

Cold Diffusion Model for Seismic Denoising

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Key Points:

- First application of a Deep Learning (DL) model using cold diffusion to denoise seismic data;
- Cold Diffusion Model for Seismic Denoising (CDiffSD) demonstrates a superior performance in scenarios with low SNR, which is particularly challenging and crucial for effective seismic data analysis;
- CDiffSD model outperforms existing methods thereby establishing a new standard in seismic data denoising.

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Abstract

Seismic waves contain information about the earthquake source, the geologic structure they traverse, and many forms of noise. Separating the noise from the earthquake signal is a critical first step in seismic waveform analysis. This is, however, a difficult task because optimal parameters for filtering noise typically vary with time and, if chosen inappropriately, they may strongly alter the original seismic waveform. Diffusion models based on Deep Learning (DL) have demonstrated remarkable capabilities in the restoration of images and audio signals. However, those models assume a Gaussian distribution of noise, which is not the case for typical seismic noise. Diffusion models trained on Gaussian noise do not perform well in seismic applications; therefore, we introduce a "cold" variant of diffusion models in which both clean and noisy seismic traces are restored. Here, we describe the first Cold Diffusion Model for Seismic Denoising (CDiffSD), including key design aspects, model architecture, and noise handling. We demonstrate that CDiffSD provides a new standard in performance, outperforming existing methods. Our model provides a significant advance for seismic data denoising and establishes a new state-of-the-art in the field.

Plain Language Summary

Seismic waves contain information about earthquakes and the earth's structure but any seismic waveform is, to a variable extent, contaminated by noise. Separating noise from earthquakes is important in order to enhance signals quality and, as a consequence, improve subsequent analyses. However, this task can be challenging because not only noise characteristics change in time, frequency and amplitude, but also because an incorrect denoising procedure might significantly alter important features of the seismic waves. Recently, deep learning techniques have proven to be valuable tools in enhancing images and audio signals. But these techniques usually expect the noise to follow a certain pattern that doesn't match the more complex noise found in seismic data. To solve this, we've developed a new approach called the Cold Diffusion Model for Seismic Denoising (CDiffSD). This model, specifically designed to handle the types of noise found in seismic data, shows better performances than previous methods in removing noise and restoring seismic signals, ultimately setting a new high standard in the field.

1 Introduction

Seismograms contain signals generated by earthquakes and by other unidentified sources categorized in general as 'noise' (e.g., oceanic waves, wind, vehicular traffic, sonic booms, quarry activities, and instrument malfunctions.) It is standard practice in seismology to denoise waveforms to improve the performance of the subsequent analyses, such as P- and S-wave onset picking, earthquake source moment tensor inversion, and techniques of exploration seismology. Most commonly and in routine analysis, denoising is performed through bandpass filtering. However recent works have proposed several more sophisticated schemes to "clean" seismic traces. These include methods based on the independent component analysis (ICA) (Comon, 1994; Cabras et al., 2010; Moni et al., 2012), beamforming methods (Gibbons et al., 2008; Boué et al., 2013; Brooks et al., 2009), and MULTiple SIGNAL Classification (MUSIC) (Schmidt, 1986; Bear et al., 1999). All of these methods, however, can fall short when the noise shares frequencies with the earthquake generated signal.

Denoising models have evolved to incorporate time-frequency methods, with techniques like the Wavelet transform (Gaci, 2014; Siyuan & Xiangpeng, 2005; W. Liu et al., 2016; Mousavi & Langston, 2016b; Mousavi et al., 2016; Mousavi & Langston, 2017), the Short-Time Fourier Transform (STFT) (Mousavi & Langston, 2016a), the S-transform (Tselentis et al., 2012), and other transformation-decomposition methods (Hennenfent & Herrmann, 2006; Bekara & der Baan, 2009; Neelamani et al., 2008; Han & van der Baan,

2015; Y. Liu et al., 2013; Chen & Ma, 2014; Shan et al., 2009; Tang & Ma, 2011). These techniques have proven useful but the emergence of deep learning (DL) has provided new strategies with improved performance. A notable development in this arena is the Deep Denoiser (DD) model (Zhu et al., 2019). The DD approach is based on a variant of the Variational Autoencoder (VAE) (Kingma & Welling, 2019), which generates dual masks for seismic and noise signals, enhancing waveform extraction. Another notable approach is that of van den Ende et al. (2021) who employed a DL to denoise Fiber-optic Distributed Acoustic Sensing (DAS) data. They demonstrate the potency of DL to enhance the quality of DAS and seismic data. Similarly, the Novoselov et al. (2022) project, utilizing a Dual-Path Recurrent Neural Network (DPRNN), led to another substantial stride in the application of deep learning for seismic signal denoising. These studies not only validate the efficacy of deep learning methods in seismic noise reduction but also pave the way for further innovations in this field.

Here we built on this topic, drawing parallels with techniques used in speech enhancement, a field closely related to seismic denoising. Speech enhancement has recently seen the use of models such as GANs (Pascual et al., 2017; Donahue et al., 2018; Cao et al., 2022; Kim et al., 2021) and VAEs (Fang et al., 2021; Leglaive et al., 2020, 2018; Bie et al., 2022). However, the recent trend points to the growing success of Diffusion Models (Sohl-Dickstein et al., 2015; Ho et al., 2020), which are now outperforming their predecessors. Using techniques like cold diffusion or Gaussian diffusion for denoising presents several advantages over approaches that use binary masks, especially in terms of flexibility, reconstruction quality, and the ability to handle complex noise; while binary generally retain advantages in terms of simplicity, speed, interpretability, and computational efficiency. Here, we investigate the application of diffusion models for seismic denoising. These models typically transform the input into an isotropic Gaussian distribution through the consistent addition of Gaussian noise. In the reverse process, diffusion probabilistic models aim to remove the anticipated noise from the corrupted input, thus recovering the original signal. However, given the non-Gaussian nature of seismic noise, traditional diffusion models are not directly applicable.

This challenge led us to explore the emerging Cold Diffusion model (Bansal et al., 2022; Yen et al., 2023), which adapts the diffusion process by replacing Gaussian noise with other types of degradation processes. The Cold Diffusion model demonstrates how diffusion models can effectively restore signals impaired by various types of degradation. Its inherent properties make it particularly suitable for tasks such as speech source separation in practical settings with non-Gaussian noise. Building on this, our research aims to adapt the cold diffusion paradigm for seismic trace denoising. This adaptation involves specific modifications, primarily in the sampling algorithm, to suit the unique challenges of seismic data. The result is a Cold Diffusion Model for Seismic Denoising (CDiffSD).

2 Methods

2.1 Formalization of the problem

We begin with problem formulation and application of a novel diffusion model specifically designed for seismic denoising. The primary challenge in seismic signal processing is to extract the earthquake signal, denoted as x_0 , from a noisy signal $y = x_0 + x_n$. This signal y consists of the desired seismic signal x_0 and an unwanted noise component x_n . The goal is to develop a model that can effectively learn to approximate the function $f(y) = x_0$, thereby isolating the earthquake signal from the noise. To address this challenge, we introduce a diffusion probabilistic model, which utilizes both forward and reverse processes for noise reduction:

1. **Diffusion process** (or Forward process) is defined as a T-step Markov chain that gradually adds Gaussian noise to the recorded earthquake x_0 :

$$D(x_{1:T}|x_0) := \prod_{t=1}^T D(x_t|x_{t-1}) = \prod_{t=1}^T N(x_t; \sqrt{1 - \beta_t}x_{t-1}, \beta_t I) \quad (1)$$

2. **Reverse process** (or Backward process) aims to restore x_0 from the latent variable x_T based on the following Markov chain:

$$R_\theta(x_{0:T}) := R(x_T) \prod_{t=1}^T R_\theta(x_{t-1}|x_t) := R(x_T) \prod_{t=1}^T N(x_{t-1}; \mu_\theta(x_t, t), \Sigma_\theta(x_t, t)) \quad (2)$$

where $\beta = 1 - \alpha$ serves as a key parameter that controls the process of adding and removing noise in signal during the training and inference process. In particular, the Markov formulation asserts that a given distribution depends only on the previous timestep, hence we can rewrite (2) as:

$$R_\theta(x_{t-1}|x_t) := N(x_{t-1}; \mu_\theta(x_t, t), \Sigma_\theta(x_t, t)) \quad (3)$$

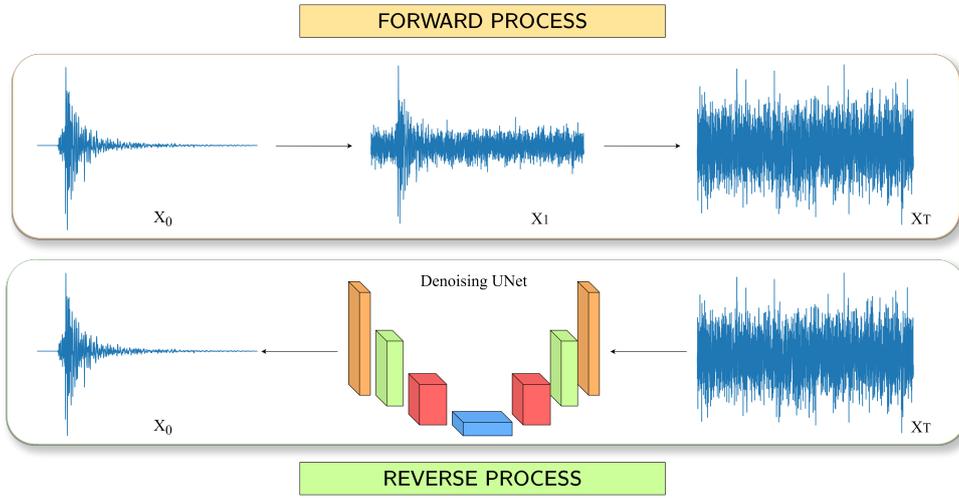


Figure 1. Sketch of how the Diffusion Process is adapted for seismic data. The standard Forward Process, which typically adds Gaussian noise, is modified to incorporate real noise, which defines so-called Cold Diffusion. The Reverse Process then employs neural networks to recover the recorded earthquake from the noise-enhanced data, illustrating the transition from noisy data back to the recorded data.

