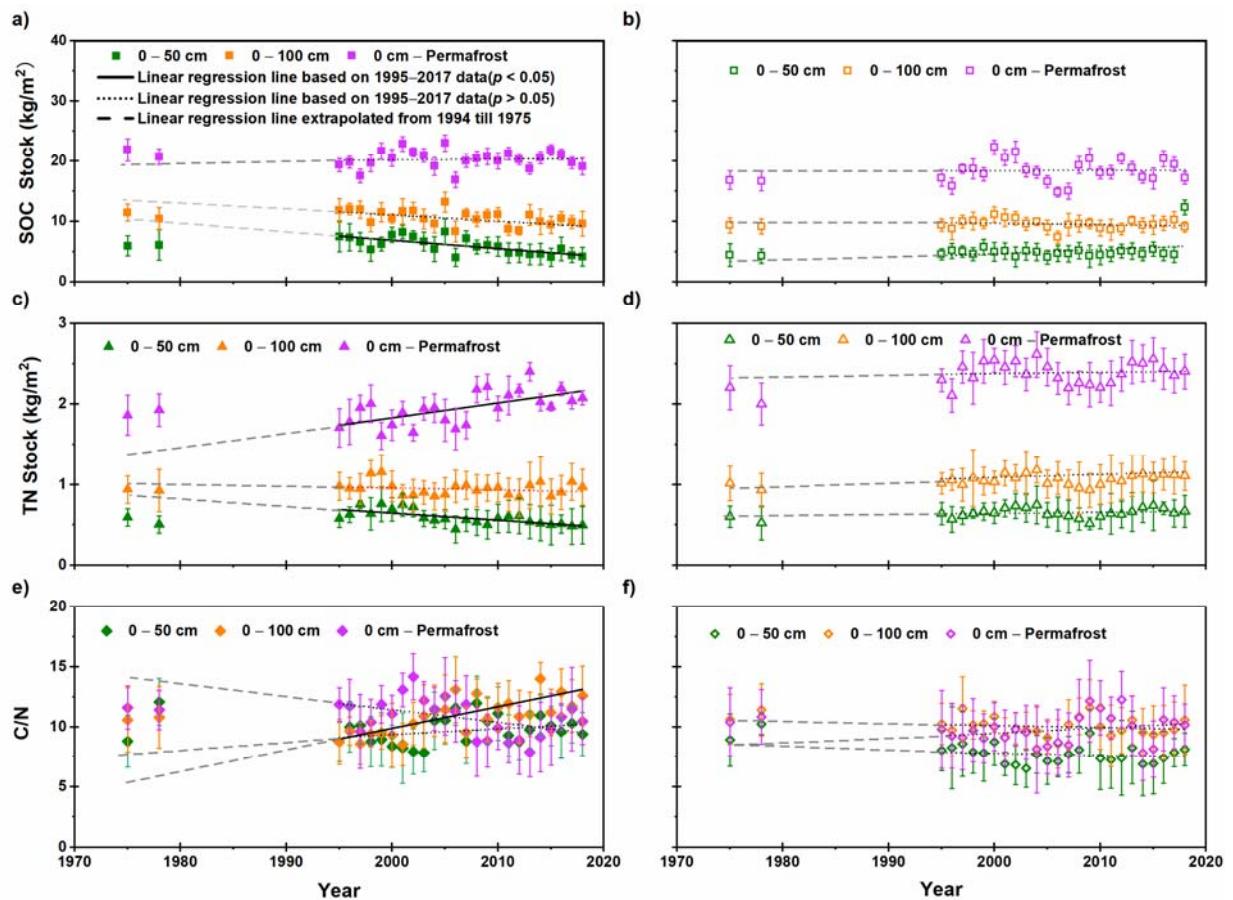
Supplementary Figure S4. Temporal changes in active layer thickness (ALT) on the Tibetan Plateau from 1975 to 2017. **a** is shown mean ALT temporal changes for group A locations and group B locations from 1975 to 2017. Error bars denote SE, $n = 7$ for 1975–1994 and $n = 10$ for 1995–2017 for group A, and $n=4$ for group B. Subfigures **b**, **c**, **d**, **e**, and **f** show the ALT data of group A locations of TSH, WQ, ZAD, HSX, QSH, SQH, GZ, MA, XD, and ZD, respectively. Subfigures **g**) and **h**) show the ALT data of group B locations of AD, KXL and BLH, QML, respectively. Solid lines for 1975–2017 indicate significant changes ($p < 0.01$, ** and $p < 0.05$, *), while dotted lines indicate non-significant changes ($p > 0.05$). The insets show the mean annual ALT for group A, group B and specific locations, which with different color columns from 1975 (locations WQ, QSH, and MA from 1995) to 2017.