2.2 Diffusion Models

We explore diffusion models in some detail in order to elucidate key aspects of the training phase for our DL model and its operational principles. Understanding these elements is essential for appreciating how diffusion models achieve effective noise reduction and signal recovery in complex data sets. Starting from Equation (3), we can define:

$$\mu_\theta(x_t, t) = \frac{1}{\sqrt{\hat{\alpha}_t}} \left(x_t - \frac{\beta t}{\sqrt{1 - \hat{\alpha}_t}} \epsilon_\theta(x_t, t) \right) \quad (4)$$

The function $\mu_\theta(x_t, t)$ predicts the mean of x_{t-1} by removing the estimated Gaussian noise $\epsilon_\theta(x_t, t)$ in x_t , and the variance of x_t is fixed to a constant $\hat{\beta}_t = \frac{1 - \hat{\alpha}_{t-1}}{1 - \hat{\alpha}_t} \beta_t$.

The employed strategy is as follows: during training, a random time step t is sampled, and the signal is progressively degraded with Gaussian noise until reaching time

Algorithm 1 Diffusion Model Training

repeat $x_0 \sim D(x_0)$ $t \sim \text{Uniform}(\{1, \dots, T\})$ $\epsilon \sim N(0, I)$ Take gradient descent step on $\nabla_{\theta} \|\epsilon - \epsilon_{\theta}(\sqrt{\alpha_t}x_0 + \sqrt{1 - \alpha_t}\epsilon_t t)\|^2$ **until** converged

135 t, after which the signal is restored. Once the model is trained in this manner, the sam-
 136 pling process follows. The model removes noise step by step from time step t to T , as
 137 in Equations 1 and 2. This process is generally motivated by two factors. First, diffu-
 138 sion models can be harnessed to generate novel synthetic data starting from a strongly
 139 degraded step. The second motivation is that, particularly in denoising tasks, the step-
 140 by-step noise removal approach is expected to yield superior performance compared to
 a direct procedure. Our initial experiments involved adding Gaussian noise to earthquake

Algorithm 2 Diffusion Model Sampling

 $x_T \sim N(0, I)$ **for** $t = T, \dots, 1$ **do** $z \sim N(0, I)$ if $t > 1$, else $z = 0$ $x_{t-1} = \frac{1}{\sqrt{\alpha_t}}(x_t - \frac{1-\alpha_t}{\sqrt{1-\alpha_t}}\epsilon_{\theta}(x_t, t)) + \sigma_t z$ **end for****return** x_0

141 seismic traces, but we found that most seismic noise does not conform to Gaussian noise
 142 characteristics. Therefore, we adopted a more general and effective approach using a model
 143 based on Cold Diffusion (Bansal et al., 2022). This model generalizes diffusion models
 144 by replacing Gaussian noise with real noise. By using real seismic noise patterns, the model
 145 can more accurately and effectively perform denoising tasks, reflecting the actual com-
 146 plexities and variations found in seismic data, aligning better with the behavior of seis-
 147 mic traces to be denoised.
 148

149 This approach marks a significant leap forward in applying diffusion models to prac-
 150 tical tasks, integrating the use of real noise. Such integration not only confirms the find-
 151 ings of previous research (Bansal et al., 2022; Yen et al., 2023), which successfully ap-
 152 plied the cold diffusion model in computer vision and speech enhancement respectively,
 153 but also highlights the limitations of traditional methods dependent on Gaussian noise.
 154 This is because Gaussian noise may not adequately represent the real-world character-
 155 istics of seismic data.

2.3 Proposed Method: Cold Diffusion Seismic Denoising (CDiffSD)

157 The core of our model involves degrading a one-dimensional earthquake, in the form
 158 of a seismic record, x_0 (the target), with recorded seismic noise x_n , to produce an out-
 159 of-domain sample (noisy signal): $x_T = x_0 + x_n * NRF$. Here, x_0 represents an earth-
 160 quake recorded by a seismometer. While x_0 serves as a 'clean' sample in our context, it's
 161 important to note that it inherently contains some level of noise, given its real, non-synthetic
 162 origin.

163 The "Noise Reduce Factor" (NRF) is a key element in our specific analysis. It's
 164 responsible for calibrating the amplitude of the noise signal (x_n) in relation to the earth-
 165 quake signal's amplitude, often indicated by the amplitude of S-waves in the data. By

166 choosing a NRF value within the range 0.4 to 0.65, we ensure that the noise does not
 167 dominate the trace compared to the earthquake. We work with data from different stations
 168 that independently record noise and earthquake signals. It's worth mentioning that
 169 we mix earthquake x_0 and noise x_n recorded from different seismic stations, to improve
 170 generalizability and robustness. This setup provides our model with an input x_T and a
 171 ground truth x_0 , enabling effective backpropagation of loss and performance measurement.
 172

173 Therefore, concerning the specific degradation introduced in diffusion models, we
 174 can rephrase the degradation at time T as follows:

$$x_t = D_{x_T}(x_0, t) = \sqrt{\alpha_t}x_0 + \sqrt{1 - \alpha_t}x_T \quad (5)$$

175 where x_0 is the recorded earthquake $x_T = x_0 + x_n * NRF$ and $\alpha \in [0, 1]$ is the param-
 176 eter interpolation weight. α can also be regarded as the amount of information retained
 177 in the diffusion process, and it can alternatively be defined as $1 - \beta$, where beta rep-
 178 represents the amount of noise introduced in the degradation, such parameters are defined
 179 a priori by a scheduler.

180 Regarding the specific operation of cold diffusion models, our approach is delineated
 181 using the improved training algorithm proposed by (Yen et al., 2023):

Algorithm 3 Cold Diffusion Enhanced Training

for $n = 1, \dots, N_{iter}$ **do**
 Sample clean data x_0
 Sample $t \sim \text{Uniform}(\{1, \dots, T\})$
 $x_t \leftarrow D(x_0, t), \hat{x}_0 \leftarrow R_\theta(x_t, t)$
 Sample $t' \sim \text{Uniform}(\{1, \dots, t\})$
 $\hat{x}_{t'} \leftarrow D(\hat{x}_0, t'), \hat{\hat{x}}_0 \leftarrow R_\theta(\hat{x}_{t'}, t')$
 Take gradient descent step on $\nabla_\theta(\|\hat{x}_0 - x_0\|_1 + \|\hat{\hat{x}}_0 - x_0\|_1)$
end for

182 This approach enhances the robustness of the training phase when applied in the
 183 presence of non-Gaussian noise. During the training phase, the model randomly selects
 184 a time step t within the range $[0, T]$. At this point, the signal is degraded by introduc-
 185 ing recorded noise, after which a restoration operation is applied. This step is crucial as
 186 it simulates the process of denoising, where the model learns to reverse the effects of noise
 187 on the signal. The algorithm further deepens its learning by reiterating this process with
 188 a new time step t' where $t' < t$. At this stage, the signal undergoing degradation is not
 189 the recorded earthquake, but rather the one that has already been restored in the pre-
 190 vious step. The signal is then degraded again up to the new time step t' and subsequently
 191 restored. This iterative process of degrading and restoring at various time steps allows
 192 the model to learn more robustly, adapting to the complexities introduced by real noise
 193 patterns. The improved training algorithm is tolerant to shifting errors during the sam-
 194 pling process. As we can observe, Algorithm 3, the training process incorporates $\hat{x}_{t'}$, which
 195 is the denoised signal. This results in $\hat{\hat{x}}_{t'}$, which now contains the misalignment error that
 196 may occur during the sampling process.

2.3.1 Input Assumptions

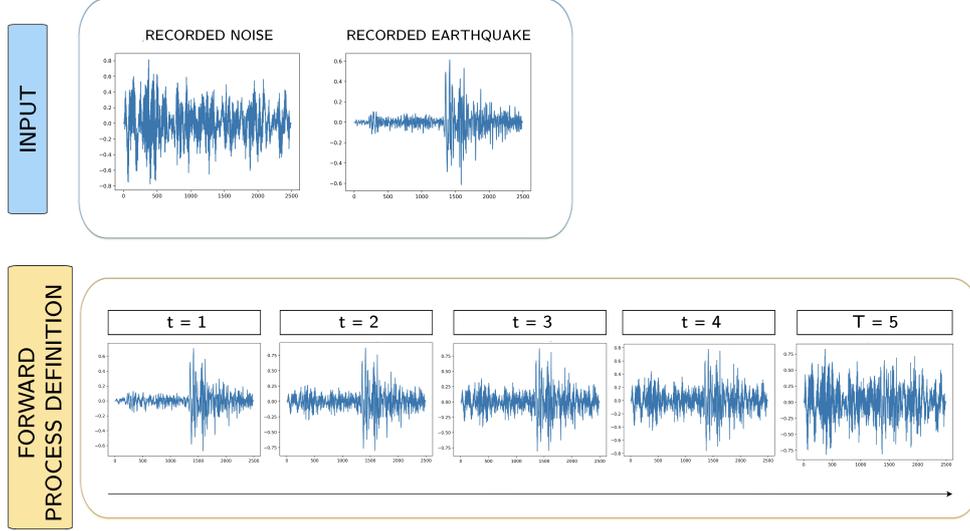


Figure 2. An illustration of the forward process with real noise for $T = 5$. The recorded noise, the recorded earthquake, and their various combinations are presented according to Equation 5. Notably, the combinations of noise and earthquake magnitudes are dictated by a scheduler, which pre-determines the levels of β and α . The level of noise at level $T=5$ is defined by the NRF.

198 In our seismic denoising approach, we separately normalize the noise and earth-
 199 quake data. We adopt a trace-specific method, normalizing each seismic trace (earth-
 200 quake and noise) across its East-West (E), North-South (N), and Vertical (Z) channels.
 201 This normalization process aligns the maximum and minimum values within these chan-
 202 nels, standardizing the data to a range of $[-1,1]$. Such an approach ensures that each com-
 203 ponent retains its relative amplitude, enabling precise and balanced analysis. This also
 204 enhances generalizability for each type of seismic trace that the end user wants to de-
 205 noise.

206 In the training phase for each seismic trace, we begin by merging a normalized earth-
 207 quake trace with a normalized noise trace. The noise component is scaled using the *NRF*,
 208 adjusting its intensity in the noisy signal before the forward process is applied.

209 The creation of the 'noisy signal' x_T , a combination of the earthquake and scaled
 210 noise signals, leads to the 'forward process'. Here, a stochastic variable t , ranging from
 211 0 to a predetermined maximum T , is chosen for further noise modulation. At $t = 0$,
 212 we have a recorded earthquake signal with no additional noise, whereas at $t = T$, the
 213 noise is at its full scale, set at 1.

2.3.2 Model Configuration

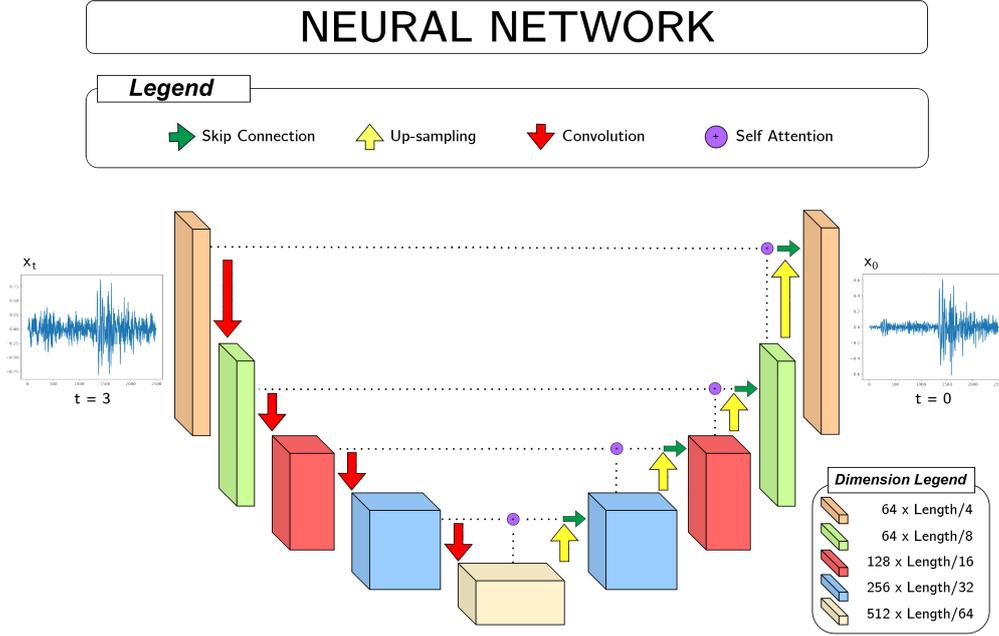


Figure 3. CDiffSD employ combinations of convolutional layers, ResNet blocks, and attention mechanisms to process one-dimensional data efficiently. The attention mechanisms are particularly important as they allow the models to selectively weigh the importance of different parts of the input data, aiding in extracting meaningful features for downstream tasks.

As a building block of the diffusion model, we adopt a neural network model inspired by the 1D U-Net (Ronneberger et al., 2015) design, for the processing of one-dimensional data streams such as time series or audio signals. The network begins with 1D convolutional layers, each equipped with 64 filters of a kernel dimension of 7, instrumental for the initial extraction of salient features. This is followed by the integration of temporal processing units that leverage sinusoidal positional encoding to effectively capture the temporal intricacies inherent within the data. These units then employ two linear layers dedicated to feature refinement and are paired with the Gaussian Error Linear Unit (GELU) (Hendrycks & Gimpel, 2016) activation function to instill the requisite non-linearity. As the architecture progresses, it introduces dimensionality manipulation layers consisting of ResNet modules (He et al., 2016) pivotal for feature conservation during down-sampling and 1D convolutional layers for further data refinement. Post-downsampling, a series of upsampling layers are implemented, designed to elevate data dimensionality by merging ResNet blocks with dedicated upsampling operations. A noteworthy feature of our design is the mid-level blocks, each outfitted with dual residual units. They exploit attention mechanisms crucial for highlighting pertinent data characteristics. The network culminates with terminal residual blocks that are succeeded by 1D convolutional layers, making definitive outputs typically manifest as singular channels. The U-Net block is applied for each iteration of the diffusion model from each t_i to 0 and then again from t_{i-1} to 0 and so on until the end of the process.

We trained models with 3 configurations: $T = 20, T = 100, T = 300$. These diverse scheduler assumptions allowed us to evaluate how performance metrics vary with increasing T , highlighting the trade-off between model performance and computation time,

238 which is a crucial consideration in seismic monitoring room operations where balancing
239 processing speed and precision is essential.

240 Particularly in the inference phase, understanding the impact of T on both model
241 performance and computational efficiency is vital. For applications requiring rapid trace
242 processing, like real-time seismic monitoring, a preference for speed may be necessary,
243 though it could impact precision. Conversely, in tasks where accuracy is the priority, such
244 as dataset cleaning, a greater emphasis on precision may be warranted, even at the ex-
245 pense of longer processing times.

246 We compared our approach using the same seismic dataset with DD, that we con-
247 sider as the reference for the state of the art. For this task, DD underwent compre-
248 hensive training for 400 epochs, while our model completed its training in just 150 epochs.
249 This difference was due to our model’s learning dynamics and efficiency. We initiated
250 our model’s training with a learning rate of 1e-3 and employed a scheduler to reduce this
251 rate gradually, ensuring controlled and stable convergence.

2.3.3 Inference with Direct and Sampling Reconstruction

252 Cold diffusion models involve distinct methods to reconstruct the signal including
253 the adoption of direct or sampling reconstruction. These methods represent approaches
254 within the framework of diffusion models, each with unique operational mechanisms and
255 implications for model performance. Understanding the nuances of these methods is cru-
256 cial for comprehending the overall efficacy and application potential of diffusion mod-
257 els.
258

259 For the range of configurations used in training our models ($T = [20, 100, 300]$),
260 we applied these configurations to both direct and sampling reconstruction. In the con-
261 text of diffusion models, the distinction between ‘direct’ and ‘sampling’ approaches is
262 pronounced, marked by their differing operational mechanisms.

263 The ‘**direct**’ method involves applying the reverse process using the U-Net archi-
264 tecture to transition from a specific timestep t_n directly to zero. Conversely, the ‘**sam-**
265 **pling**’ method incrementally applies this reverse transition from a specific timestep t_n
266 to zero, but crucially, it traverses through all intermediate timesteps t_i , where $i \in [n-$
267 $1, 0]$. This results in applying the U-Net architecture multiple times (n).

268 A key aspect of the cold diffusion paradigm is evaluating the effectiveness of the
269 sampling procedure, which is hypothesized to outperform the direct approach. If the di-
270 rect method, particularly using U-Net alone, yields comparable results, it would call into
271 question the necessity of the complex training infrastructure typically associated with
272 diffusion models. We provide a detailed comparison between the direct and sampling meth-
273 ods in section 4.