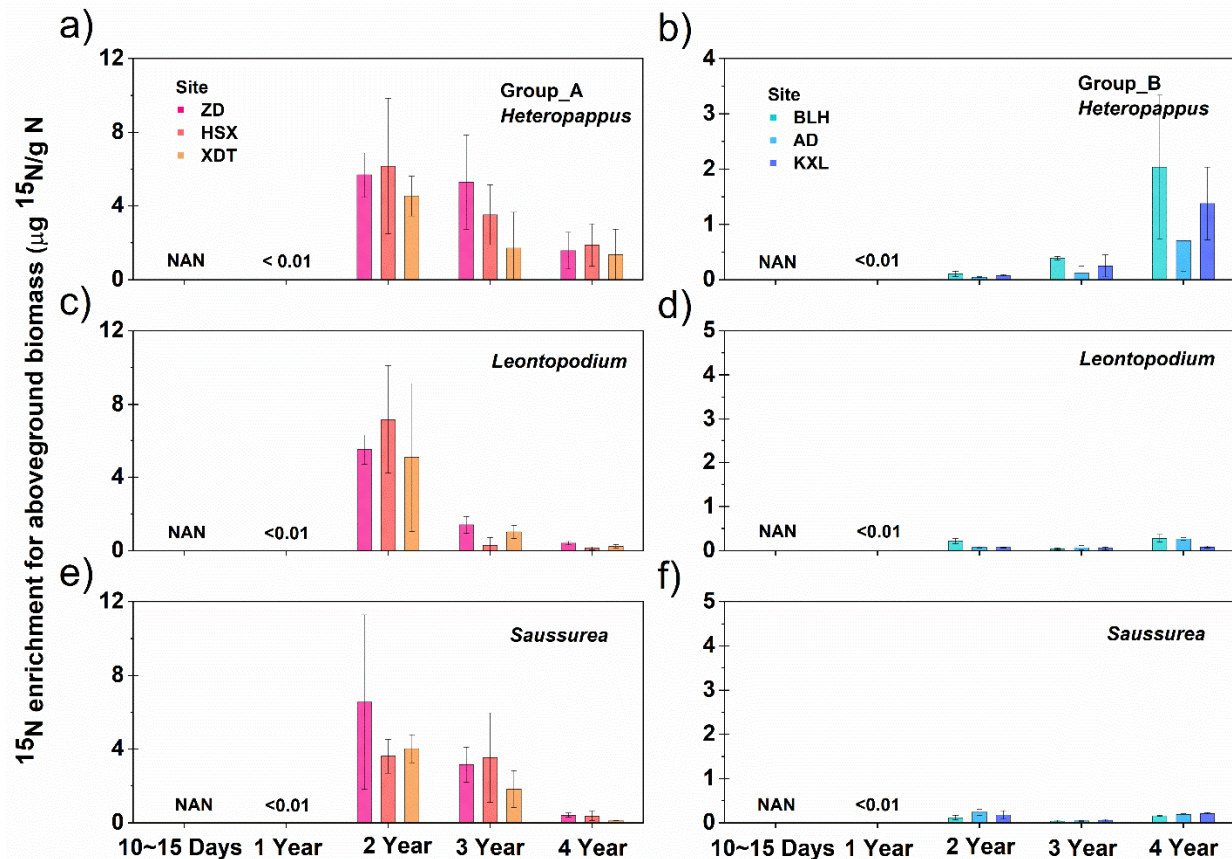
Supplementary Figure S5. Changes in air temperature and active layer thickness (ALT) per decade on the Tibetan Plateau from 1975 to 2017. Based on location-specific *t*-tests, the 14 stations are divided in group A locations (with a significant positive change in ALT with respect to air temperature; locations TSH, QSH, SQH, GZ, WQ, ZAD, XD, ZD, and HSX) and group B locations (with a significant negative in ALT with respect to air temperature, locations KXL and QML or no significant in ALT with respect to air temperature, locations AD and BLH). One standard deviation is shown as horizontal and vertical bars for $\delta T_{\text{air}}/\text{decade}$ ($n = 4$) and $\delta \text{ALT}/\text{decade}$ ($n = 4$), respectively. The red line is the linear fit line, which illustrates the sensitivity of changes in active layer thickness and changes in air temperatures ($y = 85.88x - 25.02$, $r = 0.80$; $p < 0.01$).