2.3.4 Metrics

274 For enhanced clarity, we define here the metrics used in our study now and then
275 in Section 4 we provide a detailed commentary on the results.
276

- 277 1. **Signal to Noise Ratio (SNR)** is a measure used to compare the level of a sig-
278 nal (earthquake in this case) to the level of background noise. A higher SNR in-
279 dicates that the seismic signal stands out clearly from the background noise, fa-
280 cilitating accurate analysis and interpretation. We defined SNR as in (Zhu et al.,
281 2019):

$$10 \log_{10} \frac{\sigma_{signal}}{\sigma_{noise}}.$$

- 282 where σ_{noise} and σ_{signal} are the standard deviation of waveforms before and af-
283 ter the P arrival, respectively.
- 284 2. **Cross-correlation** is a widely used measure of similarity between two signals.
285 We compute the zero-lag cross-correlation (CC) between the recorded earthquake
286 signals (before noise is added) that represents our ground truth x_0 and the denoised
287 ones to evaluate the performance of the different models in reconstructing the recorded
288 waveform.
- 289 3. To evaluate the **picking** performances of the proposed method, we applied the
290 deep learning phase picker PhaseNet (Zhu & Beroza, 2019) to the waveforms and
291 compared the retrieved arrival times with the labeled picked phases of the cata-
292 log ($\sim 70\%$ of manually picked and $\sim 30\%$ of automatic picked). In this way we
293 can assess the impact of the denoiser on P and S arrival determination, the ac-
294 curacy of which enables the calculation of a well constrained location.

3 Data Sources and Selection

In our study, we focus on a subset extracted from the STanford EArthquake Dataset (STEAD) (Mousavi et al., 2019). This section is dedicated to elucidating the composition of the subset, detailing the following components:

1. We selected specific seismic stations to gather earthquakes and others for noise, with some overlap, providing a clear trace of the data’s origin for our analysis (Figure 4).
2. The distribution of seismic events across the globe (Figure 4) is mapped out, with these events sorted into training, validation, and test sets. This classification helps us to assess the model’s effectiveness and its generalizability across different regions.
3. We applied constraints to the dataset, including the magnitude and proximity to the seismic stations.

STEAD features a significantly larger number of stations for earthquake data compared to those used for noise. Moreover, the majority of these stations are concentrated within the U.S. territory. In our study, we utilize a ratio of (1786/2613) stations for the extraction of earthquake data, representing a fraction of the total available. For seismic noise, we have selected a subset corresponding to 306 stations dedicated to noise recording.

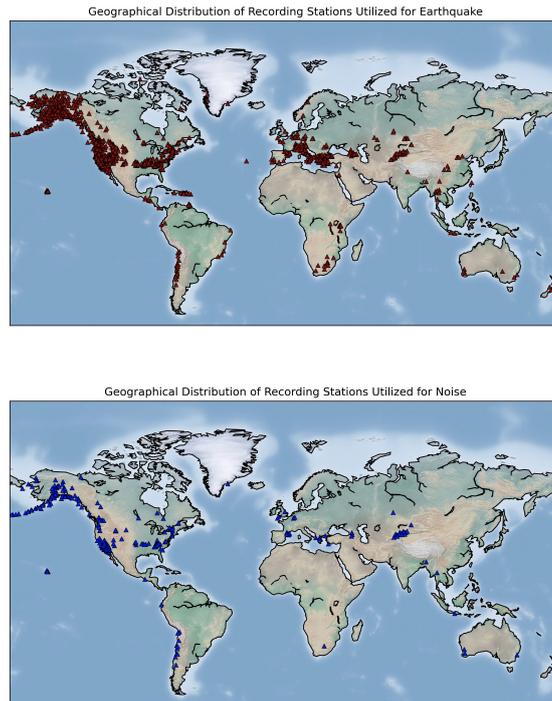


Figure 4. The maps show the subset of stations of the STEAD (STanford EArthquake Dataset) used for the recorded earthquake signal (upper) and the recorded noise (bottom).

Throughout our analysis, we consistently sample seismic traces of 30-second durations, based on the following criteria: magnitude > 2 , earthquake-station distance < 100

315 km, and P-wave arrival after 7 seconds. Figure 5 shows the frequency-magnitude statistics for our data set.
 316

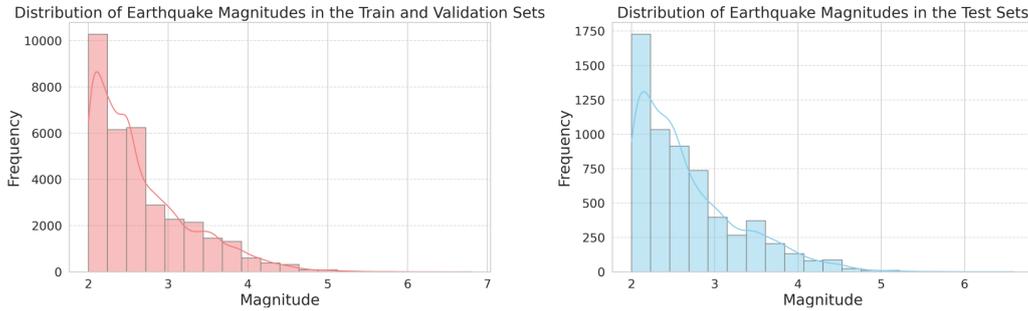


Figure 5. The histograms illustrate the frequency distribution of earthquake magnitudes within our dataset, with the left panel representing the training and validation sets and the right panel the test set.

317 We chose an inclusive approach for training, leveraging the full spectrum of avail-
 318 able data, without any SNR selection criteria. While this might seem disadvantageous
 319 initially, a model that performs well under these conditions can be versatile across var-
 320 ious scenarios. For researchers looking to retrain this model on their datasets, especially
 321 when specific datasets are limited, it may be advantageous not to put restrictive filters
 322 such as SNR.

323 Our dataset was divided into training (30491 traces), validation (3441 traces), and
 324 test (5994 traces) as illustrated in Figure 6. Such a division in machine learning ensures
 325 model reliability and generalizability. The training set aids the model’s primary learn-
 326 ing, the validation set is used for hyperparameter adjustments, and the test set objec-
 327 tively evaluates the model’s performance on unseen data.

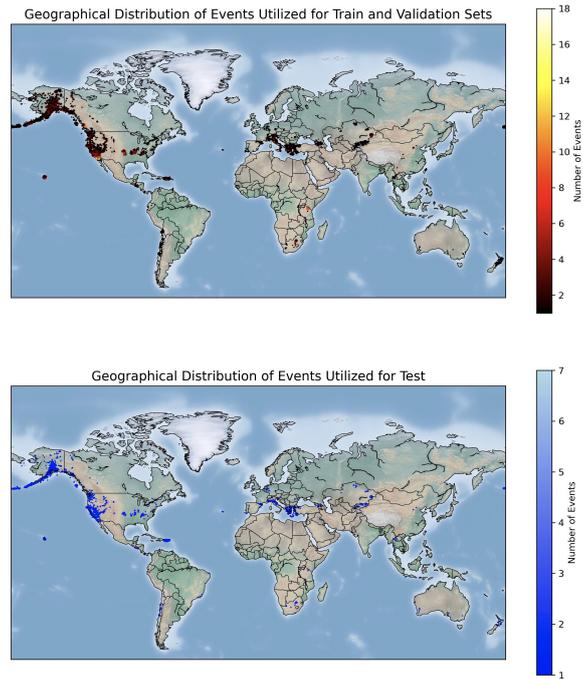


Figure 6. The image presents two maps of the geographical distribution of seismic events used in our study, with the upper map illustrating the events for the training and validation sets marked in red, and the bottom map showing the events for the test set in blue. The color intensity on each map corresponds to the number of events, with darker shades indicating a lower concentration of events in that location.

328 For more details on the specific train, validation and test configurations, please refer
329 to our GitHub repository at the following link:

330 <https://github.com/Daniele-Trappolini/Diffusion-Model-for-Earthquake>.

331 4 Results

332 In the following we present our results and discuss the validity of our model by adopt-
 333 ing quantitative and qualitative categories. The metrics used for each are provided sec-
 334 tion 2.3.4.

335 4.1 Quantitative Results

336 4.1.1 Signal to Noise Ratio (SNR)

337 A comparison of the SNR metric for the denoised waveforms obtained with differ-
 338 ent models and configurations is shown in Fig. 7. Note that Figure 7 includes the same
 339 metric for the original earthquake signals (labeled "earthquake") and those with added
 340 noise (labeled "eqk + noise"). The latter are the inputs to the denoiser algorithm. The
 341 performances of the different models appear aligned, with DD differing by a slightly lower
 median but greater variability in output SNR. In Fig. 8 we classified the noisy obser-

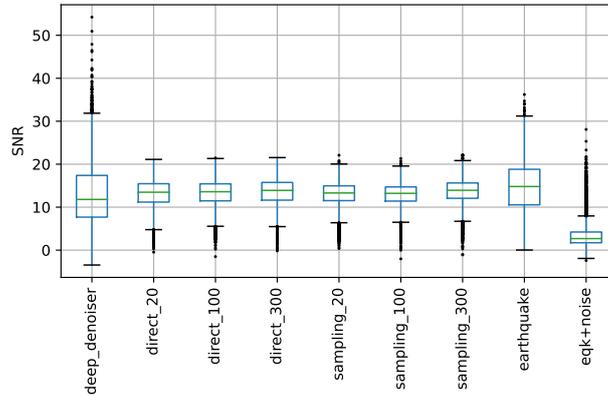


Figure 7. SNR comparisons using box-plots for various model configurations applied to the test set. The original signals (earthquake) and the ones with added noise (eqk+noise) have respectively the higher and lower SNR, as expected. The different denoising models appear overall aligned, with direct and sampling showing slightly higher median values and tighter distributions with respect to DD.

342 variations as a function of the SNR before denoising to highlight the effectiveness of our
 343 models in cleaning the seismic traces. The performance of our CDiffSD are consistently
 344 superior with respect to DD in low SNR scenarios. This aspect is crucial, given that low
 345 SNR conditions correspond to more complex and heavily noisy seismic traces precisely
 346 where an effective denoising solution is most needed. The high-quality performance of
 347 our model in these low SNR environments is demonstrated in Fig. 8. We note in partic-
 348 ular model reliability and efficacy in extracting correct signals from noisy data. This
 349 proficiency is important in real-world seismological applications, especially for discov-
 350 ering lower magnitude earthquakes often hidden in the noise.
 351

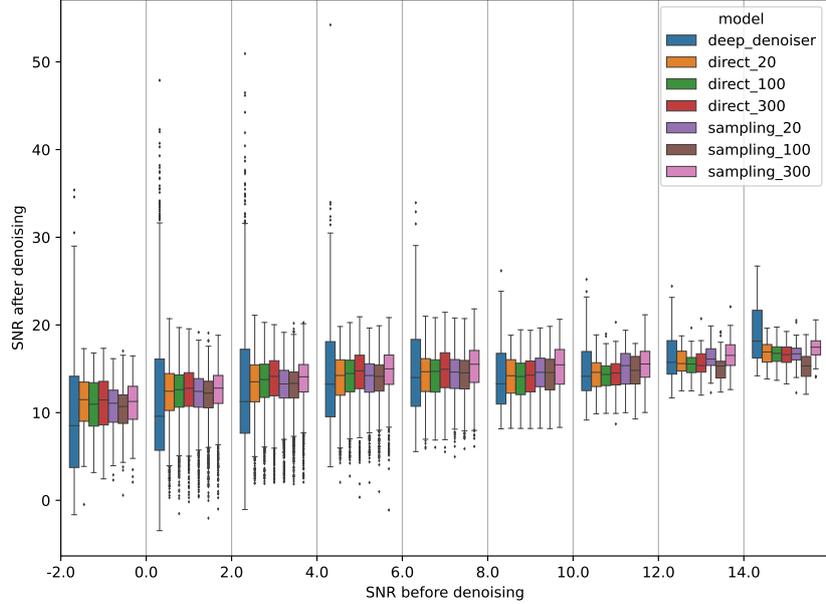


Figure 8. Distributions of SNR values of denoised waveforms for different ranges of input SNR. The SNR statistics after denoising are computed on 2dB wide ranges of input SNR. CDiffSD models show higher performances in low SNR scenarios, while DD is superior for the higher SNR signals. We study the range: $SNR < -2.0 \cup SNR > 16.0$, which covers 99% of real data. Solid bars within each model (color) show the median value.

352 While the cold diffusion approach excels in low SNR scenarios, the binary mask-
 353 based method DD exhibits greater variability and tends to perform better in higher SNR
 354 conditions, benefiting from its ability to provide a clear-cut signal delineation (Fig. 7 and
 355 Fig. 8). In particular, DD shows improved performance when the input SNR is higher
 356 than ~ 14 and is get worse at lower input SNR while our models remain consistently
 357 effective for a large range of input SNR. An example of high input SNR conditions can
 358 be found in the Supporting Information.

359 *4.1.2 Cross Correlation*

360 We evaluate the similarity between original signals and denoised signals, by show-
 361 ing the statistics of the computed CC values, in Fig. 9. A higher CC indicates a greater
 362 similarity between the denoised trace and the original signal. In this figure, we see that
 363 all CDiffSD models show similar performance and they are all consistently higher than
 364 DD. To better highlight the variability of CC values obtained from the different traces
 365 of the test set, in Fig. 10 we show the distribution of CC values between denoised and
 366 original traces as a function of CC of the noisy traces with original signals (x axis), that
 367 is, CC of traces before denoising is applied. The performances for both direct and sam-
 368 pling are higher than DD for every considered range of CC before denoising.

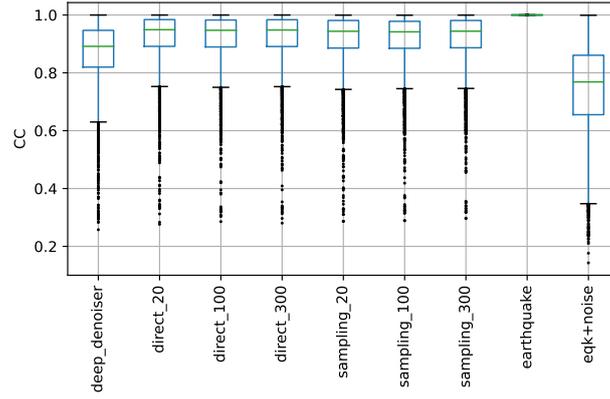


Figure 9. Cross-correlation (CC) comparisons for various model configurations applied to the test set. Higher CC values indicate greater similarity between the denoised trace and the original signal. All CDiffSD models show similar performance and that they are consistently higher than DD.

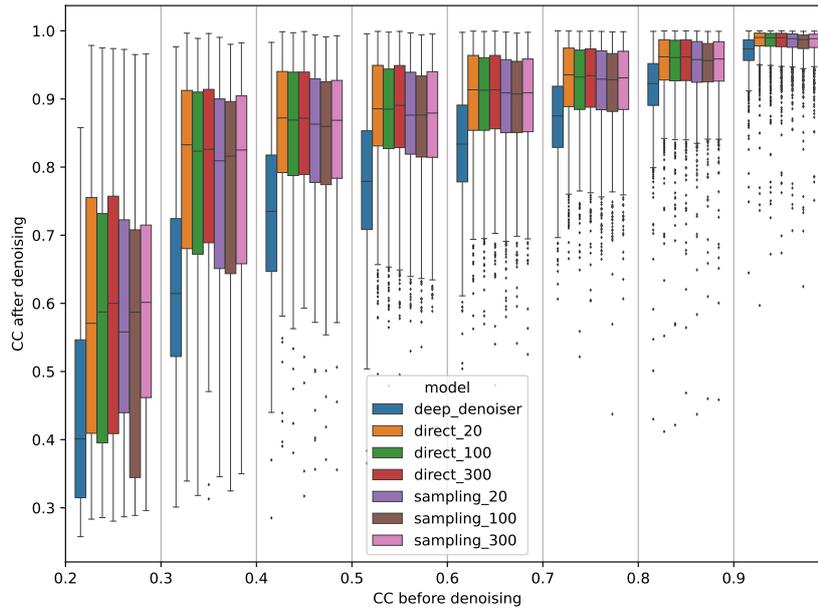


Figure 10. Distribution of CC values between original and denoised signals (y axis) as a function of CC before application of denoising. The statistics are computed for ranges of 0.1. CDiffSD models show better performances with respect to DD for all the ranges considered. The difference is more noticeable especially at low CC values before denoising.

369 For each model considered we see better performance, with higher values of CC af-
 370 ter denoising (Fig. 10). Another noteworthy aspect is that at higher noise levels, thus
 371 lower CC before denoising (values from 0.2 to 0.3), models with $T = 300$ outperform their
 372 counterparts. As expected, these performance disparities tend to converge with an in-

373 crease in CC before denoising, corresponding to a relative reduction in noise compared to the signal.
 374

375 **4.1.3 Phase arrival picks**

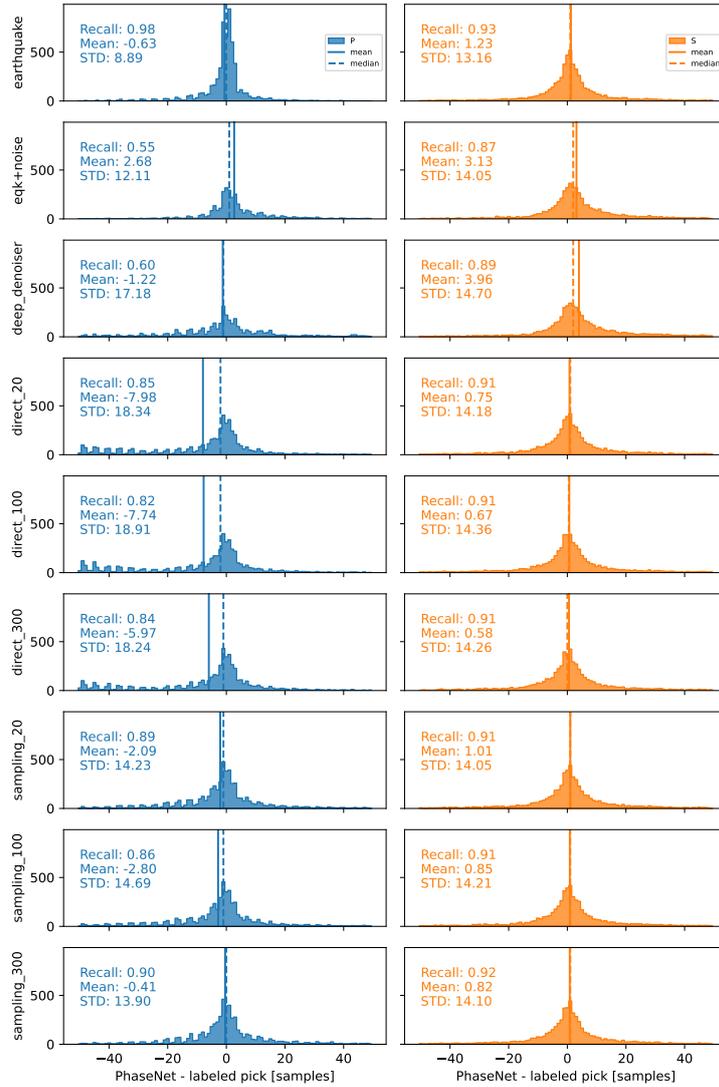


Figure 11. The histograms display the distributions of P-wave (blue) and S-wave (orange) arrival time differences between automated PhaseNet detections and label picks (in samples). The results obtained using the original seismograms, eqk+noise and DD are shown for comparison in the first, second and third row, respectively. The remaining rows show the results for different CDiffSD models applied to the same data subset, offering insights into the accuracy of wave arrival time detection by each model. Central tendency metrics, such as mean and median, are indicated in these histograms, highlighting any potential skewness in the distribution towards either early or late picks for both P and S waves.