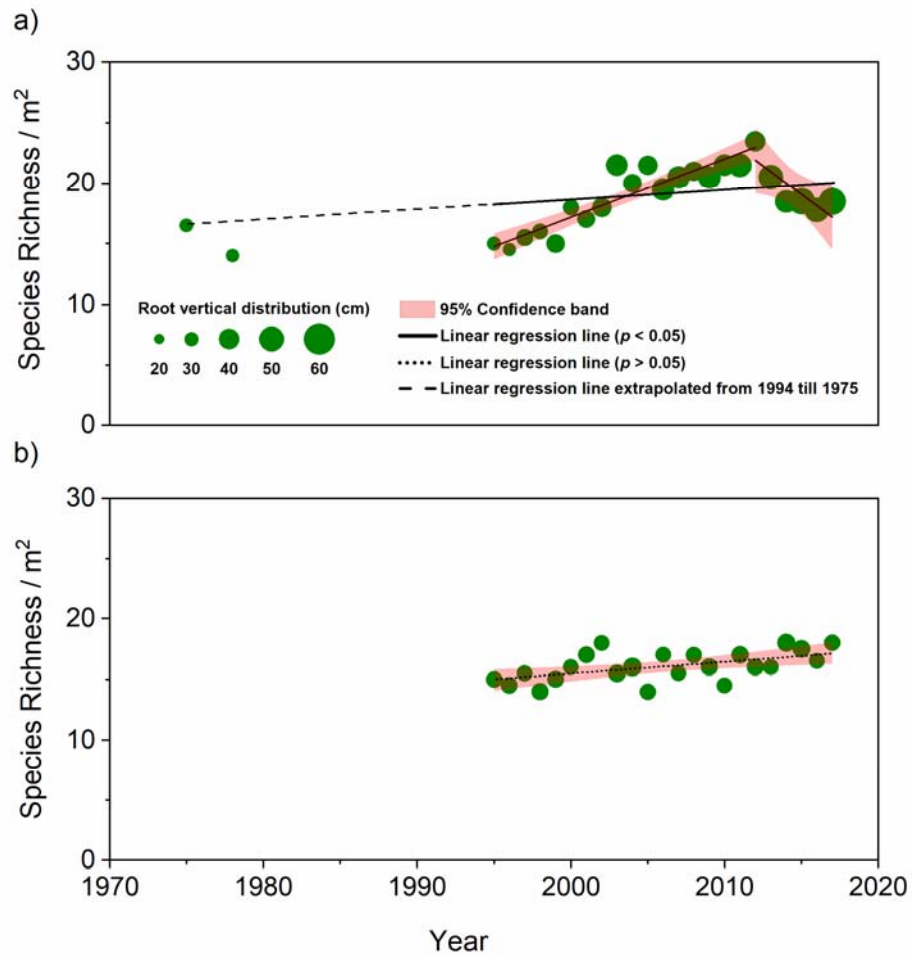
Supplementary Figure S6. Temporal changes in soil properties at three depth intervals averaged at locations across the Tibetan Plateau, which include soil bulk density, soil organic carbon (SOC) concentration, total nitrogen (TN) concentration, and soil pH. Average data are shown for group A locations (subfigures **a**, **c**, **e**, and **g**) and group B locations (subfigures **b**, **d**, **f** and **h**). Vertical bars represent one standard deviation, $n = 6$ for 1975, 1978, and $n = 10$ for 1995–2017 in group A; $n = 3$ for 1975, 1978 and $n = 4$ for 1995–2017 in group B.



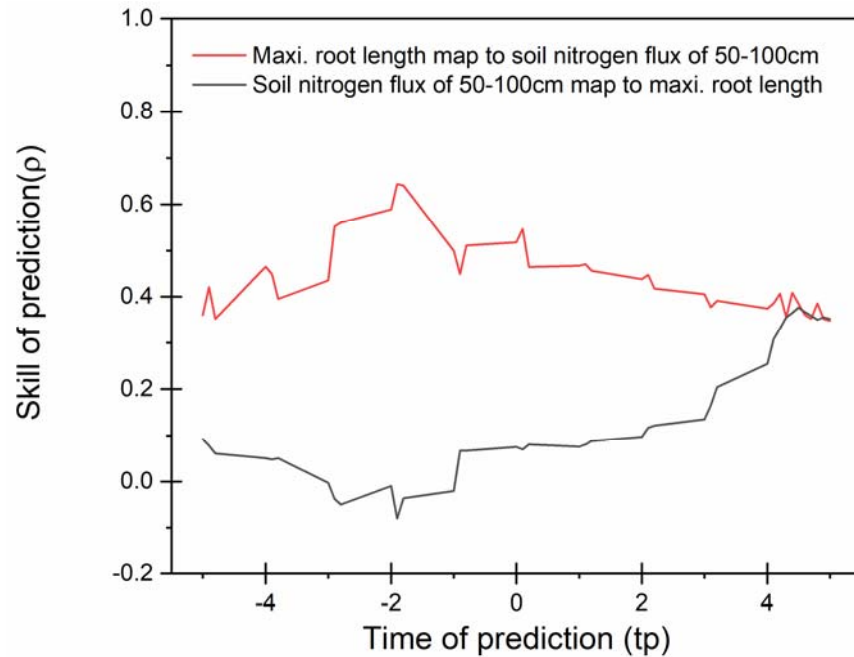
Supplementary Figure S7. Temporal changes in the soil organic carbon stock (SOC stock), total nitrogen stock (TN stock), and carbon-to-nitrogen ratio (C/N) at three depth intervals at locations across the Tibetan Plateau. Subfigures **a**, **c**, and **e** for group A locations; **b**, **d**, and **f** for group B locations. Worth notes, three depth intervals were 0–50 cm, 0–100 cm, and 0 cm–permafrost. Vertical bars represent one standard deviation, $n = 6$ for 1975, 1978, and $n = 10$ for 1995–2017 in group A; $n = 3$ for 1975, 1978 and $n = 4$ for 1995–2017 in group B.