376 The histograms in Fig. 11 provide a visual representation of the efficacy of differ-
377 ent seismic signal denoising methods — "direct", "sampling", and DD — in retrieving
378 a signal and preserve P- and S-wave onsets. The accuracy of automated P and S-wave
379 arrival time picks by PhaseNet is compared to label picks. The histograms are organized
380 by method and parameter variations, displaying the distribution of arrival time discrep-
381 ancies measured in samples.

382 In the case of earthquake (i.e., no noise added, top histogram), the P-wave pick dif-
383 ference distribution exhibits spreads that are narrower than those of the S-wave and this
384 is in full agreement with the expected behavior.

385 When noise is introduced, the pick difference distributions for P-waves and S-waves
386 tend to converge towards a more similar pattern. This convergence can be attributed to
387 the primary impact of noise on P-waves, owing to their lower amplitude compared to S-
388 waves. As a result, the performance with added noise on P waves detection is much more
389 degraded than on S waves detection with the same level of noise because P-waves have
390 also smaller amplitudes. This observation is further supported by the recall values for
391 S waves, which remain greater than 0.85 not only for all the denoising methods, but also
392 for the noisy traces (earthquake + noise). In contrast, the recall rate for P-waves is con-
393 sistentlly lowered by the presence of noise (Fig. 12). For these reasons we focus our anal-
394 ysis on P-wave picks.

395 As seen in Fig. 11 the distribution of the "direct" methods show pronounced neg-
396 ative skews, with mean values far from 0. This indicates a tendency of PhaseNet to pick
397 P-waves slightly before the labeled picks for the waveforms denoised with "direct" meth-
398 ods. The reason of this behavior is most likely to be attributed to noise remaining in the
399 denoised traces processed with the "direct method". This in turn can mislead PhaseNet
400 to an early detection (see the "direct" example in Fig. 13). This tendency, however, is
401 mitigated completely by the CDiffSD "sampling" method, as shown in Fig. 13. In par-
402 ticular, we see that the "sampling" methods display recall rates that are consistently high
403 for both P and S, especially the 300 configuration, indicating a good denoising perfor-
404 mance and the ability to recover the labeled phases.

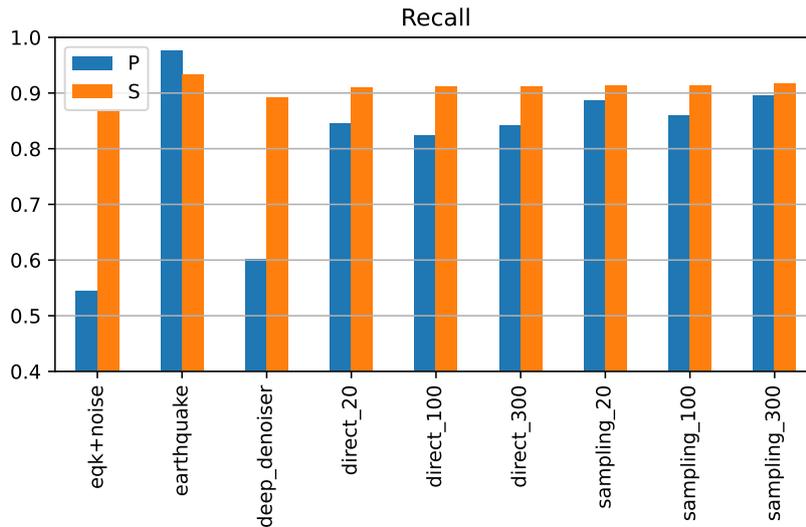


Figure 12. Comparison of recall rates for P and S waves between the different methods within a fixed window of 50 samples. S-waves recall rates are aligned almost for all models, indicating that the noise level is not enough to affect the S-waves because of the greater amplitude. P-waves recall rates instead show significant disparities between the DD approach and other methods, suggesting a lower performance of DD in preserving the P onset in these cases. The 'sampling-300' method is confirmed as the one with better performances.

405 From the comparison of the results obtained with the "direct", "sampling", and
 406 DD methods, it is evident that each method influences the automated pick accuracy dif-
 407 ferently. The "sampling" method, particularly at higher parameter settings, demonstrates
 408 a notable alignment with label picks, suggesting its superiority in mitigating noise and
 409 enhancing the precision of automated picking systems. It is also noteworthy that the re-
 410 call values for P-waves shown in Fig. 12 are higher than DD for both "sampling" and
 411 "direct" methods, which suggests that in these cases DD does not preserve accurate P-
 412 wave onsets.

413 4.2 Qualitative Results

414 Qualitative factors are useful as side-by-side comparisons between the outcomes
 415 of different approaches. In the supplementary materials, we delve deeper into the anal-
 416 ysis of seismic traces, examining the impact of denoising on picking. This includes a thor-
 417 ough examination of both the strengths and limitations of our model. We highlight in-
 418 stances where our model excels in denoising, as well as situations in which it does not
 419 perform optimally.

420 The examples below and those in the supplementary are organized with the same
 421 layout: in the top panel we compare the noisy signal (grey) with the denoised signal (black);
 422 in the middle panel we compare the original signal (green) with the denoised signal (black);
 423 the bottom panel is a zoom on the P-wave arrival.

4.2.1 Qualitative Picker Analysis: Direct Vs Sampling

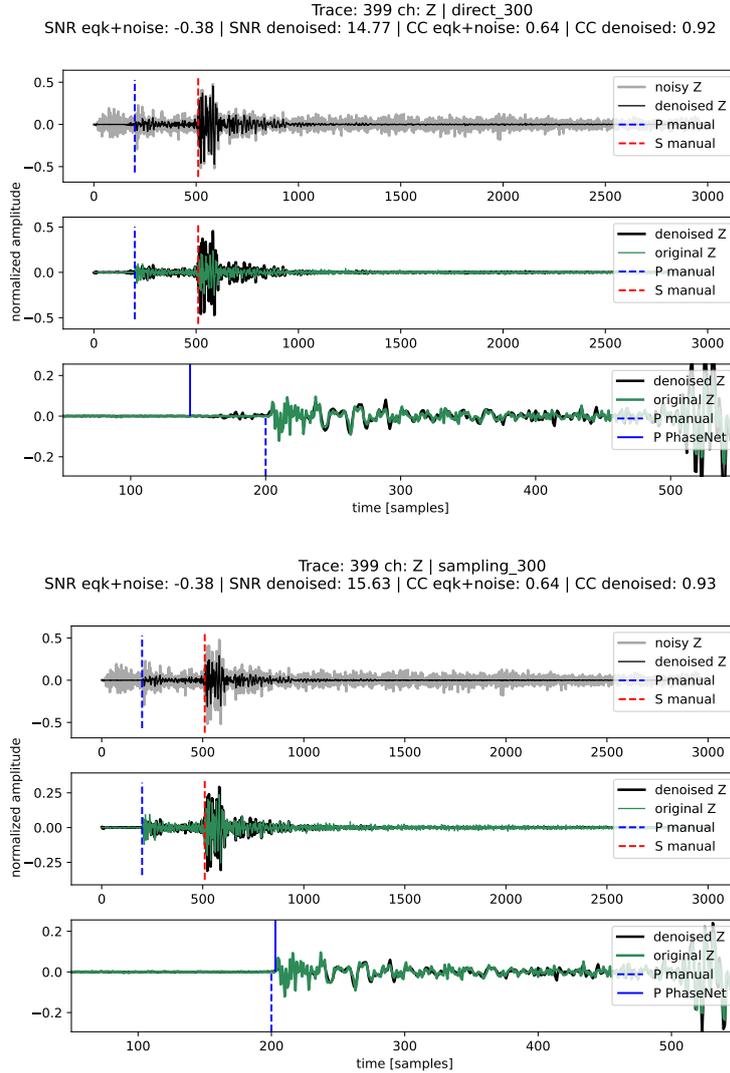


Figure 13. Comparison of a trace processed using 'direct_300' and 'sampling_300' methods. Notably, the 'direct_300' retains some of the noise preceding the P wave arrival, which is instead filtered out in the 'sampling_300' results. This noise before the P-wave retained with the the 'direct_300' method explains the tendency for this model to cause early picks (as seen in Fig. 11), as the residual noise can lead to earlier detections.