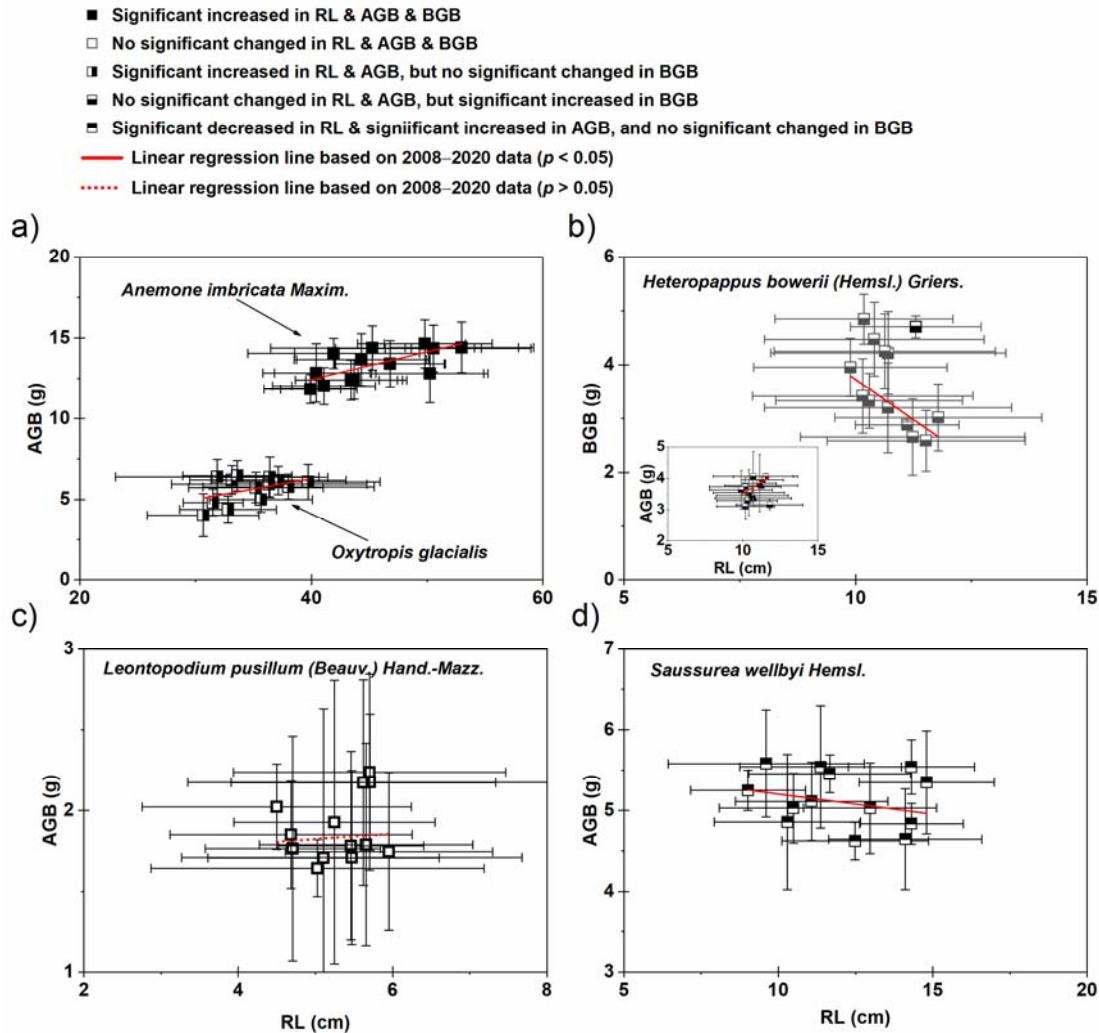
Supplementary Figure S8 Enrichment $^{15}\text{N_NH}_4\text{Cl}$ in aboveground mass (including leaves and stem) of shallow-rooted species (average root length less than 20 cm) by treatment after 10~15days, 1 year, 2 years, 3 years, and 4 years for group A (a, c, e) and group B (b, d, f). a and b were *Heteropappus bowerii* (Hemsl.) Griens. (*Heteropappus*), c and d were *Leontopodium nanum* (Hook. f. et Thoms.) Hand.-Ma (*Leontopodium*), and e and f were *Saussurea arenaria* Maxim (*Saussurea*). vertical bars representing one standard deviation, n = 15, excluded the 1 year, n= 12.



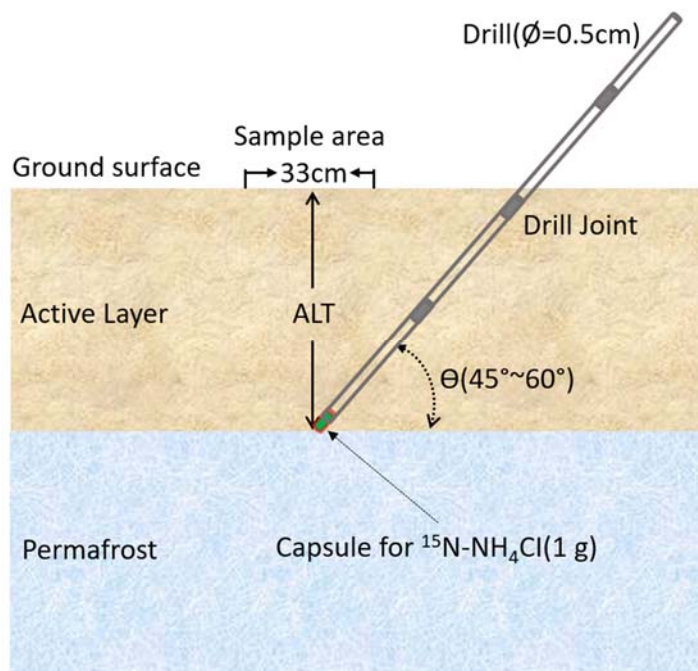
Supplementary Figure S9. Changes in specie richness and root vertical distribution from 1975 to 2017. The size of green dot indicates the difference of root vertical distribution. Subfigure **a** show the group A locations and **b** for the group B locations.



Supplementary Figure S10. Causality test running for the variation of root length map to soil nitrogen flux of 50-100 cm (red line) and soil nitrogen flux of 50-100 cm map to root length (grey line), respectively. Skill of prediction (ρ) for red line reaches the peak when time prediction (tp) around -2, demonstrate that the soil nitrogen flux of 50cm caused the root length change (increase), and the time lag is about 2 years. This result consisted with field observations of isotope ^{15}N that label N rejected to the front of permafrost is could be uptake by the deep-rooted plant in two years. In contrast, ρ of the grey line peak *arrived* maximum when tp around 4, which means the hypothesis of variation of root length caused soi nitrogen flux of 50cm variation is impossible. The optimal value of the embedding dimension $E=2$, estimated by function of SSR pred boot (R).



Supplementary Figure S11. Changes in root length (RL), aboveground biomass (AGB), and belowground biomass (BGB) for species type at the location XD (belonging for group A) from 2008 to 2020. **a** is two types deep-rooted species *Anemone imbricata* Maxim. and *Oxytropis glacialis*; **b**, **c**, and **d** are three shallow-rooted species *Leontopodium pusillum* (Beauv.) Hand.-Mazz., *Heteropappus bowerii* (Hemsl.) Griens., and *Saussurea wellbyi* Hemsl. **Worth note**, **b** is the changes in RL and BGB, the **inset** is the changes in RL and AGB for *Heteropappus bowerii* (Hemsl.) Griens. Vertical and horizontal bars represent one standard deviation, $n = 4$.



Supplementary Figure S12. Schematic of field experiment for $^{15}\text{N-NH}_4\text{Cl}$ capsule rejection to the front of permafrost. Collection of plants were made only immediately above the rejection point. Tests have been made previous (11) to document that the method ensure that residue is not left at shallow depth which otherwise might be accessible for plant roots.

514 **Supplementary Table ST1.** Species abundance or species turnover (presence or absence) over time of group A.

No.	HP (cm)	RL (cm)	Year																										
			1975	1978	1995	1996	1997	1998	1999	2000	2001	2002	2003	2004	2005	2006	2007	2008	2009	2010	2011	2012	2013	2014	2015	2016	2017		
1	27.4	37.2	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓		✓	✓	✓	✓		✓	✓		✓	✓	✓		
2	27.4	38.2	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓			✓		✓								
3	34.8	40.2	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓				✓	✓	✓	✓						
4	26.5	33.5	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓		✓	✓			
5	18.3	23.5	✓	✓	✓		✓	✓		✓	✓		✓	✓	✓								✓	✓					
6	15.4	18.5	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓					
7	23.6	38.2	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓		✓	✓	✓										
8	6.5	14.0	✓	✓	✓	✓		✓	✓		✓	✓	✓	✓	✓						✓	✓							
9	2.8	29.9	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓					✓	✓									
10	1.8	23.7	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓			✓	✓	✓	✓									
11	6.2	15.8	✓	✓	✓	✓	✓	✓	✓	✓		✓		✓		✓				✓									
12	8.0	14.6	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓		
13	5.9	5.1	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓																	
14	4.7	16.6	✓	✓	✓	✓		✓			✓		✓		✓							✓		✓					
15	8.1	11.4	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓			✓		✓									
16	6.5	7.3	✓	✓			✓	✓	✓	✓	✓	✓								✓									
17	6.8	13.8	✓	✓	✓		✓		✓	✓		✓	✓			✓													
18	5.4	5.2	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓		✓	✓				✓										
19	11.1	12.1	✓	✓		✓	✓	✓	✓	✓	✓	✓	✓	✓	✓		✓	✓											
20	6.3	5.5	✓	✓	✓	✓				✓				✓		✓													
21	9.8	10.8	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓		✓			✓	✓											
22	7.8	10.6	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓		✓		✓	✓	✓				
23	5.7	5.1	✓	✓	✓	✓		✓		✓		✓		✓	✓														
24	7.3	6.8	✓	✓	✓	✓	✓	✓		✓	✓	✓		✓		✓	✓	✓											
25	6.4	6.2	✓	✓		✓	✓	✓	✓	✓	✓	✓		✓		✓													
26	4.7	5.1	✓	✓	✓		✓		✓	✓	✓		✓																