425 In the first example shown in Fig. 13 we compare the 'direct_300' and the 'sam-
 426 pling_300' methods. Here "sampling" method is found to be more effective than the "di-
 427 rect" method in denoising the seismic signal, and this is particularly evident from the
 428 middle and bottom panels, where the denoised signal in the "sampling" method match
 429 more closely the original signal. In contrast, the "direct" method shows more significant
 430 deviation from the original, especially before the P-wave arrival. This example is also
 431 useful because it provides insight into the tendency of the "direct" methods to cause spu-
 432 rious early P-picks. The direct method, in fact, retains some pre-arrival noise, which can
 433 trigger an early pick in automatic approaches such as PhaseNet. This is less of an issue

434 in the sampling method, as seen in the lower set of traces, where the denoised signal is
 435 cleaner, and the P-wave arrivals are closer to the labels. The implication for seismic pro-
 436 cessing is significant since the sampling method appears to produce cleaner signals and
 437 more accurate P-wave arrival times as a direct consequence. We note that this is crucial
 438 for various seismological applications such as earthquake location and tomographic
 439 imaging.

440 **4.2.2 Qualitative Picker Analysis: DD Vs Sampling**

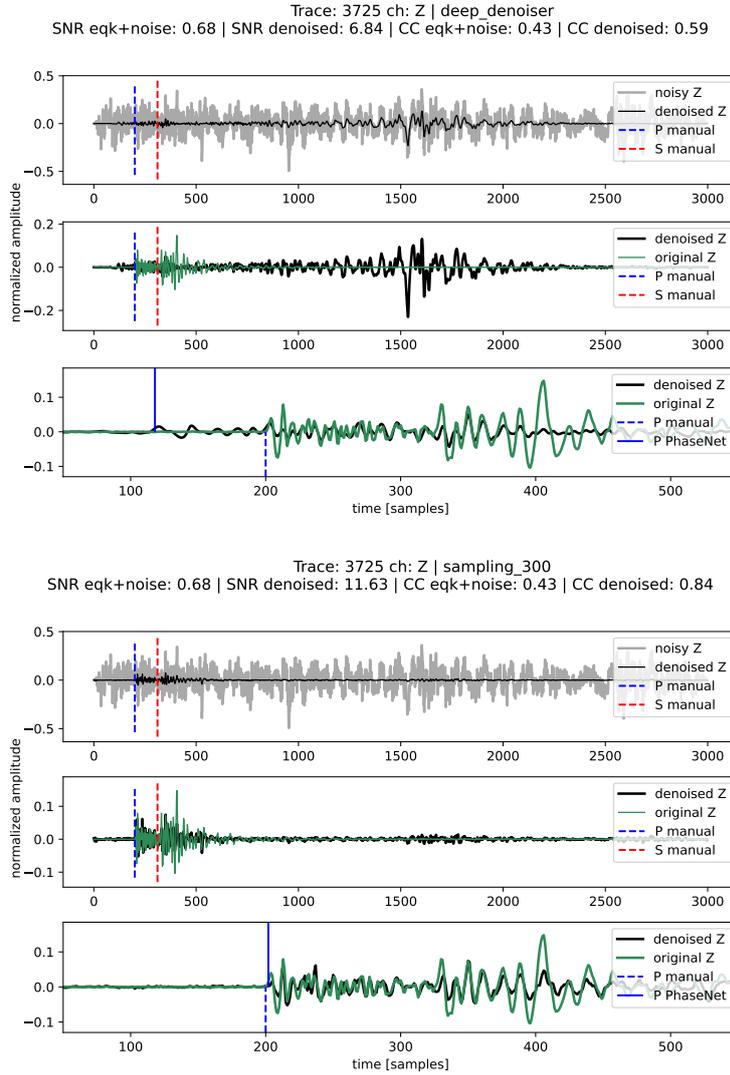


Figure 14. Comparison between a seismic trace processed with 'deep.denoiser' and 'sampling_300' methods. The 'sampling_300' method demonstrates a closer match to label phase picks and a more precise amplitude preservation, despite the substantial noise present in the original signal. DD also retains a high amplitude noise signal at around 1500 samples that 'sampling_300' manages to filter out almost completely.

441 Fig. 14 exemplifies the concepts previously discussed in Fig. 8, highlighting the per-
 442 formance of our model compared to that of the 'deep denoiser' in scenarios with very low
 443 Signal-to-Noise Ratio (SNR) before denoising. The figure demonstrates clearly how an
 444 extreme noise situation can lead to an error in phase picking for the 'deep denoiser', whereas
 445 the 'sampling' method is capable to reconstruct accurately the correct P wave arrival
 446 despite the presence of significant noise.

447 **4.2.3 Qualitative Amplitude Analysis: Direct Vs Sampling**

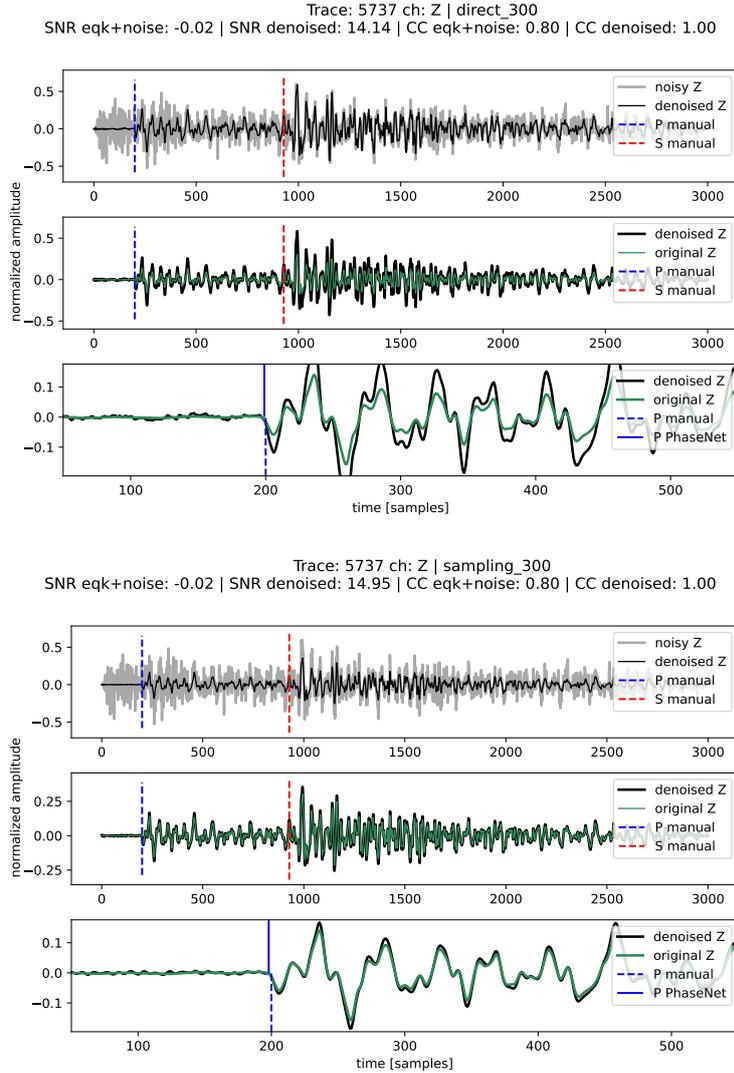


Figure 15. Comparison of a seismic trace processed with 'direct_300' and 'sampling_300' methods. It is particularly significant that the 'sampling_300' technique demonstrates an enhanced ability for amplitude reconstruction compared to the 'direct_300' method.

448 In Fig. 15 we show a comparative analysis of seismic signal denoising methods to
 449 investigate the importance of amplitude preservation. The Cold Diffusion Model employ-
 450 ing a sampling strategy ('sampling_300') demonstrates a superior performance in main-
 451 taining the amplitudes of the seismic signal. In practice, the denoised signal aligns more

452 accurately with the original waveform, preserving the integrity of the amplitude across
 453 the signal’s duration. This is particularly evident in the detailed zoomed-in analysis, where
 454 the ‘sampling_300’ method displays remarkable congruence with the original signal, as
 455 evidenced by the minimal and consistent residuals. In contrast, the direct application
 456 of a U-Net model (‘direct_300’) displays a slight but discernible attenuation in ampli-
 457 tude, most noticeable in segments with higher amplitude peaks. The increased residu-
 458 als associated with the ‘direct_300’ method suggest a more significant alteration of the
 459 signal after the denoising process. Therefore, the Cold Diffusion Model with sampling
 460 stands out as the most effective method for seismic data denoising (amongst those tested
 461 here), especially where the preservation of amplitude is critically important.

462 5 Model assessment: Assessing the Impact of Exclusive Noise Input

463 In this section we aim to test the behaviour of the model in no-earthquake scenar-
 464 ios, i.e. with inputs containing only noise. This is done in order to verify whether the
 465 model doesn’t generates any artifacts in the absence of signal generating false earthquakes.

466 Cold Diffusion is based on the model’s ability to learn the broad data distribution
 467 during training, which generally includes a variety of seismic traces with different lev-
 468 els of noise. Therefore, the model should be able to generalize and identify traces that
 469 are entirely dominated by noise, even without direct exposure to specific types of earth-
 470 quake samples where there is no earthquake signal. Based on these assumptions, we seek
 471 to verify if our results align with the theoretical expectations.