27	5.1	6.1	✓	✓	✓	✓	✓		✓	✓	✓	✓			✓			✓	✓						
28	13.2	17.2	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓		✓	✓	✓					
29	18.6	39.2	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓			✓				
30	21.1	42.4	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓
31	10.4	12.4	✓	✓	✓	✓	✓		✓	✓		✓	✓												
32	7.6	12.3	✓	✓	✓					✓			✓	✓	✓			✓		✓		✓	✓	✓	✓
33	12.9	10.0	✓				✓			✓				✓				✓	✓	✓		✓	✓	✓	✓
34	15.7	8.3	✓	✓	✓	✓		✓		✓	✓	✓	✓	✓		✓	✓								
35	13.6	17.3	✓	✓	✓		✓	✓	✓	✓	✓	✓	✓	✓	✓	✓		✓							
36	15.6	42.8	✓	✓		✓	✓	✓	✓		✓	✓	✓		✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓
37	6.9	33.3	✓	✓					✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓
38	25.7	43.6	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓		✓			✓				✓			✓
39	16.8	13.5	✓	✓	✓	✓	✓	✓	✓	✓	✓		✓	✓	✓	✓	✓	✓	✓	✓	✓				
40	11.3	39.3	✓	✓	✓	✓	✓	✓	✓	✓	✓			✓				✓		✓				✓	✓
41	5.4	7.3	✓	✓	✓	✓	✓	✓	✓	✓		✓			✓	✓		✓	✓	✓	✓	✓	✓	✓	
42	12.6	42.4	✓	✓	✓	✓	✓	✓	✓	✓		✓				✓	✓	✓	✓	✓		✓	✓		✓
43	8.4	31.1	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓		✓	✓	✓		✓
44	7.3	32.3	✓	✓	✓	✓		✓			✓	✓			✓		✓		✓						✓
45	3.3	35.1	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓			✓		✓	✓	✓	✓	✓	✓	✓	
46	13.4	42.5	✓	✓			✓	✓	✓	✓	✓	✓	✓	✓		✓	✓	✓			✓		✓	✓	✓
47	11.1	15.2	✓		✓		✓		✓	✓	✓	✓	✓	✓		✓	✓		✓			✓		✓	✓
48	17.7	38.3	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓		✓	✓	✓		✓	✓	
49	6.2	4.7	✓	✓		✓	✓	✓	✓	✓	✓	✓	✓		✓	✓			✓	✓	✓	✓	✓	✓	✓
50	13.4	10.1	✓	✓	✓	✓					✓	✓	✓	✓	✓	✓	✓			✓		✓	✓	✓	✓
51	24.2	42.3	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓
52	12.6	12.1	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓		✓	✓	✓		✓	✓				
53	13.5	27.5	✓		✓	✓		✓		✓		✓	✓	✓				✓							✓
54	7.1	17.3	✓	✓	✓	✓	✓	✓		✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓
55	3.7	46.0	✓	✓		✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	
56	16.8	42.3		✓	✓		✓		✓	✓	✓	✓	✓	✓					✓	✓	✓		✓	✓	✓
57	35.1	47.2		✓	✓	✓	✓		✓		✓	✓	✓	✓	✓	✓	✓	✓	✓	✓	✓				
58	32.1	44.2		✓	✓	✓	✓	✓	✓	✓	✓		✓	✓	✓	✓	✓	✓				✓			

[illegible]

515 **Note** that HP is acronymic of mean height of plant, RL is acronymic of mean root length, values reported as annual means, not show
 516 the standard error. Species presence is used to designate √, species absence is used to designate blank, and No. is used to designate the
 517 species name. 1, *Kobresia littledalei* C. B. Clarke 2, *Kobresia tibetica* Maxim 3, *Kobresia robusta* Maxim 4, *Carex nudicarpa* (Y. C.
 518 Yang) S. R. Zhang (2015) 5, *Triglochin maritimum* 6, *Polygonum viviparum* L. 7, *Astragalus purdomii* 8, *Ranunculus longicaulis* C.
 519 A. Mey. var. *nephelogenes* (Edgew.) L. Liou 9, *Androsace tangulashanensis* Y. C. Yang et R. F. Hua. 10, *Androsace integra* (Maxim.)
 520 Hand – Mazz. 11, *Lagotis brachystachya* Maxim. 12, *Leontopodium pusillum* (Beauv.) Hand.-Mazz. 13, *Primula minor* Balf. f. et
 521 Ward 14, *Saussurea stella* Maxim. 15, *Aster alpinus* L. 16, *Thalictrum alpinum* L. 17, *Oxygraphis glacialis* (Fisch.) Bunge 18,
 522 *Saxifraga unguiculata* Engl. 19, *Taraxacum brevirostre* Hand.-Mazz. 20, *Silene gonosperma* 21, *Saussurea wellbyi* Hemsl. 22,
 523 *Polygonum sibiricum* Laxm. 23, *Gentianopsis paludosa* (Hook. f.) Ma 24, *Gentiana crenulatotruncata* (Marq.) T. N. Ho 25, *Gentiana*
 524 *futtereri* Diels et Gilg 26, *Glaux maritima* L. 27, *Callianthemum pimpinelloides* 28, *Heteropappus bowerii* (Hemsl.) Griens. 29, *Carex*
 525 *parvula* O. Yano 30, *Kobresia deasyi* C. B. Clarke 31, *Taraxacum tibetanum* Hand.-Mazz. 32, *Hedinia tibetica* (Thomson) Ostenf. 33,
 526 *Delphinium candelabrum* Ostf. var. *monanthum* (Hand.-Mazz.) W. T. Wang 34, *Delphinium candelabrum* Ostf.
 527 var. *monanthum* (Hand.-Mazz.) W. T. Wang 35, *Allium cyaneum* Regel 36, *Oxytropis pauciflora* Bunge 37, *Kobresia humilis* (C. A.
 528 Mey. ex Trautv.) Sergiev 38, *Iris potaninii* Maxim. 39, *Saussurea melanotrica* Hand.-Mazz. 40, *Erysimum chamaephyton* Maxim 41,
 529 *Parnassia trinervis* 42, *Corydalis dasyptera* Maxim. 43, *Potentilla multicaulis* Bge. 44, *Potentilla saundersiana* 45, *Arenaria*
 530 *brevipetala* Y. W. Tsui et L. H. Zhou. 46, *Aconitum tanguticum* (Maxim.) Stapf 47, *Dimorpha emon landulosus* Kar. et Kir 48,
 531 *Potentilla bifurca* 49, *Ajania tibetica* 50, *Aster flaccidus* Bge. subsp. *glandulosus* (Keissl.) Onno 51, *Pedicularis cheilanthesifolia* 52,
 532 *Cremanthodium humile* 53, *Cortiella caespitosa* Shan et Sheh 54, *Ajania khartensis* 55, *Thylacospermum* Fenzl T. rupifragum
 533 Schrenk. 56, *Ceratoides compacta* (Losinsk.) Tsien et C. G. Ma 57, *Rheum moorcroftianum* Royle 58, *Rhodiola algida* (Ledeb.)
 534 Fisch. et Mey. 59, *Rhodiola quadrifida* 60, *Heraeleum millefolium* 61, *Polygonum sibiricum* Laxm. var. *thomsonii* Meisn. Ex 62,
 535 *Potentilla parvifolia* Fisch. ex. Lehm. 63, *Myricaria prostrata* Hook. f. et Thoms. ex Benth. 64, *Oxytropis glacialis* 65, *Poa*
 536 *litwinowiana* Ovcz. 66, *Elymus nutans* Griseb. 67, *Stipa purpurea* 68, *Dracocephalum heterophyllum* Benth 69, *Allium carolinianum*

537 DC. 70, *Carex orbicularinucis* 71, *Carex moorcroftii* Falc. Ex Boott 72, *Meconopsis horridula* Hook. f. et Thoms. 73, *Anemone*
538 *imbricata* Maxim. 74, *Microula tibetica* 75, *Sibbaldia adpressa* 76, *Pleurospermum hedinii* 77, *Saussurea tibetica* C. Winkl 78,
539 *Saussurea Arenaria* 79, *Oxytropis melanocalyx* Bunge 80, *Trisetum spicatum* 81, *Trisetum tibeticum* P. C. Kuo et Z. L. Wu 82, *Carex*
540 *capillifolia* (Decne.) S. R. Zhang 83, *Draba altaica* (C. A. Mey.) Bunge 84, *Hypecoum erectum* L. 85, *Stipa roborowskyi* Roshev 86,
541 *Oxytropis stracheyana* Benth. ex Baker 87, *Astragalus melanostachys*.

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545 **Supplementary Table ST2. Species abundance or species turnover (presence or absence) over time of group B.**

No.	HP (cm)	RL (cm)	Year																							
			1975	1978	1995	1996	1997	1998	1999	2000	2001	2002	2003	2004	2005	2006	2007	2008	2009	2010	2011	2012	2013	2014	2015	2016
1	30.8	37.8	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√
2	29.3	37.0	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√
3	37.7	37.2	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√
4	29.7	31.6	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√
5	22.3	23.0	√	√	√	√	√		√	√	√		√	√	√		√	√	√	√		√	√	√	√	√
7	26.6	33.5	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√
8	11.4	14.7	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√
9	8.3	28.9	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√
10	6.1	25.6	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√
12	13.2	16.1	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√
13	9.8	9.9	√	√	√	√	√	√	√	√	√			√	√	√	√	√	√	√	√	√	√	√	√	√
15	12.4	12.0	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√
16	10.4	8.2	√			√					√	√				√							√	√	√	√
17	10.8	15.0	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√
18	10.4	9.4	√																			√				
19	15.9	16.0	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√
20	12.3	8.1	√							√	√										√					
21	12.9	13.8	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√
23	8.4	8.7	√	√		√	√		√	√	√		√	√	√	√		√		√	√	√	√		√	√
24	9.9	8.7	√	√	√	√																				
25	10.7	6.8	√	√	√	√	√	√	√																√	√
26	12.1	6.4	√				√		√		√	√	√		√	√	√	√		√	√	√		√	√	
27	13.0	8.3	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√
29	23.8	37.5	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√
30	26.8	38.2	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√	√

557 **Supplementary Table ST3.** Mean aboveground biomass for September (Biomass of Aboveground), mean plant coverage, and mean
558 ratio of maximum root depth to the thickness of the active layer (Max root depth: ALT) for groups A (G–A) and B (G–B) over time.