472 We have used the entire noise test set as input, without combining it with the earth-
 473 quake data. Theoretically, with a perfect denoising, the expected output would be a trace
 474 composed exclusively of zeros, in the real context the trace should approach zero.

475 We applied the model without retraining, meaning the model’s weights have never
 476 been exposed to the absence of earthquake traces as ground truth. To assess the correct-
 477 ness of the output we set an amplitude threshold between ± 0.02 to decide whether the
 478 output could resemble a trace of zeros. The direct and sampling methods have correctly
 479 reconstructed the expected signal in 60.3% and 88.6% of cases, respectively. This dif-
 480 ferent performance highlights the sampling method’s superior capability in recognizing
 481 the absence of earthquake signals and adapting to it.

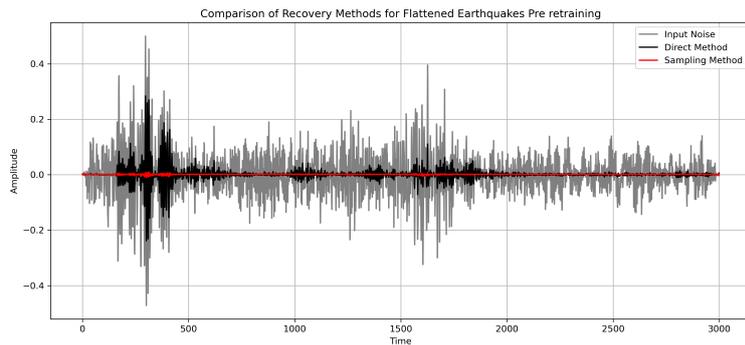


Figure 16. Example of the outputs of direct (in black) and sampling (in red) methods in case of a noise only input (in grey). No retraining is performed here, i.e. the models have never been exposed to zero-traces as ground truth for noise-only input. The direct method fails in recovering a zero-trace since it introduces artificial signals. In contrast, the sampling method reconstructs successfully an output that resembles a zero-trace.

482 Given the promising results just described, we further explored this scenario by re-
 483 training the model including no-signal traces as ground truth. We focused only on a sin-
 484 gle channel for this test and incorporated 3% of the entire training set with zeroed traces
 485 to represent the absence of seismic events. The results align with our expectations, in-
 486 dicating an improvement in performance in the presence of noise alone. Specifically, the
 487 cases where zero traces are retrieved increases to 68.2% and 90.5% for direct and sam-
 488 pling methods, respectively. The direct method exhibits a more substantial improvement,
 489 starting from a lower baseline performance, whereas the sampling method shows a smaller
 490 increase, likely due to its performance already approaching saturation.

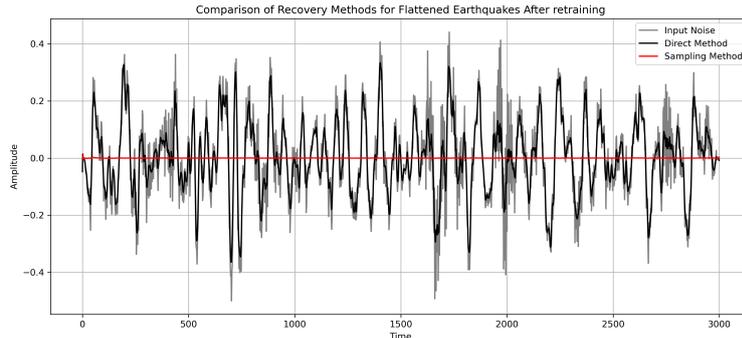


Figure 17. Example of the comparison between the sampling method (in red) and the direct method (in black), the input (in gray) for both methods is only noise. In this case the models have been retrained with zero-traces as ground truth for noise-only traces. The sampling method succeeds in reconstructing a zero-trace. On the other hand, the direct method outputs noise, indicating a less accurate reconstruction in this scenario.

491 Regarding the results post-retraining, it should be noted that the output trace of
 492 the sampling method shown in Figure 17 is indeed close to the expected zero-trace. On
 493 the contrary, low amplitude noise was still present in the output of the non retrained-
 494 case shown in Figure 16. This highlights the importance of including flat traces during
 495 the training.

496 In this evaluation of the CDiffSD on these cases comprised solely of noise, we proved that
 497 it is not imperative to include such examples in training to accurately discern between
 498 noise and genuine seismic signals. However, including these kind of signals in training,
 499 improves the capability of effectively identifying traces that are comprised solely of noise.
 500

501 6 Conclusion

502 Our study has demonstrated promising results in pick accuracy (4.1.3), Signal-to-
 503 Noise Ratio (SNR) enhancement (4.1.1), and Cross-Correlation metrics (4.1.2), thus af-
 504 firming the validity of cold diffusion denoising for seismological applications. In addition
 505 to these achievements, it is important to emphasize that, despite SNR and Cross-Correlation
 506 metrics aligning with other models, the sampling with $T = 300$ demonstrates its su-
 507 periority in practical, applied contexts, such as P-phase onset picking.

508 While SNR and Cross-Correlation are critical metrics for assessing the quality of
 509 the reconstructed signal, not every part of the seismic trace holds equal significance. In
 510 fact, the preservation of the integrity of the P- and S-wave arrivals is fundamental. As

511 highlighted in Section 4, the most effective model in this regard is the one utilizing sam-
 512 pling with $T = 300$. This model’s ability to maintain the aspects of the seismic trace,
 513 particularly the arrival times of these key waveforms, underscores its practical superi-
 514 ority in applied seismological contexts.

515 The findings discussed in Section 4, while serving as a good base, should be regarded,
 516 however, as a preliminary step towards addressing a broader spectrum of open questions
 517 and potential model enhancements.

518 A significant direction for future advancement lies in the broadening of our dataset.
 519 Our initial explorations aimed to establish the feasibility of these methods. Moving for-
 520 ward we could potentially develop a more generalized model by retraining on the full STEAD
 521 and INSTANCE (Michelini et al., 2021) datasets, encompassing collectively several mil-
 522 lion traces compared to the $\sim 40k$ traces used in this study. This expanded model would
 523 be capable of effectively treating noise in a wide range of seismological contexts with-
 524 out the need for further retraining, thus significantly boosting its applicability and ro-
 525 bustness across diverse seismic scenarios.

526 In conclusion, the model presented exhibits significant potential for enhancing seis-
 527 mic traces, facilitating more accurate onset picking of P- and S-waves. Moreover, it holds
 528 promise for extracting earthquakes from noise—events that may have eluded human de-
 529 tection or other approaches. Such capability could contribute to expanding seismic cat-
 530 alogs. While further refinements are conceivable, this method, which is borrowed from
 531 speech enhancement tasks, has proven its validity in the intricate domain of seismolog-
 532 ical analysis. This cross-disciplinary innovation underscores the model’s versatility and
 533 suggests broader applicability in extracting and analyzing subtle seismic signals.

534 Acronyms

535	CC Cross Correlation
536	CDiffSD Cold Diffusion Model for seismic denoising
537	DAS Distributed Acoustic Sensing
538	DD Deep Denoiser
539	DL Deep Learning
540	DM Diffusion Model
541	DPRNN Dual-Path Recurrent Neural Network
542	E East-West
543	eqk Earthquake
544	ERC European Research Council
545	GAN Generative Adversarial Network
546	GELU Gaussian Error Linear Unit
547	ICA Independent Component Analysis
548	INGV Istituto Nazionale di Geofisica e Vulcanologia
549	INSTANCE Italian Seismic Dataset For Machine Learning
550	MUSIC MUltiple SIgnal Classification
551	N North-South
552	NRF Noise Reduce Factor
553	ResNet Residual Neural Network
554	SNR Signal to Noise Ratio
555	STEAD STanford EArthquake Dataset
556	STFT Short-Time Fourier Transform
557	VAE Variational Autoencoder
558	Z Vertical

559 **Open Research Section**

560 The STEAD (Mousavi et al., 2019)(Seismological Tools for Earthquake Analysis
561 and Detection) dataset is openly accessible at the following link: <https://github.com/smousavi05/STEAD>
562 or by utilizing ObsPy, a Python library for processing seismological data (for more in-
563 formation on ObsPy, refer to their official site: <https://docs.obspy.org/>).

564 To replicate the data accurately, it is necessary to apply the filters described in the
565 Section 3 to chunk2 of the STEAD dataset. Furthermore, specific data related to this
566 research will soon be made available on the GitHub repository at:

567 <https://github.com/Daniele-Trappolini/Diffusion-Model-for-Earthquake>.

568 **Acknowledgments**

569 This research was made possible by the generous support of both the Istituto Nazionale
570 di Geofisica e Vulcanologia (INGV) and the European Research Council (ERC) grant
571 835012 (TECTONIC). We acknowledge partial funding from the MUR PNRR FAIR (PE00000013)
572 project. Complementary funding was provided by the project INGV Pianeta Dinamico
573 2021 Tema 8 SOME (CUP D53J1900017001) funded by the Italian Ministry of Univer-
574 sity and Research “Fondo finalizzato al rilancio degli investimenti delle amministrazioni
575 centrali dello Stato e allo sviluppo del Paese, legge 145/2018”. The funding and resources
576 provided by these institutions have been instrumental in advancing the scope and depth
577 of our study. We extend our sincere gratitude to the INGV for their valuable contribu-
578 tions and to the ERC for their commitment to fostering scientific research and innova-
579 tion.

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