Year	Biomass of aboveground (g/m ²)		Plant coverage (%)		Ratio of max. root depth and ALT		n	
	G–A	G–B	G–A	G–B	G–A	G–B	G–A	G–B
1975	NA	NA	> 60	> 65	0.11 ± 0.09	0.14 ± 0.2	52	27
1978	NA	NA	> 60	> 65	0.10 ± 0.07	0.14 ± 0.19	48	20
1995	237.7 ± 8.5	250.4 ± 8.5	68.2 ± 5.8	75.4 ± 5.5	0.09 ± 0.05	0.15 ± 0.18	59	22
1996	241.9 ± 9.0	249.8 ± 9.7	69.8 ± 5.5	79.3 ± 7.4	0.08 ± 0.03	0.16 ± 0.18	55	23
1997	231.3 ± 9.4	251.9 ± 10.2	63.3 ± 5.6	61.8 ± 6.9	0.09 ± 0.01	0.18 ± 0.17	50	25
1998	243.6 ± 9.7	247.6 ± 10.4	63.8 ± 6.2	61.3 ± 4.0	0.08 ± 0.01	0.15 ± 0.16	45	29
1999	243.6 ± 9.6	245.7 ± 10.0	73.3 ± 5.5	65.3 ± 6.0	0.09 ± 0.02	0.15 ± 0.16	79	31
2000	237.7 ± 8.7	246.2 ± 8.5	70.8 ± 5.2	79.7 ± 7.5	0.07 ± 0.02	0.15 ± 0.15	74	31
2001	245.7 ± 8.2	248.1 ± 7.9	69.1 ± 6.3	69.5 ± 8.2	0.09 ± 0.03	0.15 ± 0.15	69	29
2002	249.9 ± 7.	248.9 ± 7.4	65.2 ± 6.4	67.2 ± 5.6	0.09 ± 0.04	0.19 ± 0.14	64	30
2003	232.8 ± 7.8	252.7 ± 8.3	67.0 ± 5.5	57.1 ± 4.9	0.10 ± 0.05	0.24 ± 0.13	60	33
2004	241.2 ± 8.7	253.4 ± 9.6	67.8 ± 6.7	59.9 ± 6.9	0.08 ± 0.06	0.18 ± 0.12	78	34
2005	238.1 ± 9.1	252.7 ± 10	56.6 ± 7.4	63.1 ± 6.2	0.10 ± 0.07	0.21 ± 0.12	74	37
2006	236.0 ± 9.	253.7 ± 10.1	69.3 ± 6.3	54.2 ± 5.3	0.10 ± 0.08	0.18 ± 0.11	70	39
2007	235.7 ± 9.2	254.6 ± 9.7	73.8 ± 5.6	56.8 ± 6.8	0.11 ± 0.04	0.19 ± 0.10	66	20
2008	239.5 ± 9.4	254.4 ± 10.1	65.7 ± 5.6	53.7 ± 7.8	0.09 ± 0.02	0.21 ± 0.17	62	24
2009	228.3 ± 9.2	252.4 ± 9.6	58.7 ± 6.2	61.7 ± 8.6	0.11 ± 0.00	0.21 ± 0.17	51	28
2010	233.2 ± 8.8	249.5 ± 8.9	59.6 ± 4.1	64.4 ± 8.0	0.09 ± 0.03	0.16 ± 0.16	53	30
2011	221.0 ± 8.6	244.4 ± 9.1	54.4 ± 5.4	69.1 ± 6.3	0.12 ± 0.05	0.16 ± 0.15	52	33
2012	219.8 ± 8.8	233.4 ± 9.3	65.6 ± 7.0	63.1 ± 7.7	0.10 ± 0.04	0.17 ± 0.15	53	33
2013	227.8 ± 8.8	247.5 ± 9.3	63.2 ± 5.9	62.4 ± 9.3	0.11 ± 0.03	0.18 ± 0.14	63	34
2014	223.8 ± 9.2	229.4 ± 10.0	68.0 ± 5.8	74.3 ± 10	0.11 ± 0.03	0.16 ± 0.14	69	38
2015	229.0 ± 9.4	261.5 ± 9.9	64.7 ± 6.8	66.3 ± 6.6	0.12 ± 0.02	0.18 ± 0.13	57	36
2016	237.5 ± 8.6	256.5 ± 8.5	69.1 ± 5.2	67.6 ± 7.6	0.13 ± 0.02	0.20 ± 0.13	52	33
2017	233.8 ± 7.1	251.5 ± 8.6	70.1 ± 6.4	65.4 ± 7.1	0.12 ± 0.02	0.17 ± 0.12	50	35

559 Note: values reported as annual mean \pm SE. Data of the thickness of the active layer (ALT) during 1975-1995 based on the re-analysis
560 from air temperature (ref. 17), during 1996-2017 based on borehole soil temperature (method by ref. 5).

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